

Moistening processes before the convective initiation of Madden-Julian Oscillation events during the CINDY2011/DYNAMO period

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Backgrounds

- **Convection initiation** of MJO over the western **Indian Ocean (IO)** is one of most difficult tasks in understanding and prediction of MJO. Major target of **CINDY2011/DYNAMO** field program [October 2011-March 2012 (Yoneyama et al. 2013)].
- **moist static energy (MSE)/moisture** is a key variable that controls the MJO initiation, development, eastward propagation [e.g., Blade and Hartmann 1993; Kemball-Cook and Weare 2001; Malone 2009] . Observational studies evidenced the **buildup of moisture** associated with MJO (Kikuchi and Takayabu 2004; Kiladis et al. 2005; Yoneyama et al. 2008, 2013; Johnson and Ciesielski 2013)
- Triggering of MJO convection initiation vary with cases

Circumnavigation (tropical dynamics)

Knutson and Weickmann 1987; Wang and Li 2004; Matthews 2008; Straub 2013

Extratropical forcing (e.g., Rossby wave activity in the upper troposphere, cold surge event) Lau and Peng 1987; Hsu et al. 1990; Ray et al. 2009; Ray and Li 2013; Zhao et al. 2013; Wang et al. 2012

Moistening??

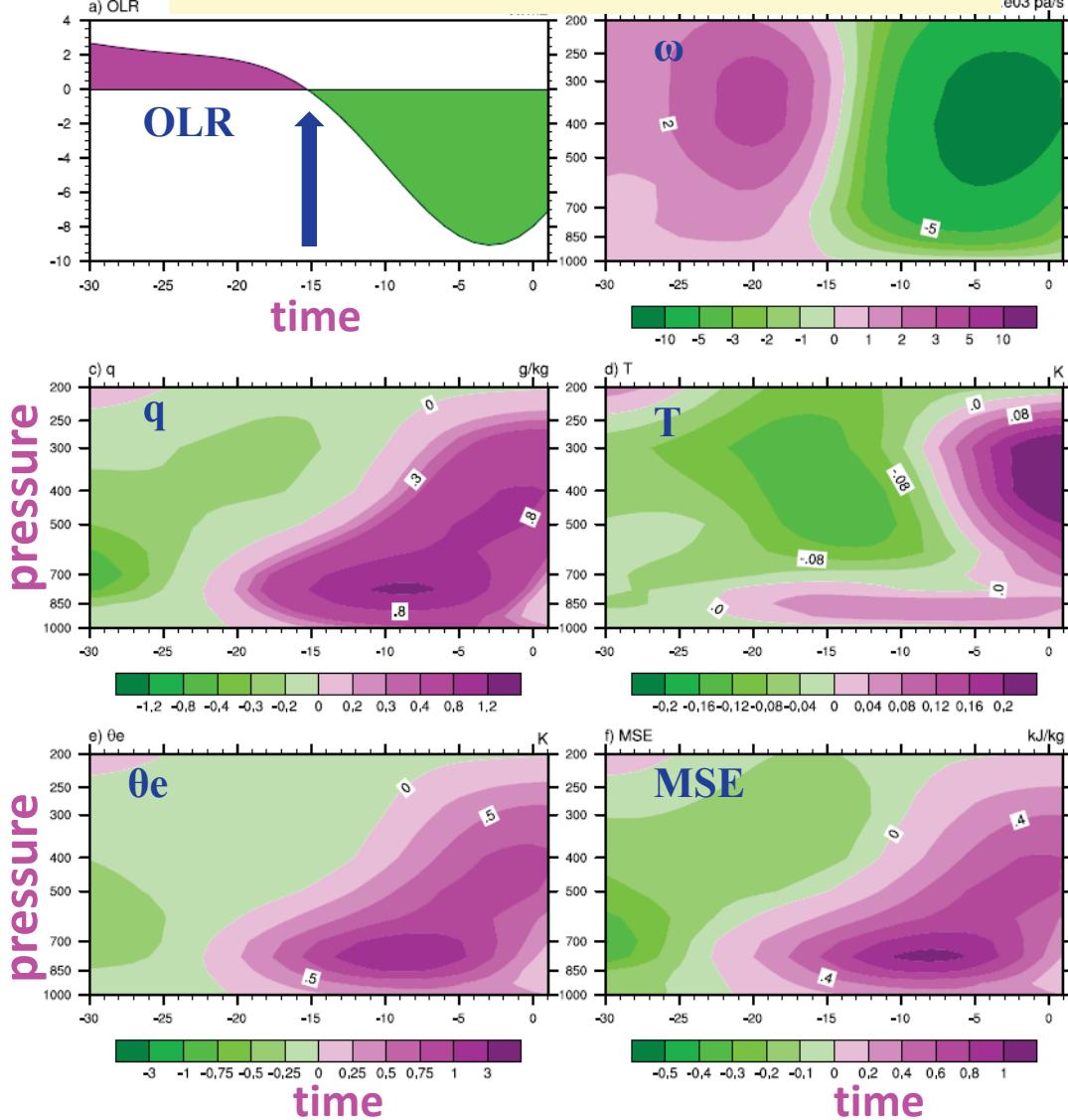
Goal:

- Gain insight into MJO convection initiation processes over the IO by case studies; MJO events during the CINDY2011/DYNAMO period
- A case study has an advantage in examining the roles of transient/local scale processes which tend to be smoothed out in statistical analysis of a large number of samples.
- Similarities and differences in the two MJO events (MJO-1 and MJO-2) will help understand the key processes.

Precursor signals prior to MJO initiation

Zhao et al. 2013

Time Evolution at (20S-0N, 50E-70E)



ERA-40, NOAA OLR (2.5deg)

Boreal winter **1982-2001**

MJO Initiation Region
($20^{\circ}\text{S} - 0^{\circ}\text{N}$, $50^{\circ}\text{E} - 70^{\circ}\text{E}$)

day -15 is regarded as the initiation date

Major precursor features:
Positive specific humidity & temperature anomalies in lower troposphere prior to the MJO initiation

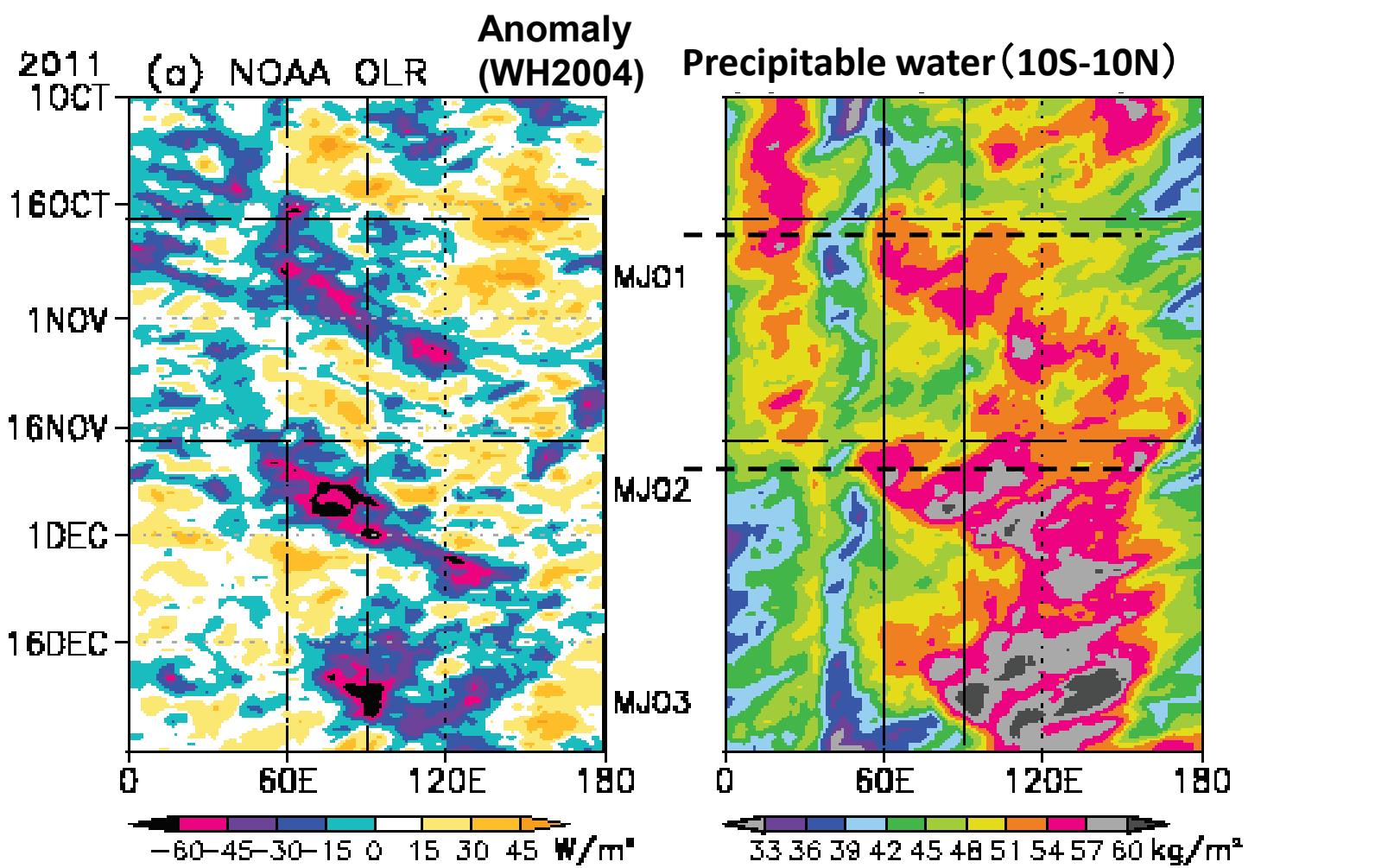
Increase of equivalent potential temperature θ_e & moist static energy MSE

Atmosphere becomes convectively more unstable

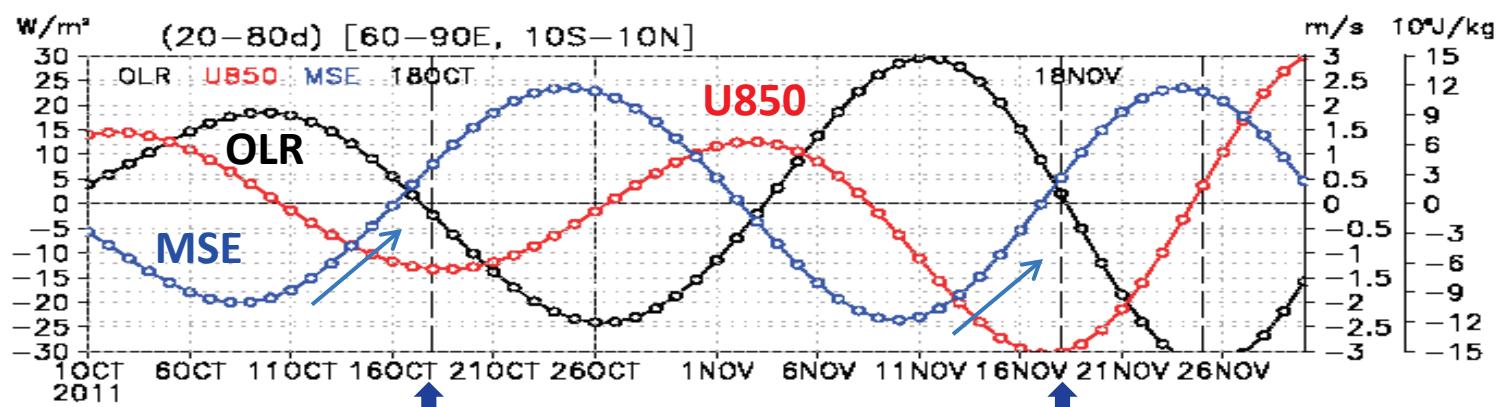
Data and methods

- **ERA-Interim** (1.0 x 1.0) and NOAA daily interpolated Outgoing Long-wave Radiation (OLR) (2.5 x 2.5) for the period of August 2011–January 2012 were used.
- Daily-mean variables were separated into low-frequency basic state (>80 day), intraseasonal (20–80 day), and high-frequency (<20 day) components by Lanczos filtering.
- **Moisture and temperature budgets** for the intraseasonal tendencies were diagnosed following Zhao et al. (2013).

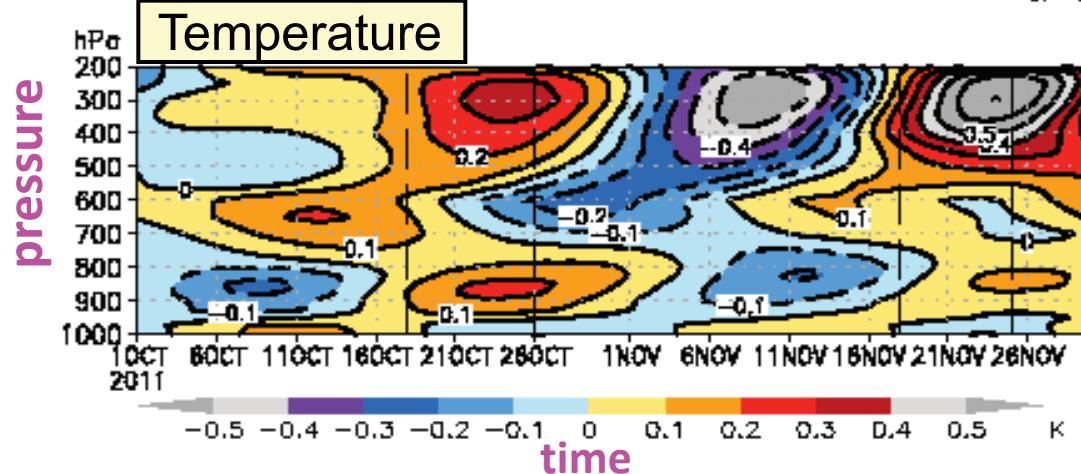
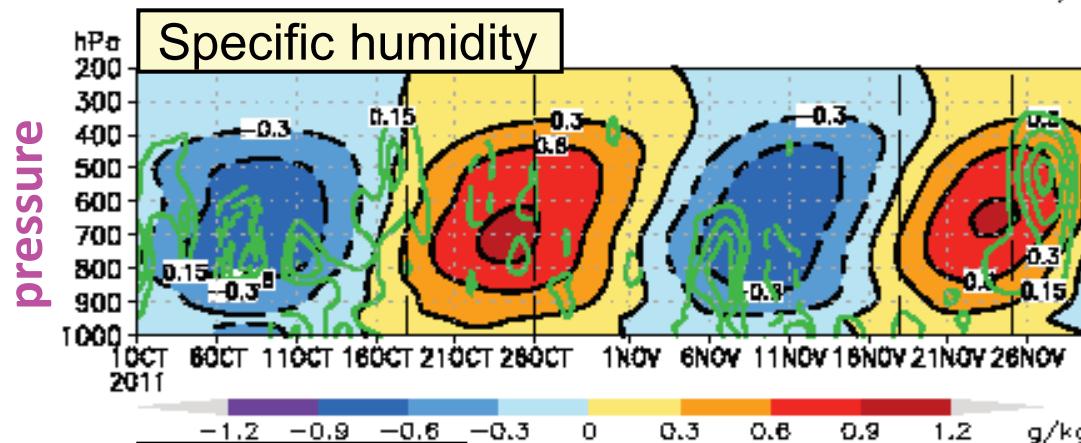
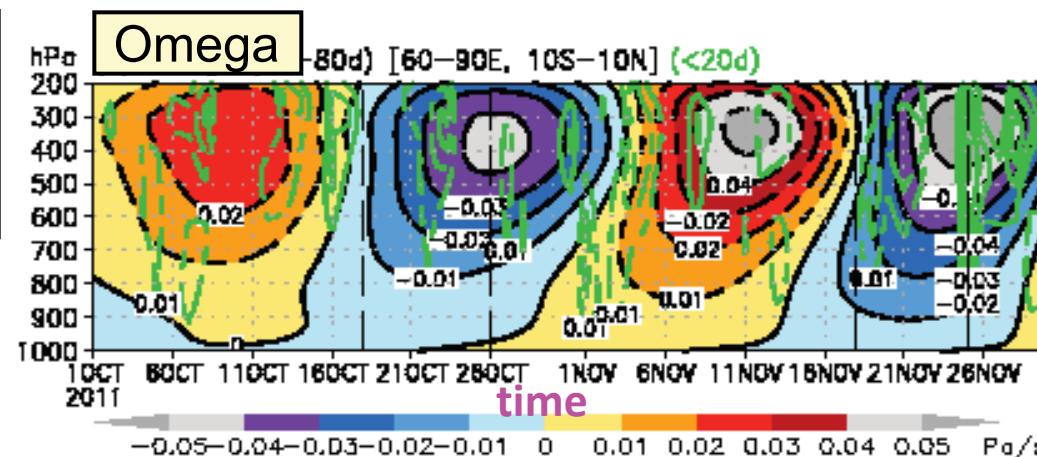
$$\left(\frac{\partial q}{\partial t} \right) \neq -V \nabla q' + -\omega \frac{\partial q}{\partial p}' + -\frac{Q_2}{L}, \quad \text{prime: 20-80 day (ISO)}$$



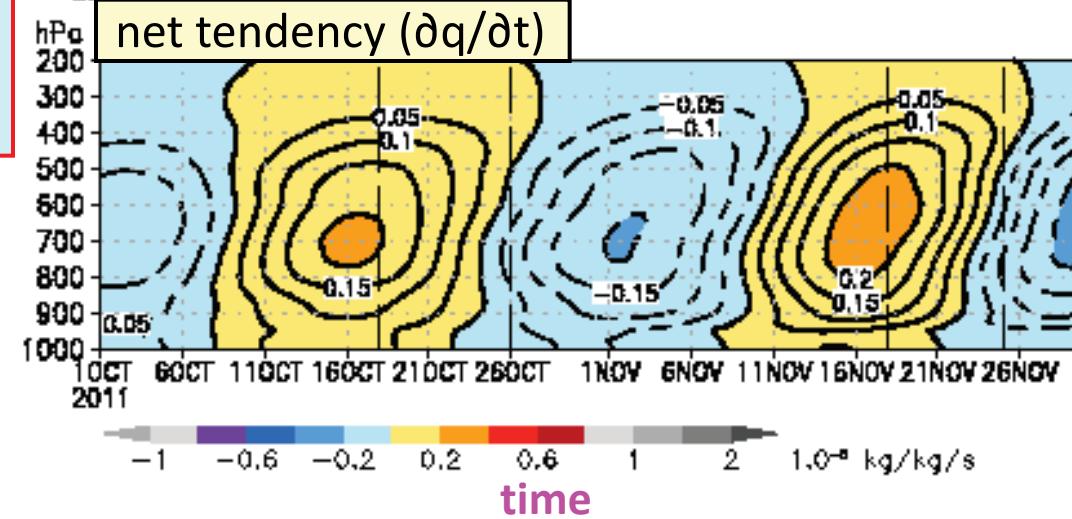
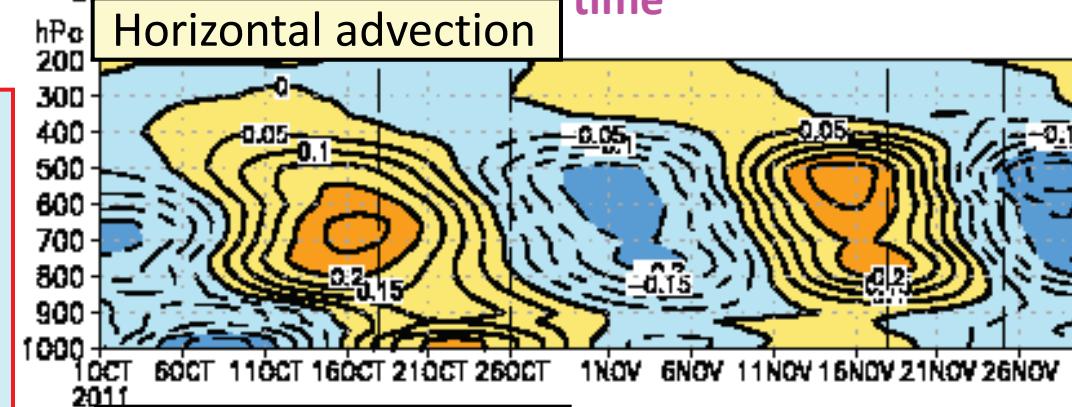
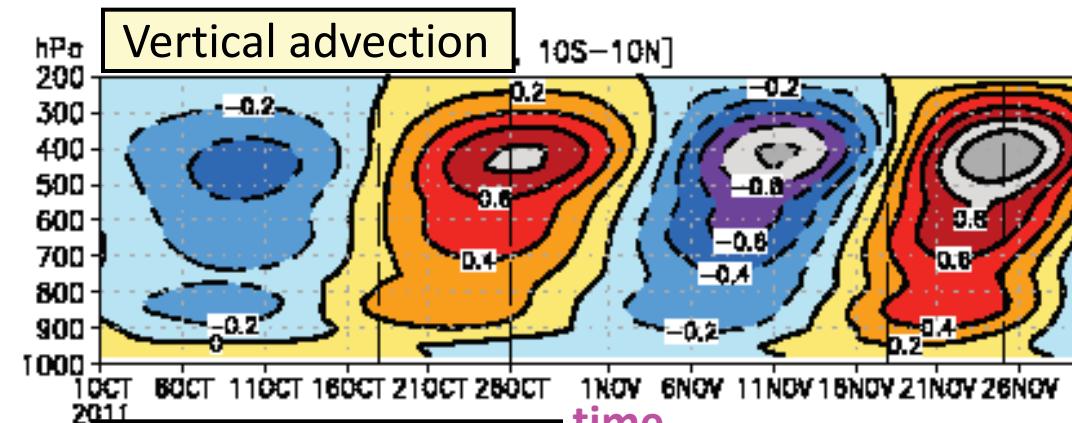
**MSE start
Increasing
7-10 days
before
OLR<0**



ERA-Interim
20-80d filtered
60-90E, 10S-10N



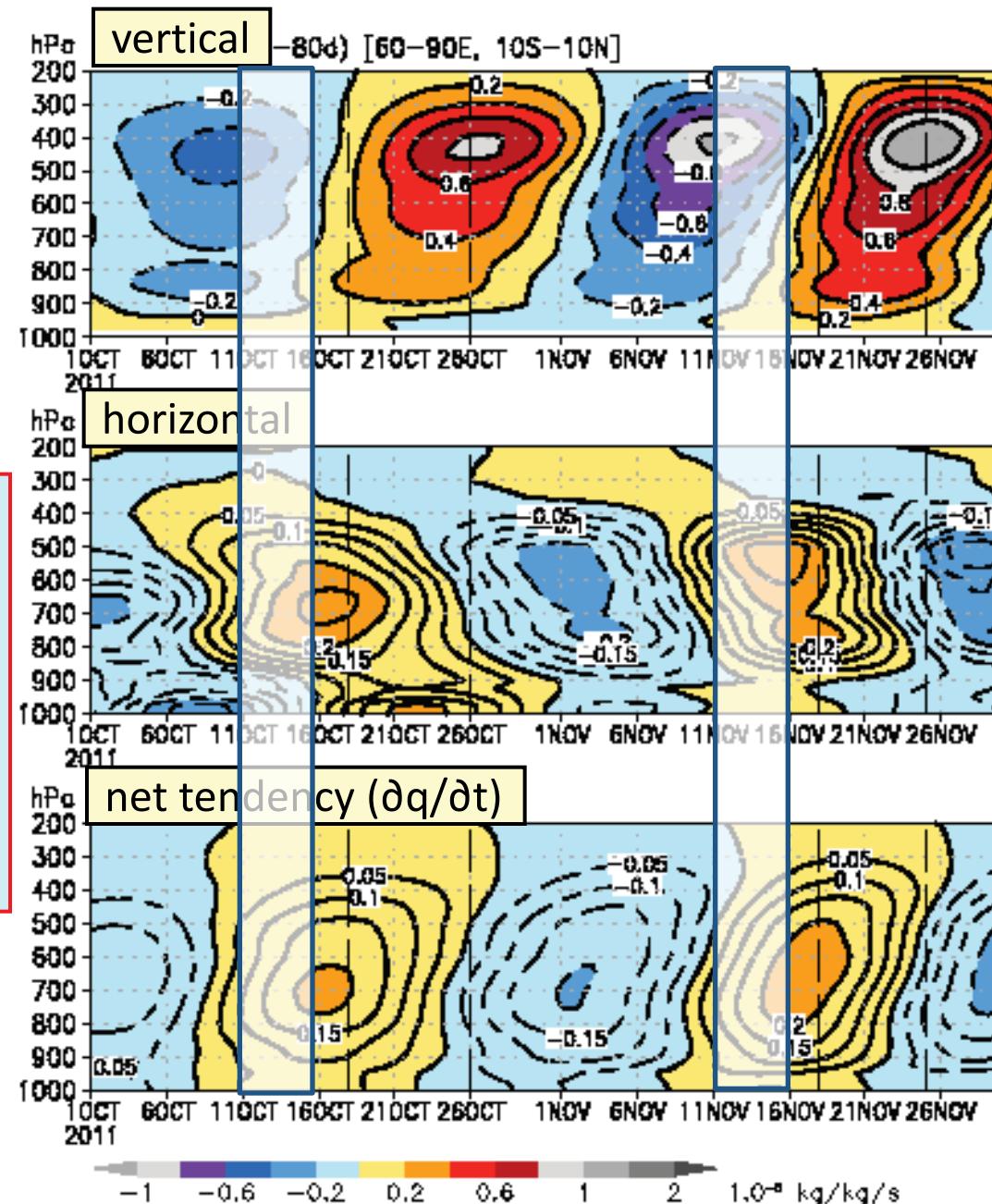
ERA-Interim
Moisture
advection
20-80d filtered
60-90E, 10S-10N



Horizontal advection
Preceded major
Moistening by
Vertical advection

Magnitude
comparable to
net tendency

ERA-Interim
Moisture
advection
20-80d filtered
60-90E, 10S-10N

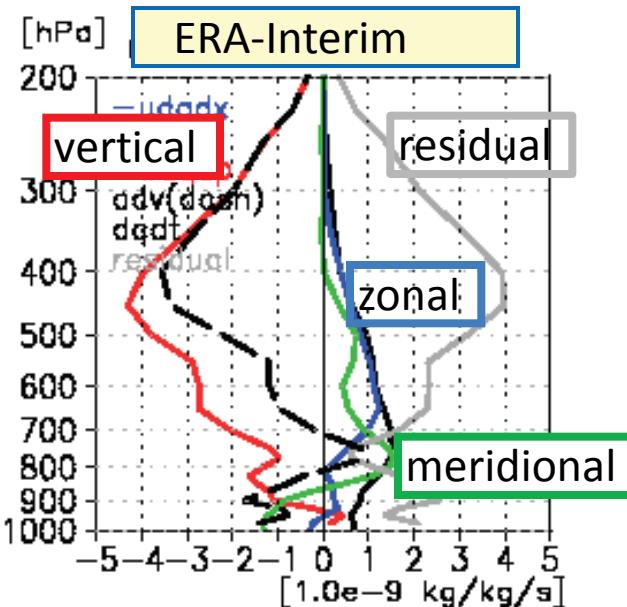


Horizontal advection
Preceded major
Moistening by
Vertical advection

Magnitude
comparable to
net tendency

Moisture budgets

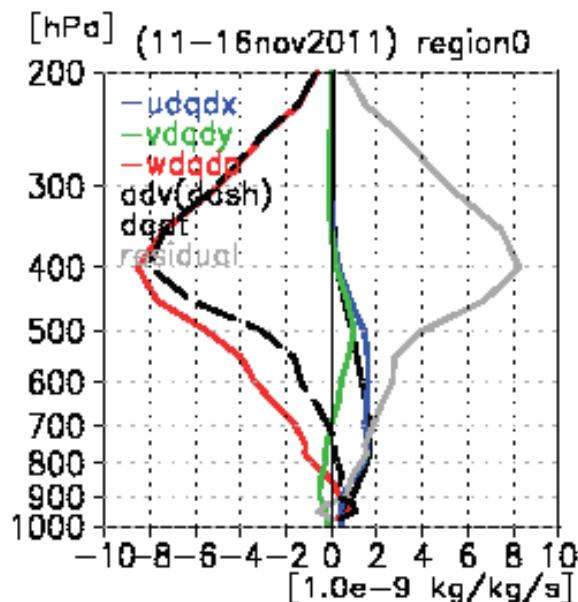
60-90E
10S-10N



$$(\frac{\partial q}{\partial t} \neq -V \nabla q' + -\omega \frac{\partial q}{\partial p} + -\frac{Q_2}{L})$$

MJO1
11-14 Oct

MJO2
11-16 Nov



NICAM CINDY2011 simulations

- Realtime forecasts
- 14~56 km mesh
- 40L
- 7-days (daily updated)
2011/7/27-2012/1/31
- SST: nudged to initial value

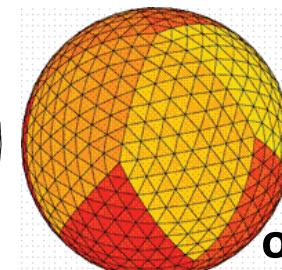
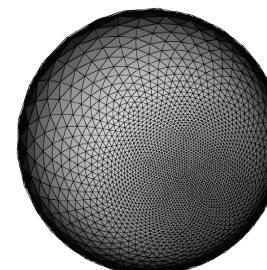
Nasuno (2013), SOLA
- Hindcasts (previous talks in this workshop)
- 14 km, 7 km, 3.5km mesh
- 40L
- 60 (30) days
init: 2011/10/16
(init: 2011/11/17)
- SST: fixed to observation etc.

NICAM simulation (real-time forecasts) Nasuno (2013, SOLA, 9, 69-73)

Nonhydrostatic ICosahedral Atmospheric Model

Satoh et al. (2008)

stretched grid
Tomita (2008)



original grid

Model: regionally stretched NICAM

Resolution: 14~56 km mesh (center: 80E, 8S), 40 levels

Domain of output data: 20-140E, 53S-37N

Length of forecasts: 7-days (6.5-days useful)

Period: 1 Oct. 2011-31 Jan. 2012

Frequency: daily

initialization: NCEP_FNL, interpolation

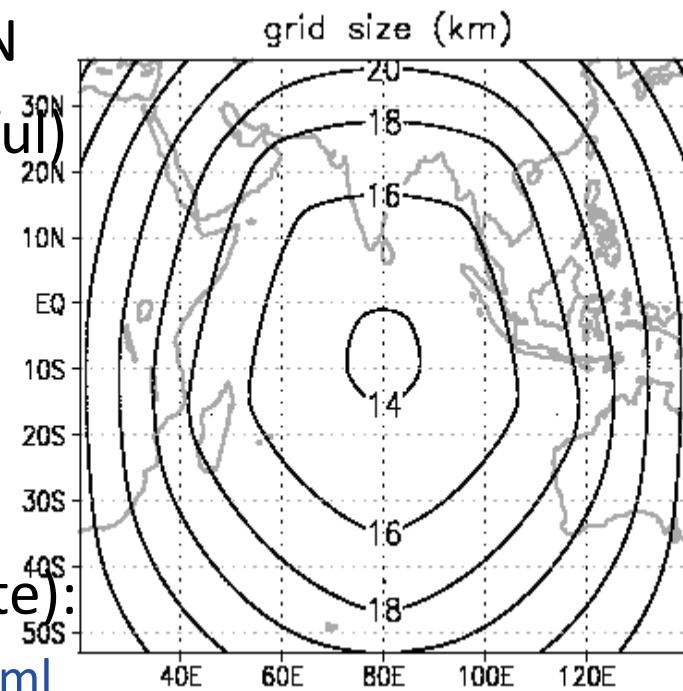
SST: slab ocean model (1-layer)

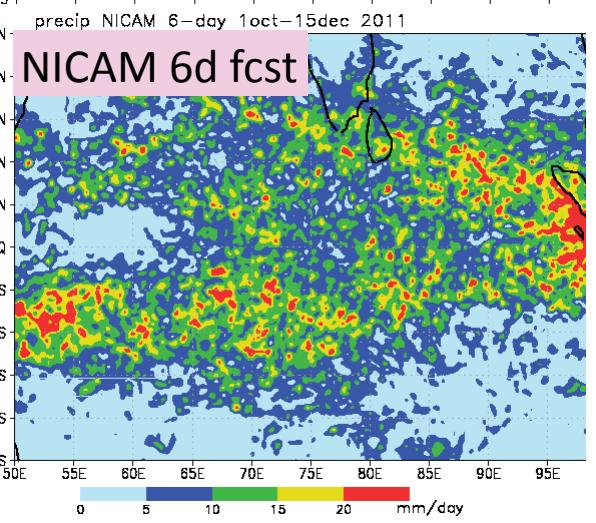
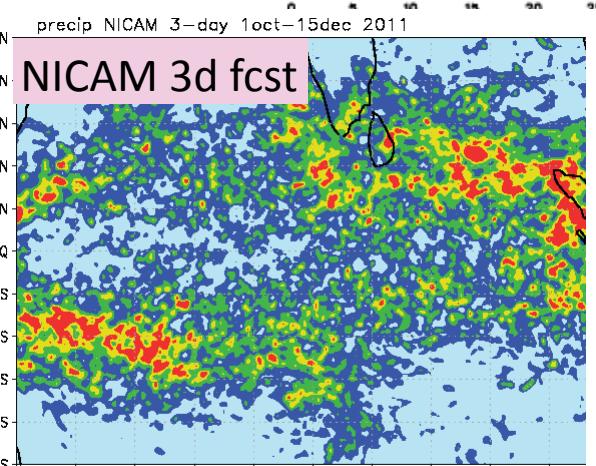
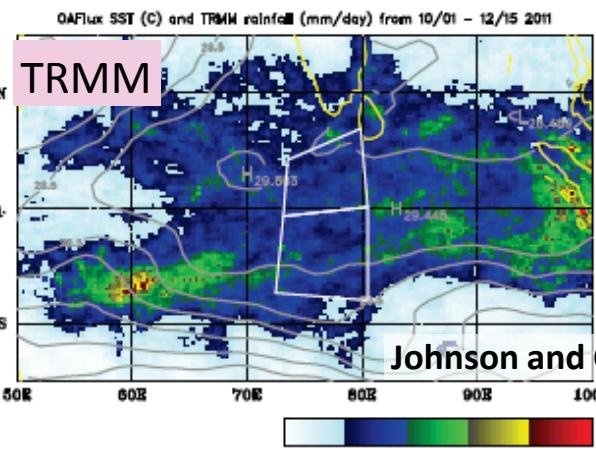
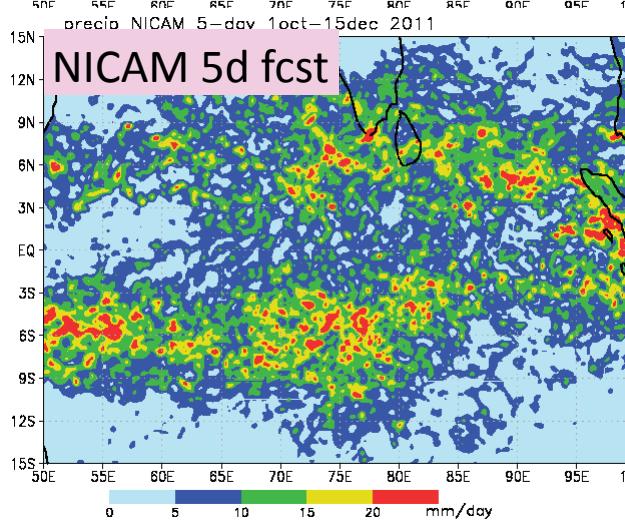
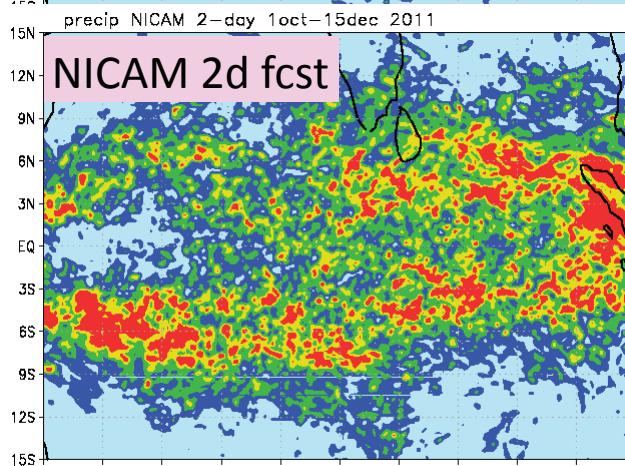
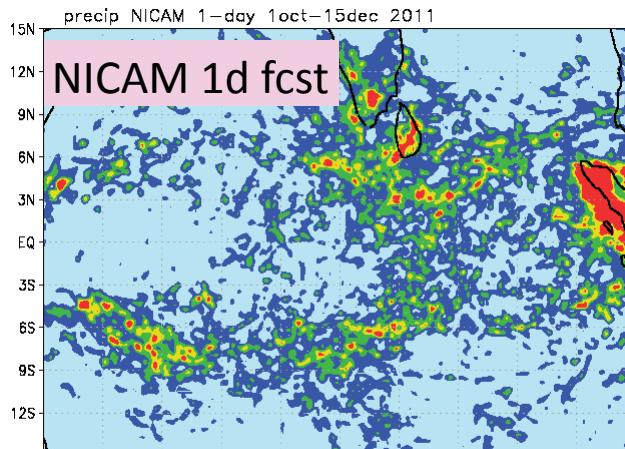
Data access (CINDY2011 JAMSTEC web site):

<http://www.jamstec.go.jp/iorgc/cindy/obs/obs.html>

link from DYNAMO page

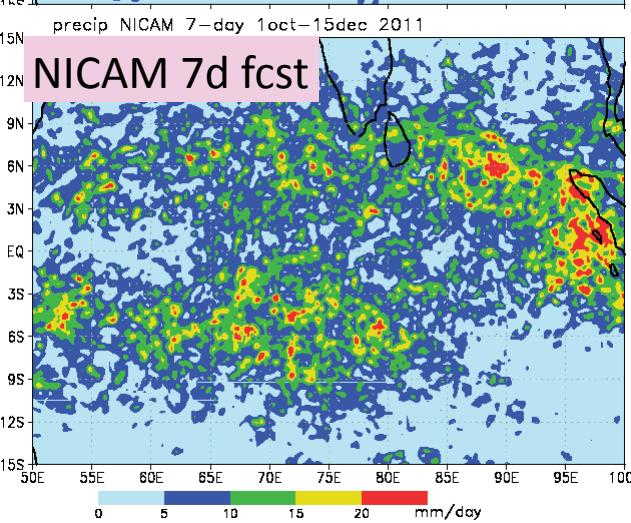
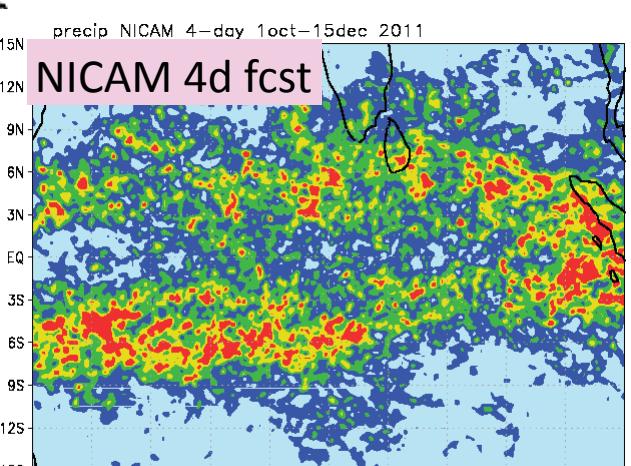
(http://data.eol.ucar.edu/master_list/?project=DYNAMO)

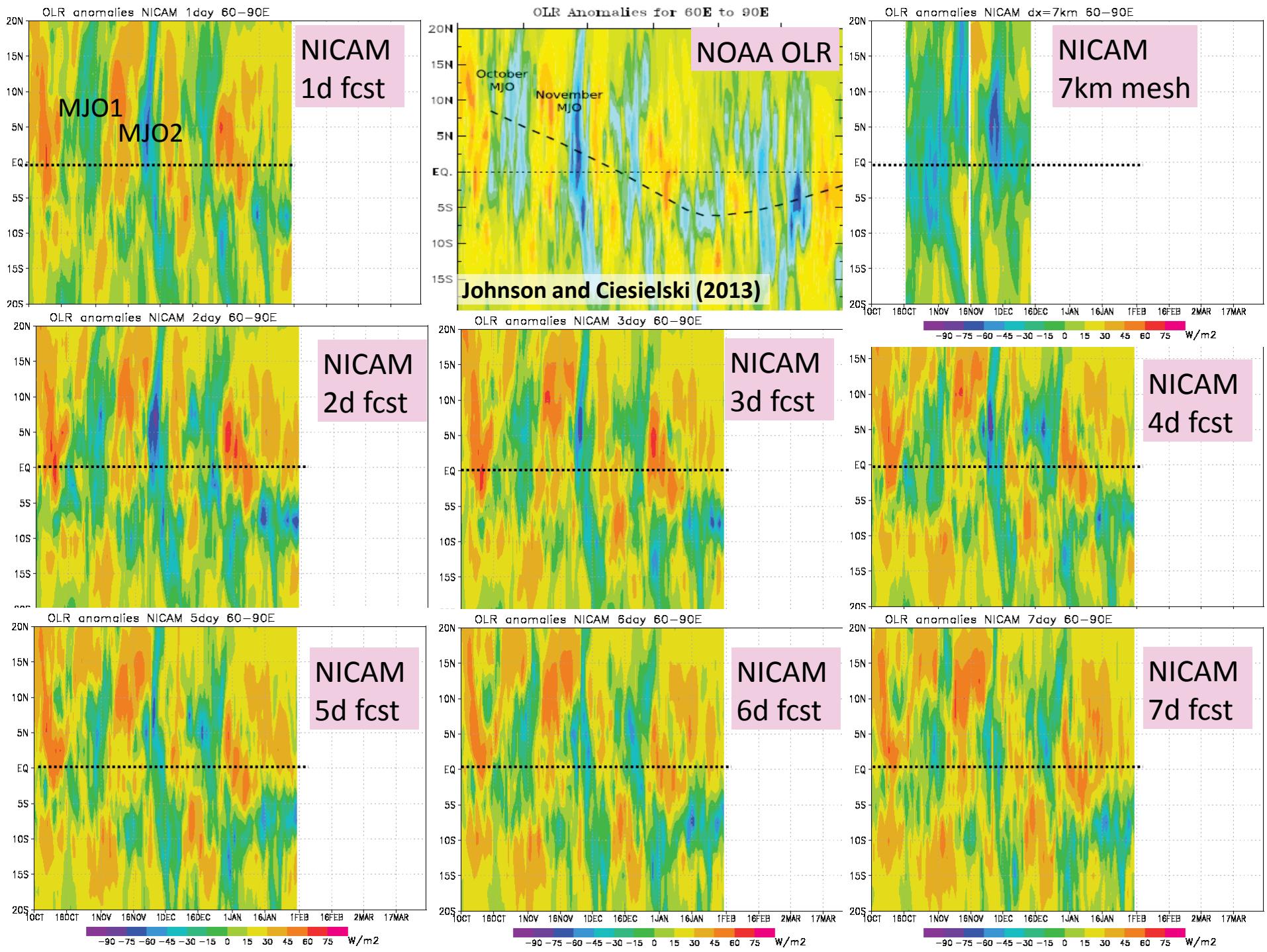




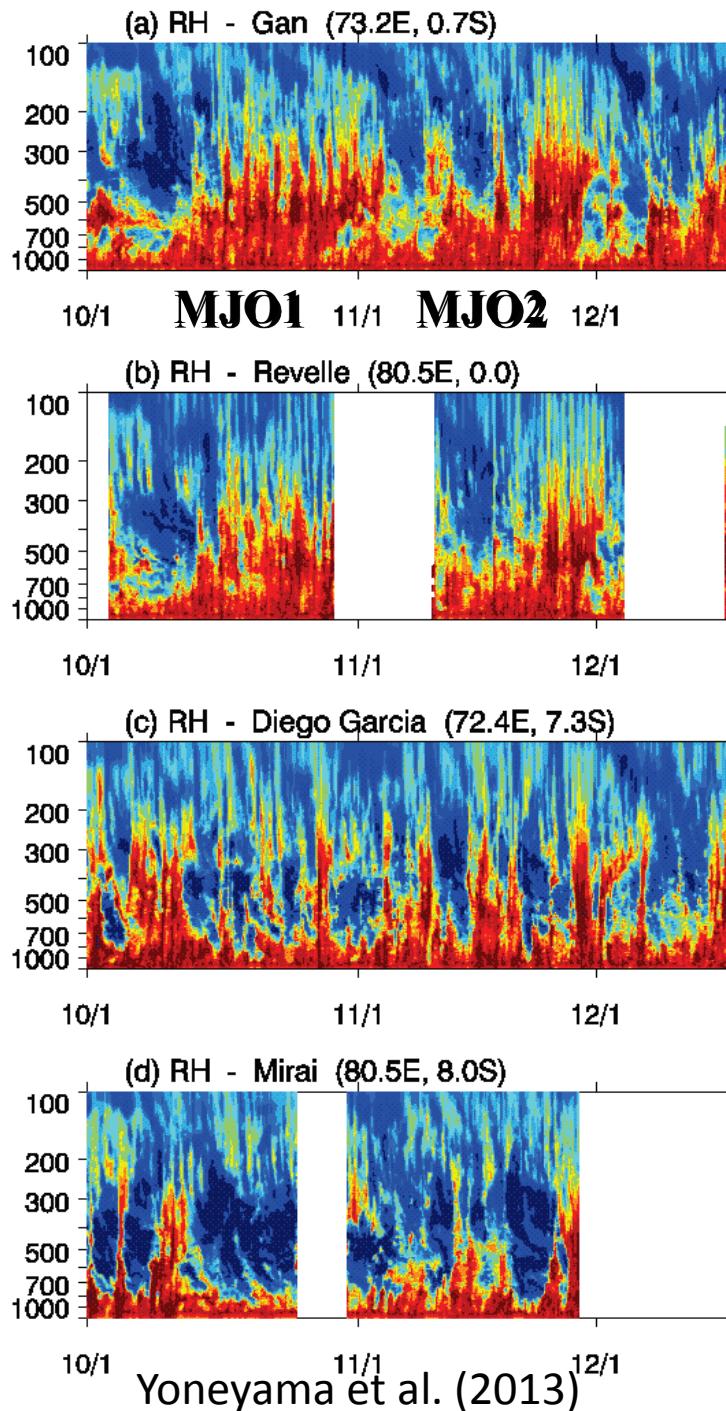
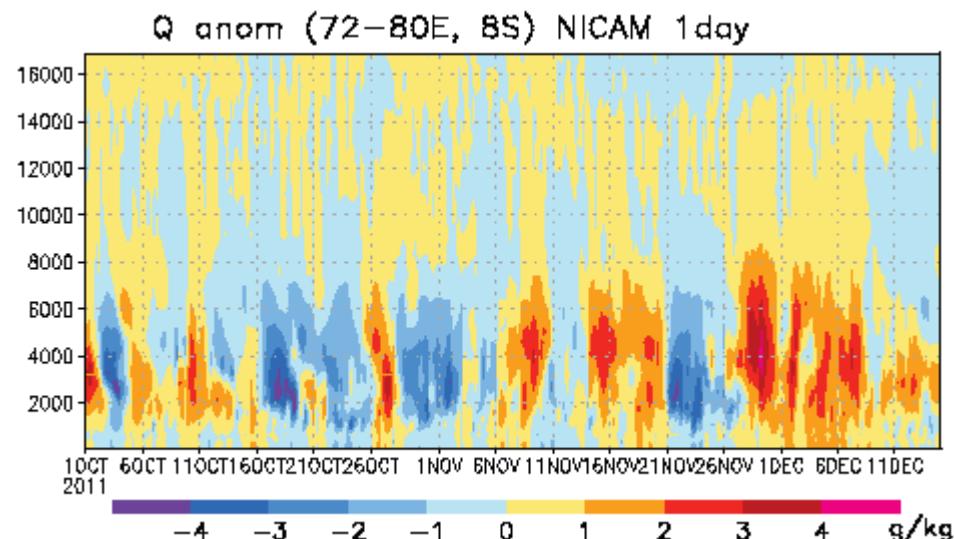
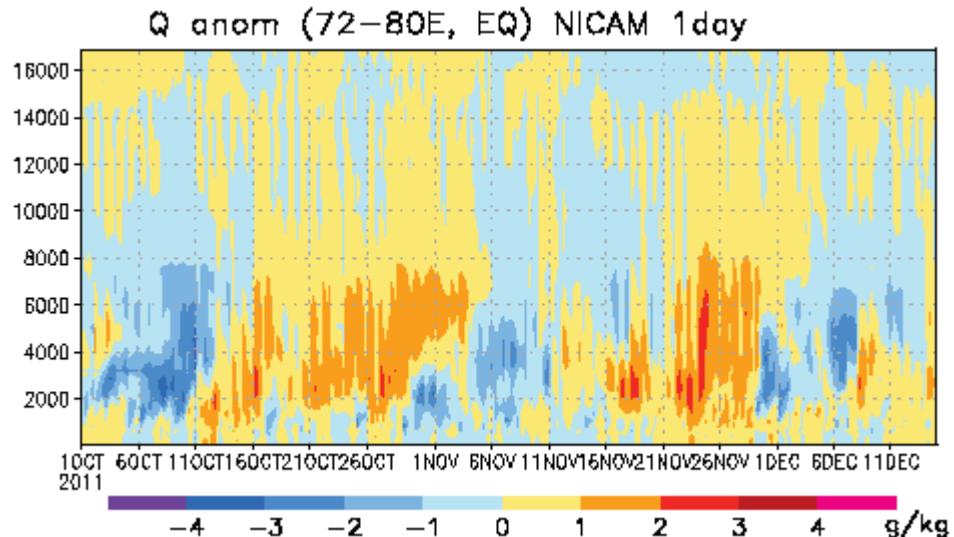
IOP mean (10/01-12/15 2011)
Daily precipitation
in stretch NICAM and TRMM

Johnson and Ciesielski (2013)



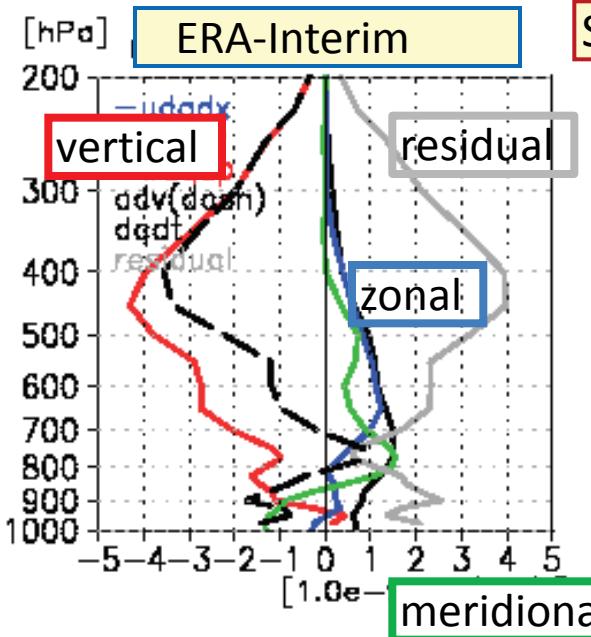


Moisture buildup in CINDY2011/DYNAMO Array and Stretch NICAM (1d forecast)

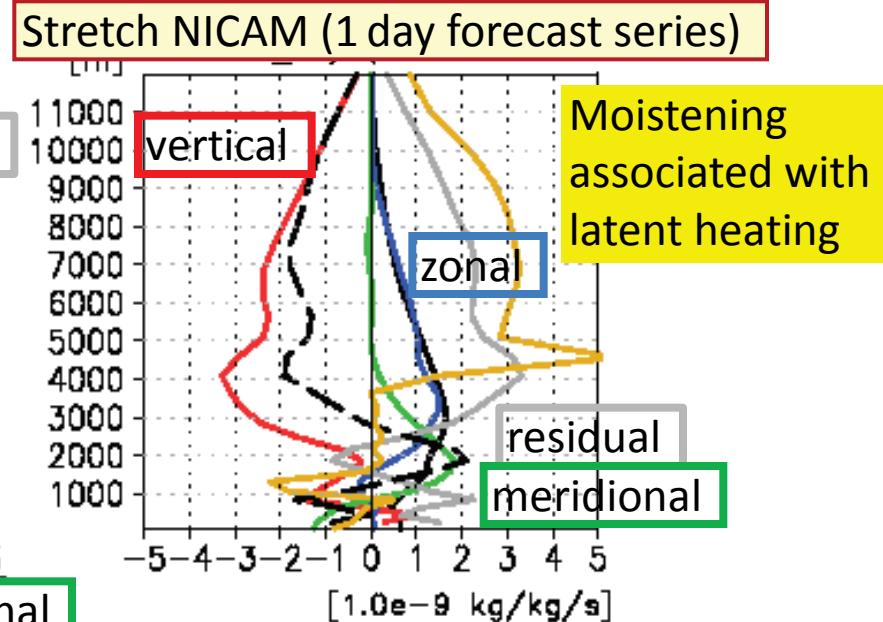


Moisture budgets

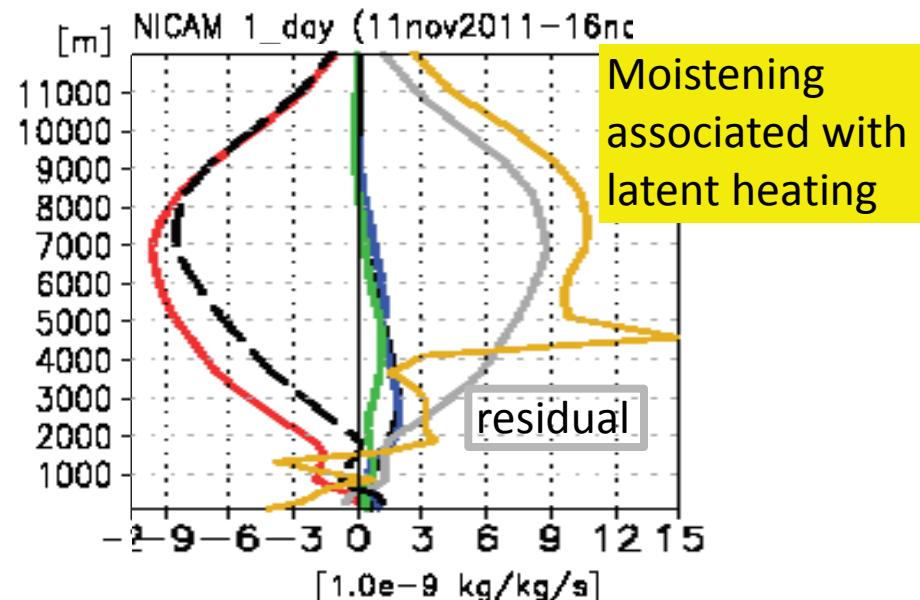
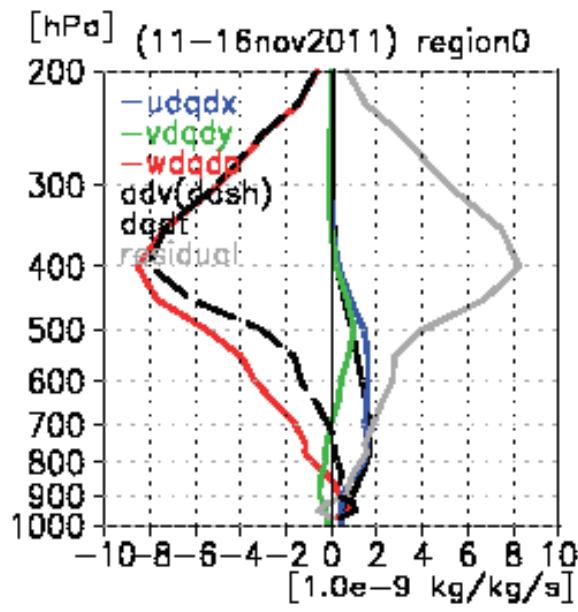
60-90E
10S-10N



MJO1
11-14 Oct



MJO2
11-16 Nov

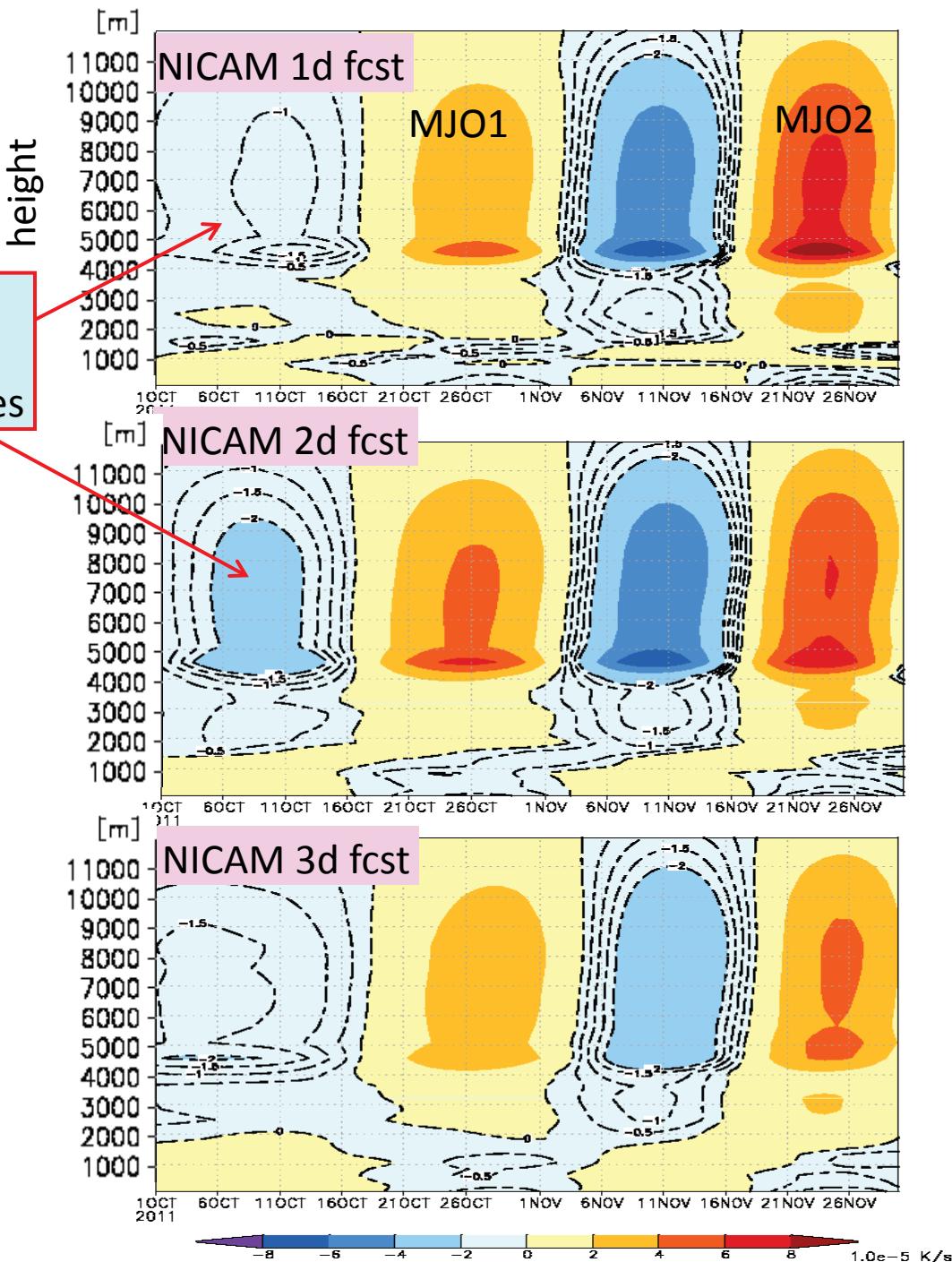


Latent heating

20-80d filtered

60-90E, 10S-10N

Evaporation
of preexisting
Water condensates



Decomposition of horizontal moisture advection

Zhao et al. 2013

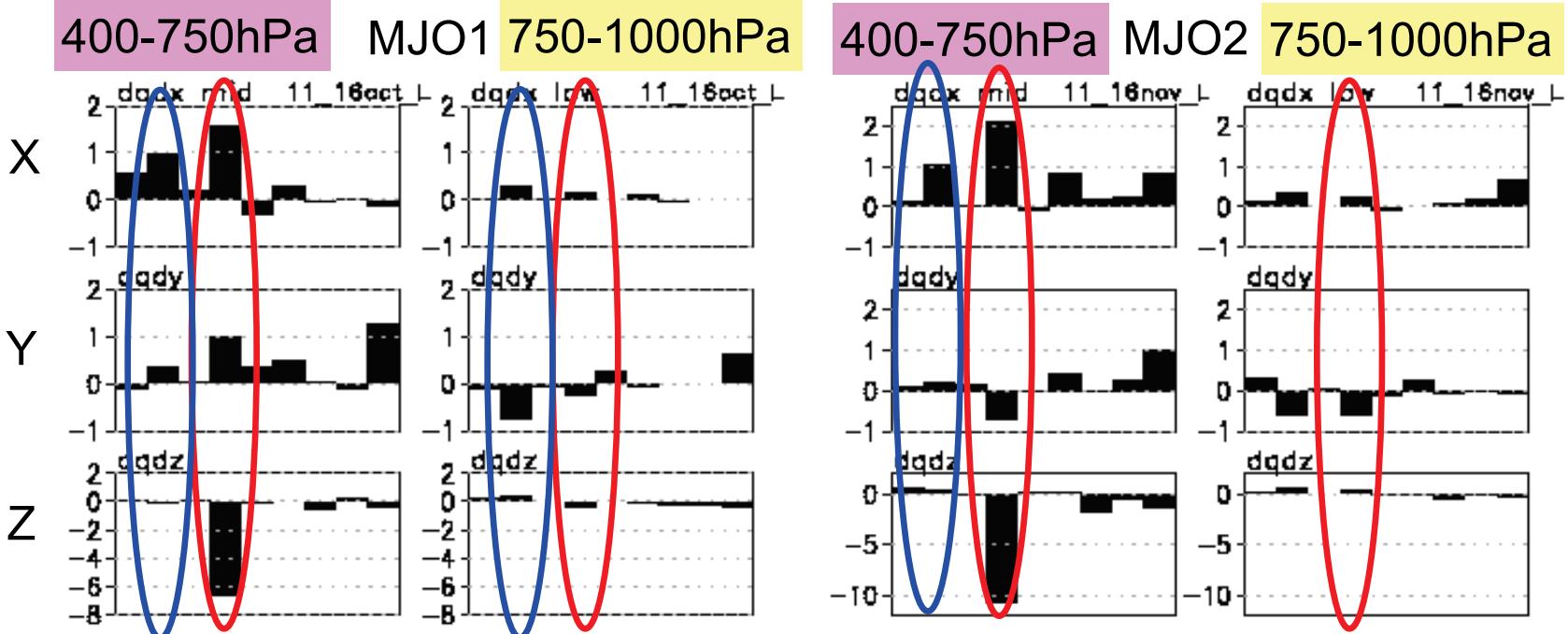
$$\left(\frac{\partial q}{\partial t} \right) \neq (-V \nabla \bar{q})' + -\omega \frac{\partial q}{\partial p} + -\frac{Q_2}{L}$$

60-90E, 10S-10N

$$u = \bar{u} + u' + u^*, \quad v = \bar{v} + v' + v^* \quad q = \bar{q} + q' + q^*$$

overbar: >80 day (basic state) prime: 20-80 day (ISO) asterisk: < 20 day (high-freq)

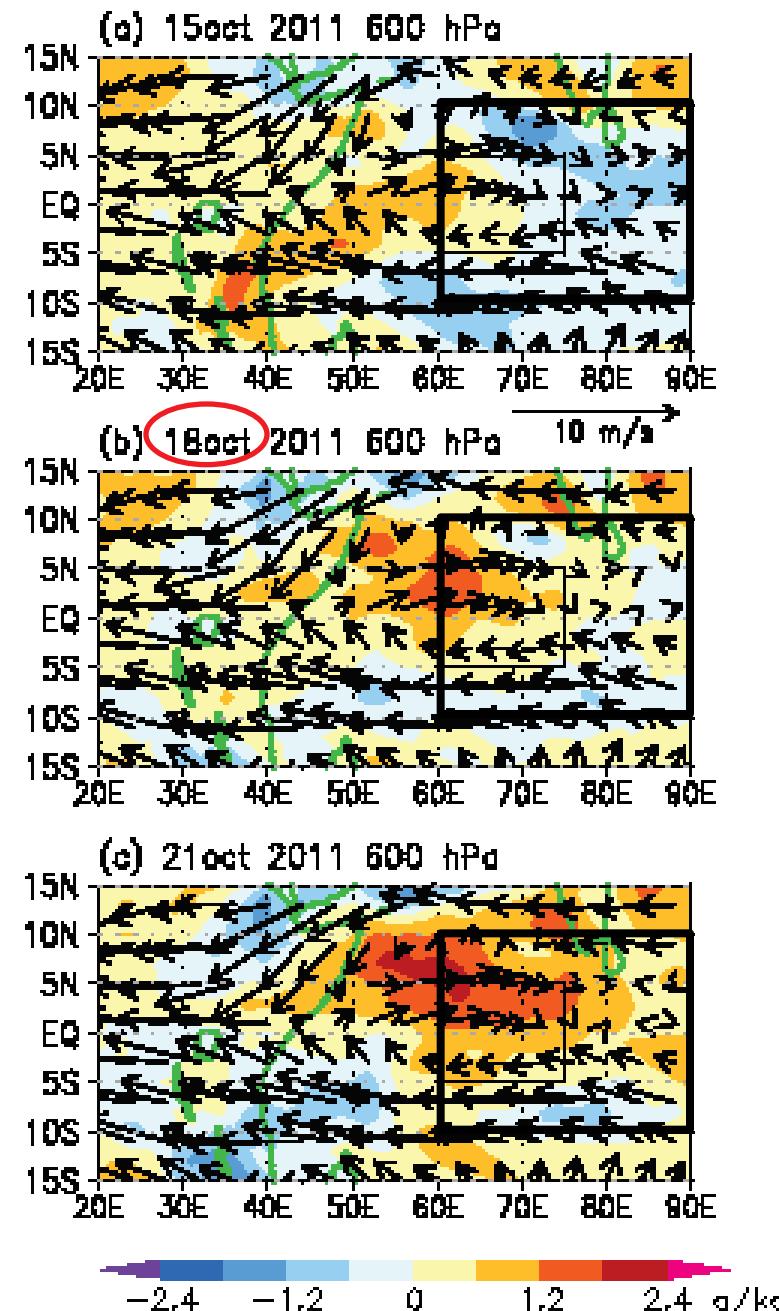
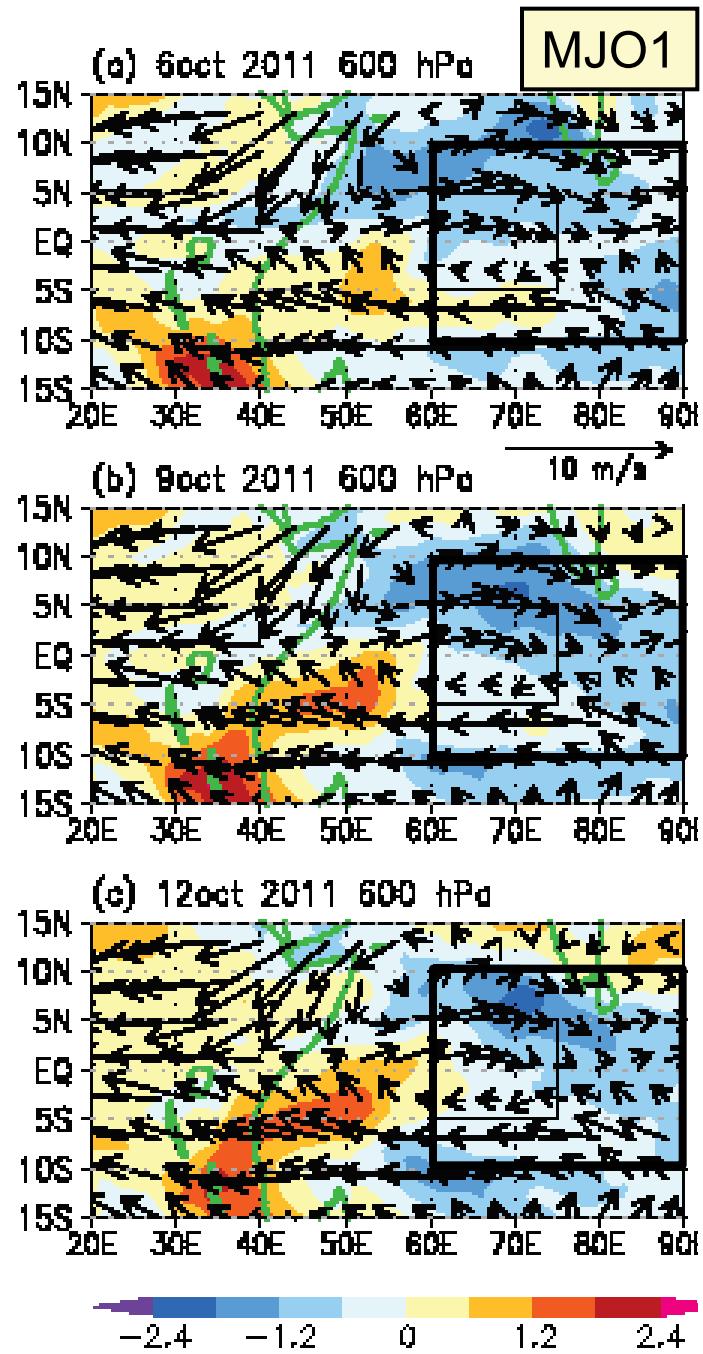
$(-\bar{V} \cdot \nabla \bar{q})'$	$(-\bar{V} \cdot \nabla q^*)'$	$(-\bar{V} \cdot \nabla q^*)'$	$(-\bar{V}' \cdot \nabla \bar{q})'$	$(-\bar{V}' \cdot \nabla q^*)'$	$(-\bar{V}' \cdot \nabla q^*)'$	$(-\bar{V}^* \cdot \nabla \bar{q})'$	$(-\bar{V}^* \cdot \nabla q^*)'$	$(-\bar{V}^* \cdot \nabla q^*)'$
1	2	3	4	5	6	7	8	9



Advection of ISO moisture by basic winds Basic moisture by ISO winds

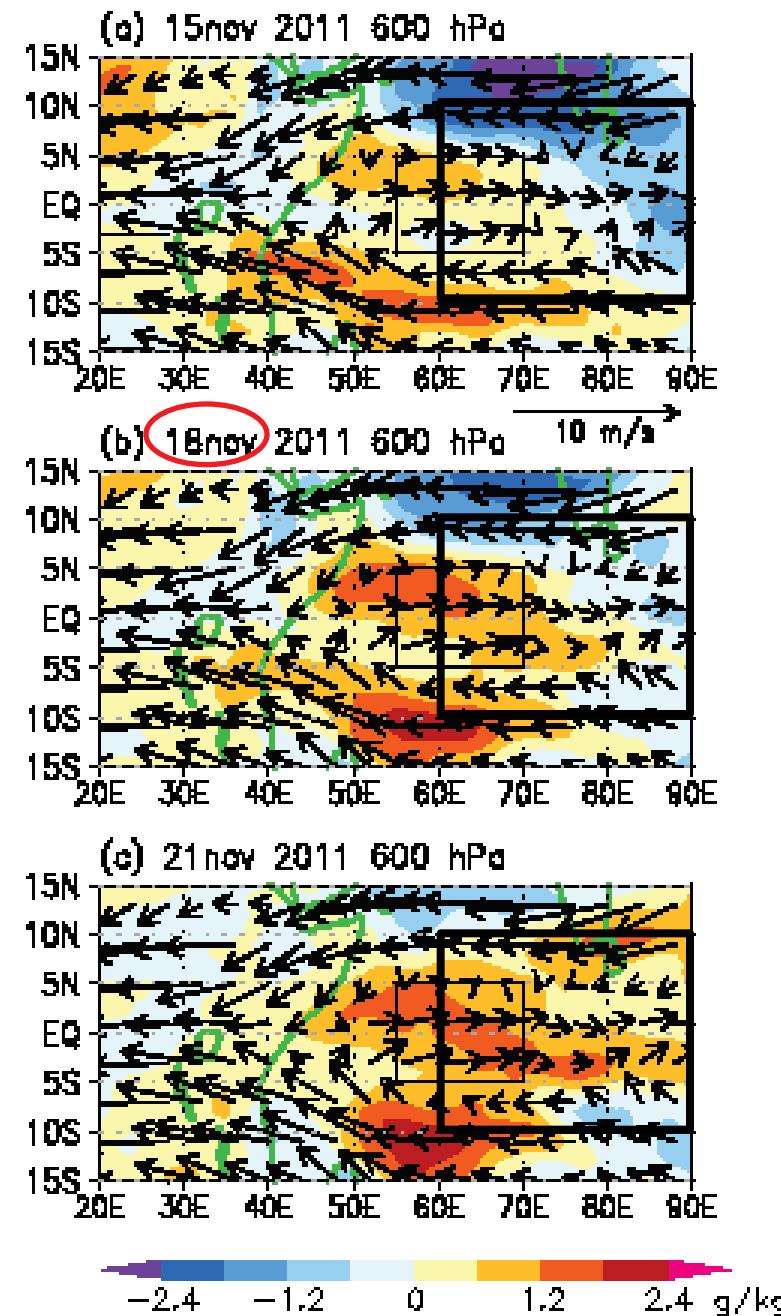
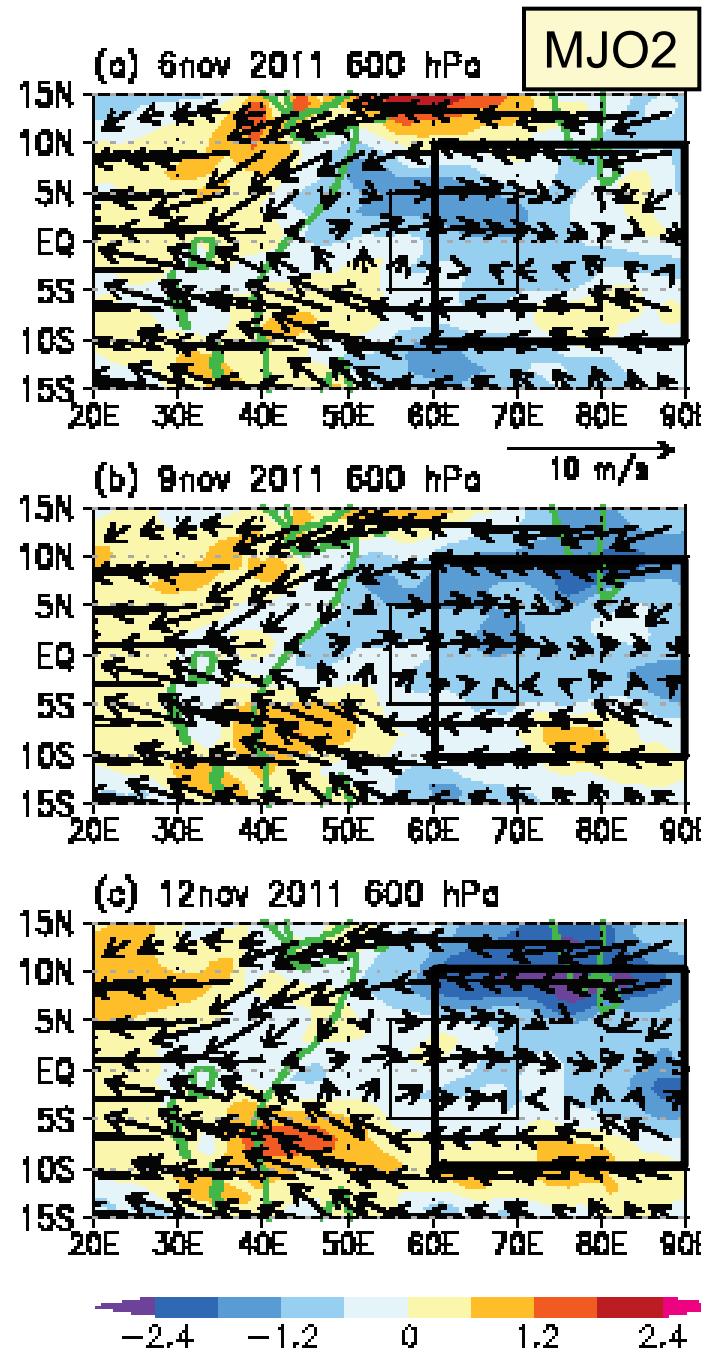
Era-Interim

600 hPa
Basic wind
(> 80d)
ISO moisture
(20-80 d)



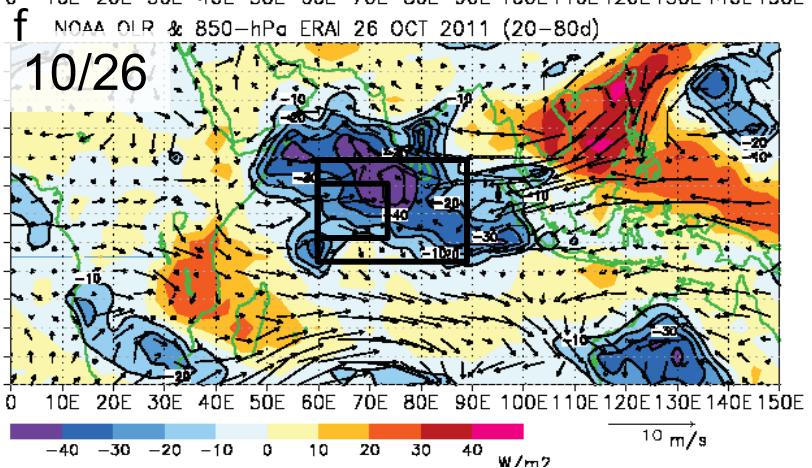
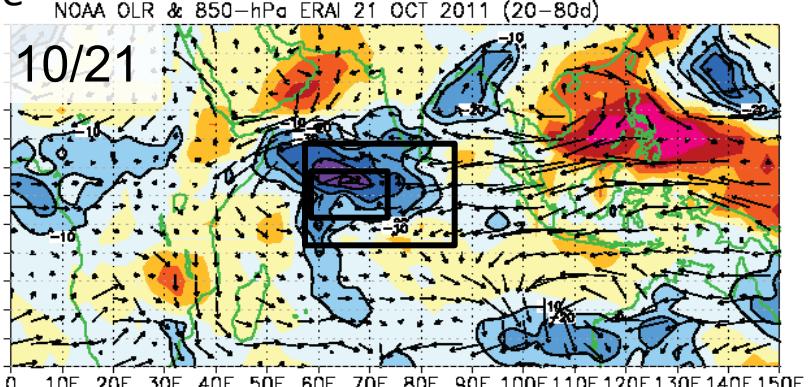
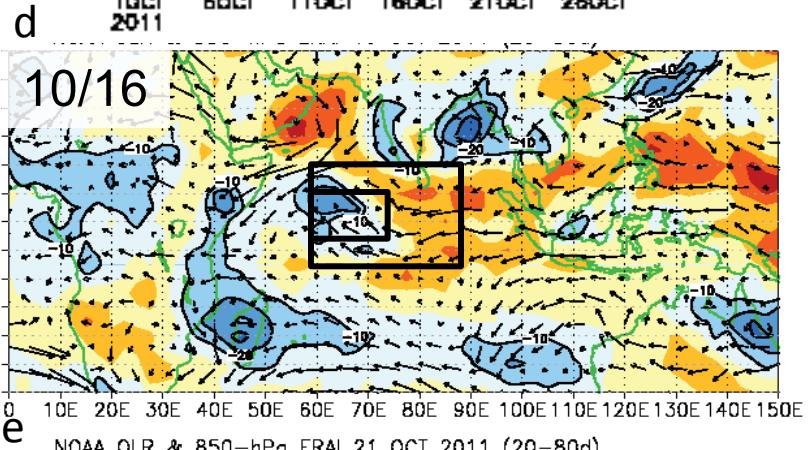
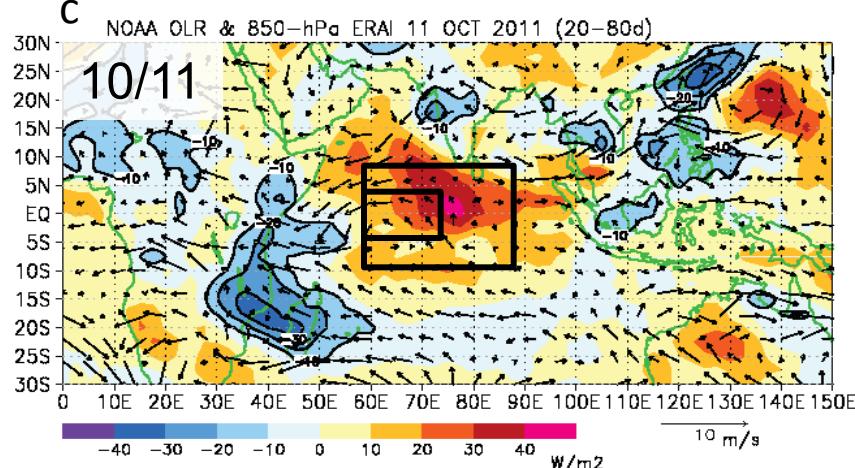
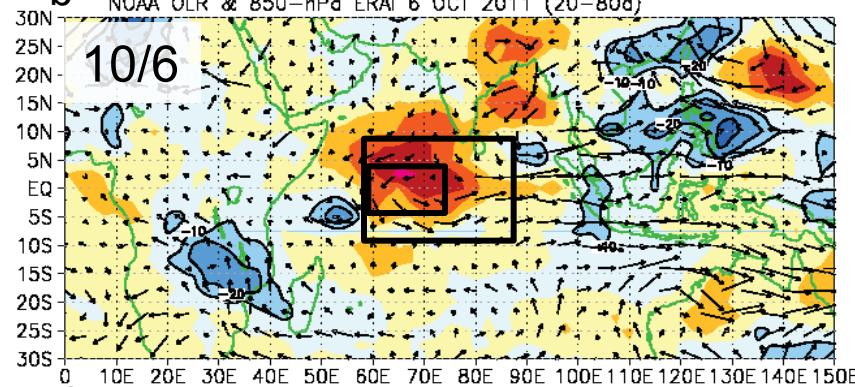
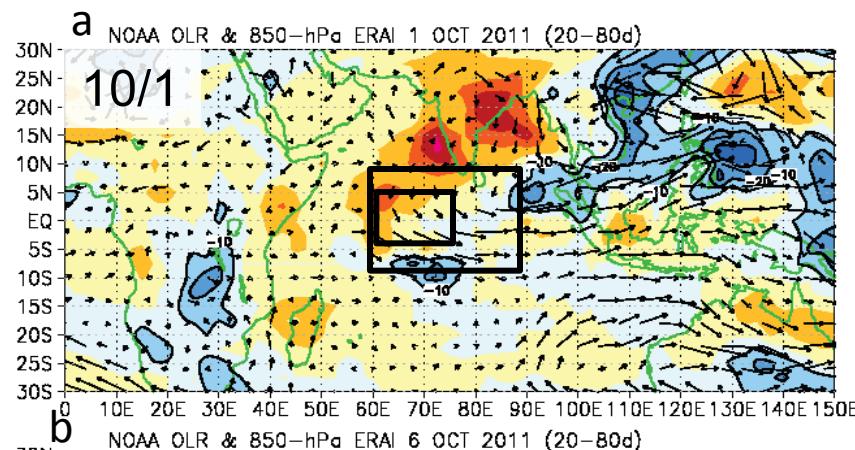
Era-Interim

600 hPa
Basic wind
(> 80d)
ISO moisture
(20-80 d)

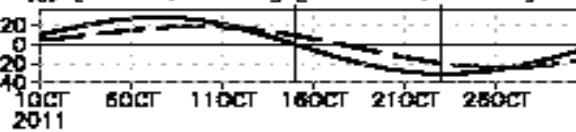


OLR (20-80 day) 850hPa wind

MJO1

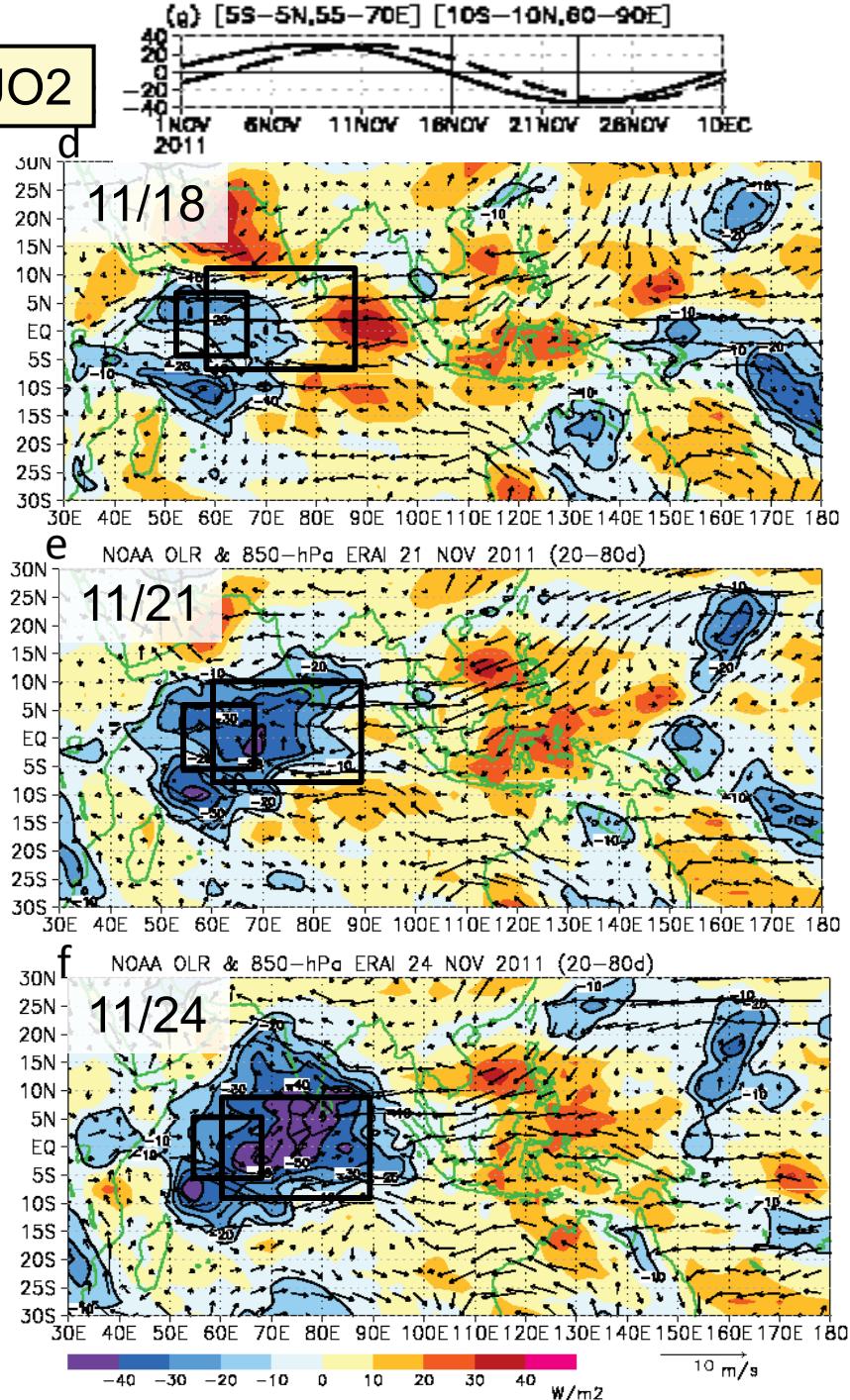
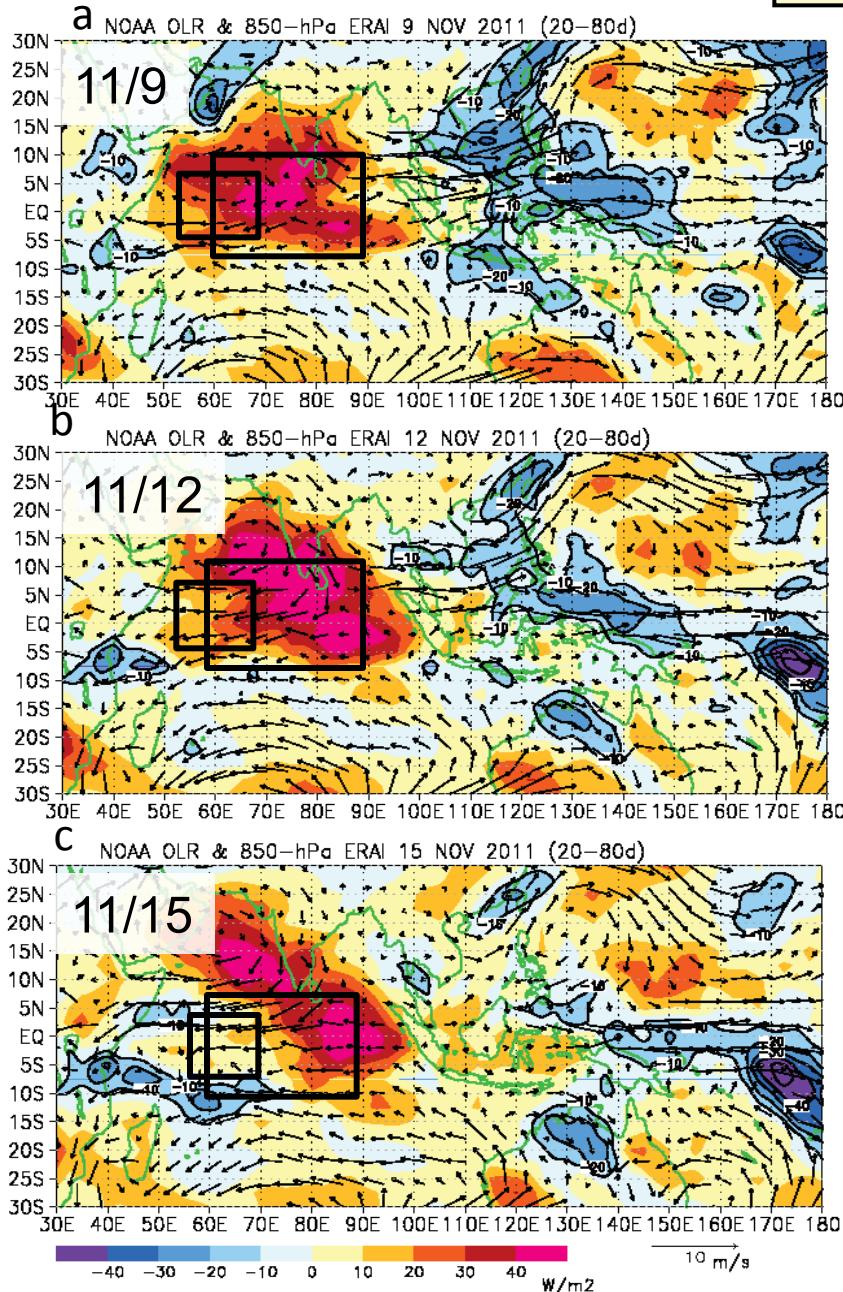


(e) [5S-5N,60-75E] [10S-10N,60-90E]



OLR (20-80 day) 850hPa wind

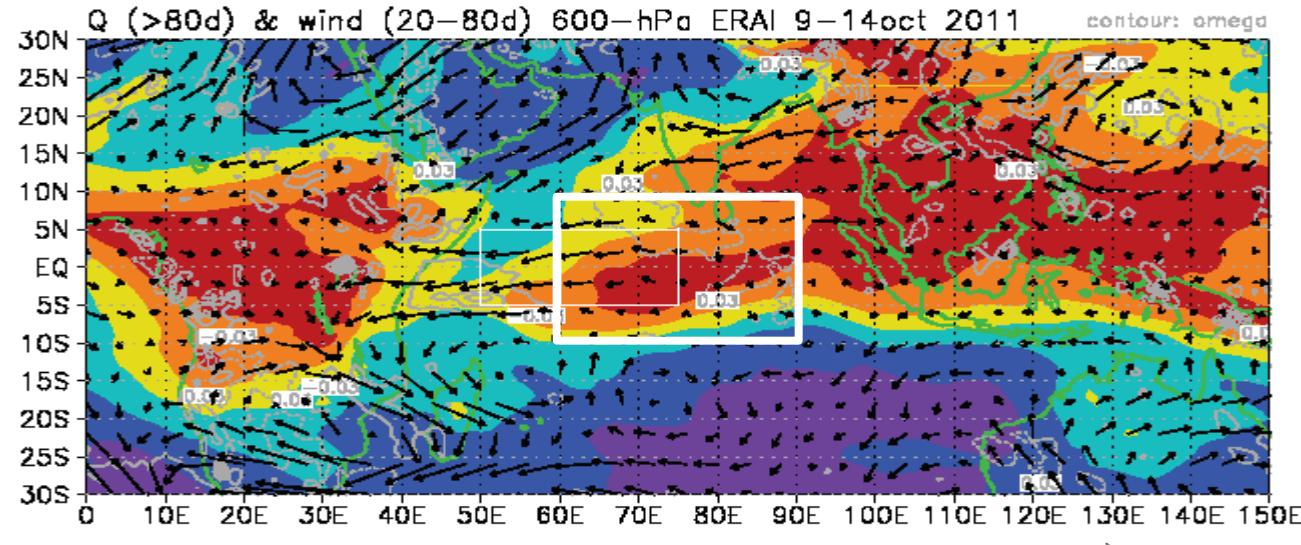
MJO2



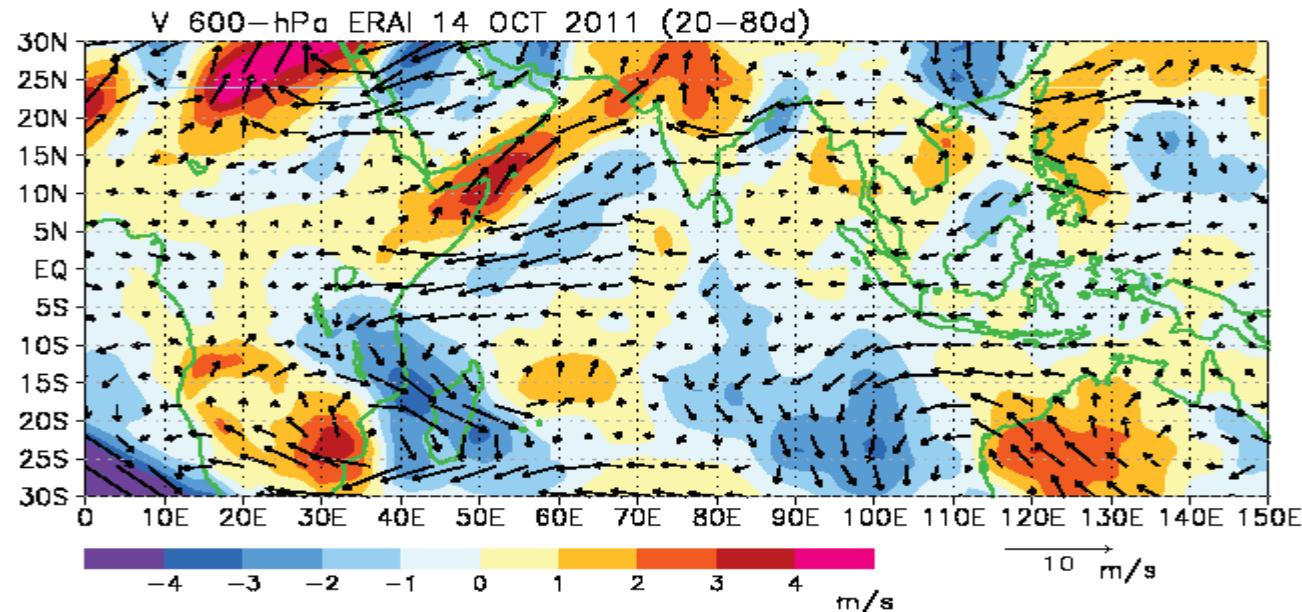
Era-Interim

MJO1
600 hPa

Basic
moisture
(> 80d)
ISO wind
(20-80 d)



V (20-80d)



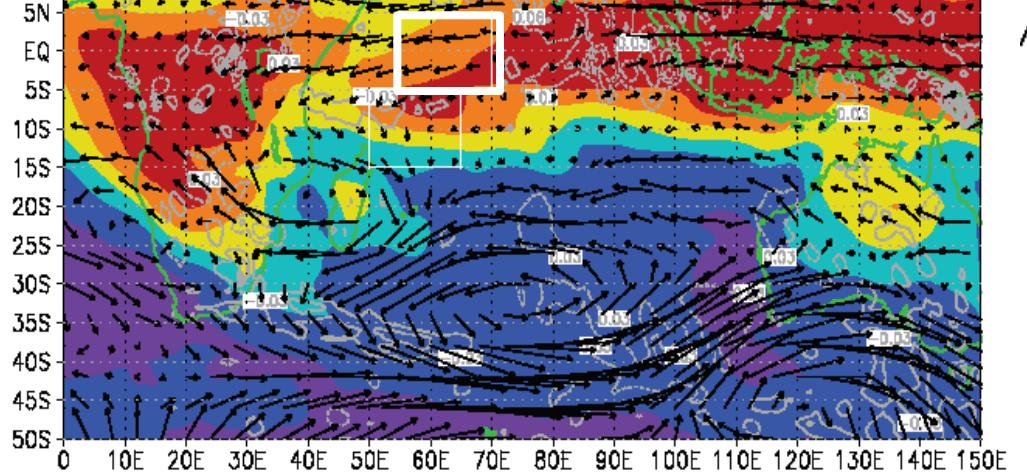
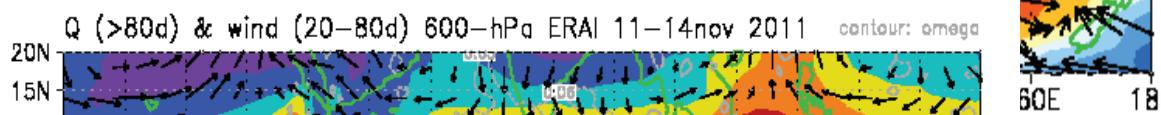
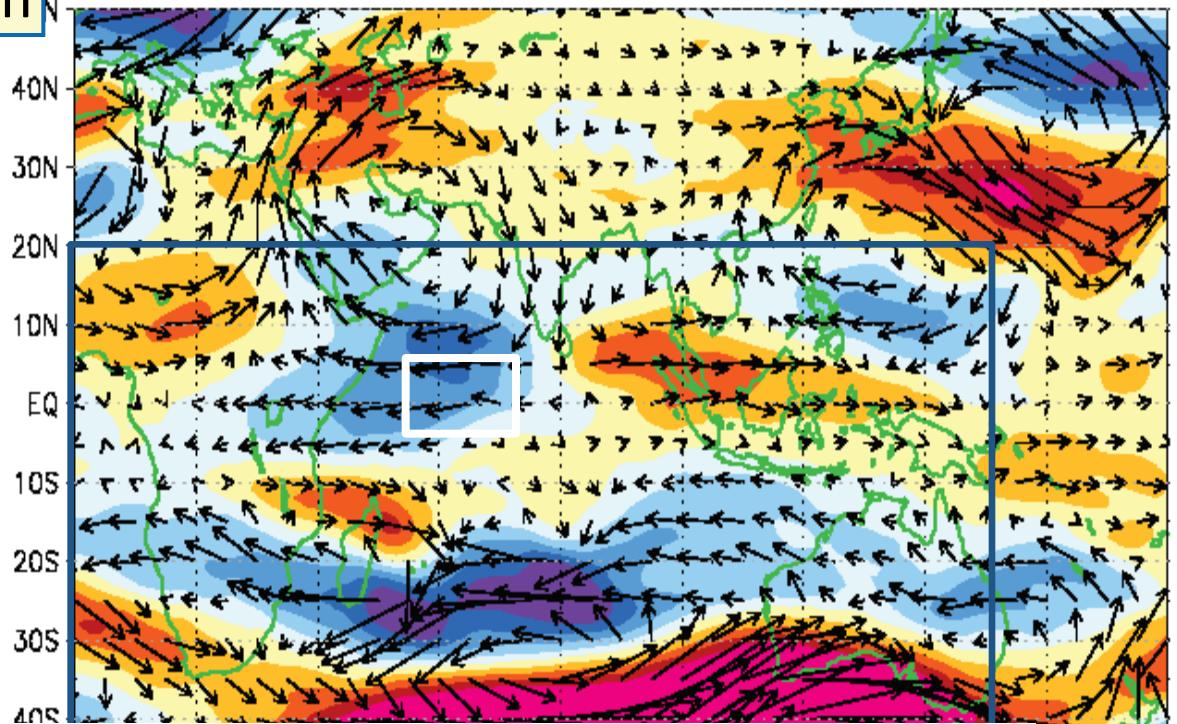
Era-Interim

MJO2

500 hPa
U(20-80d)

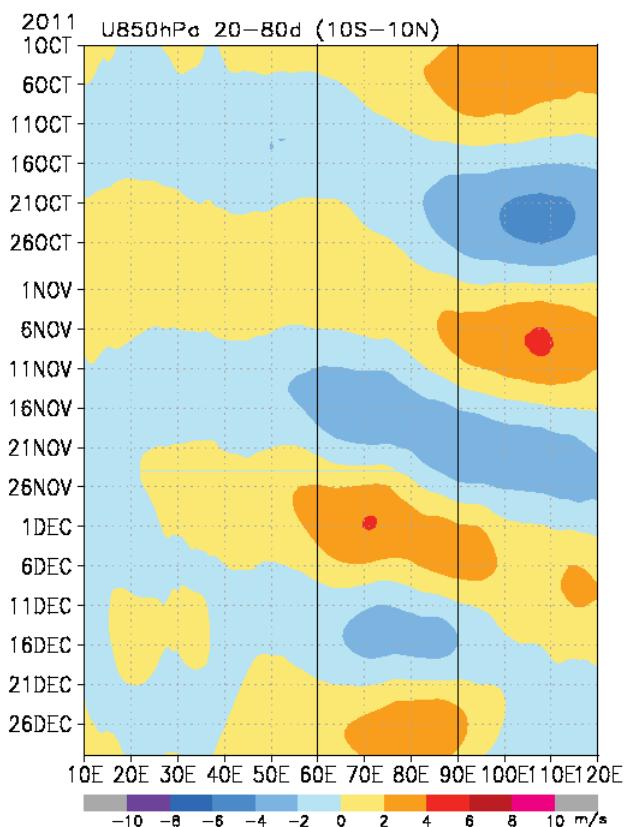
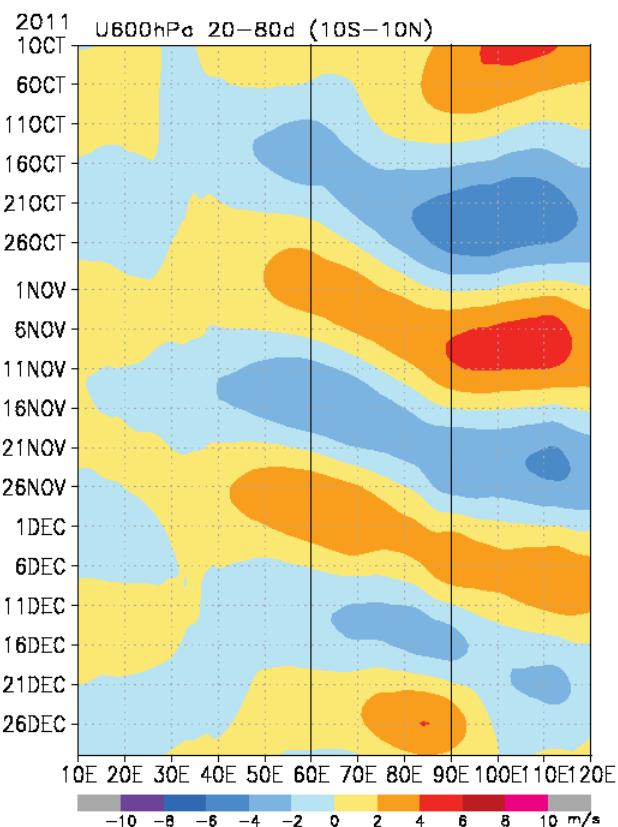
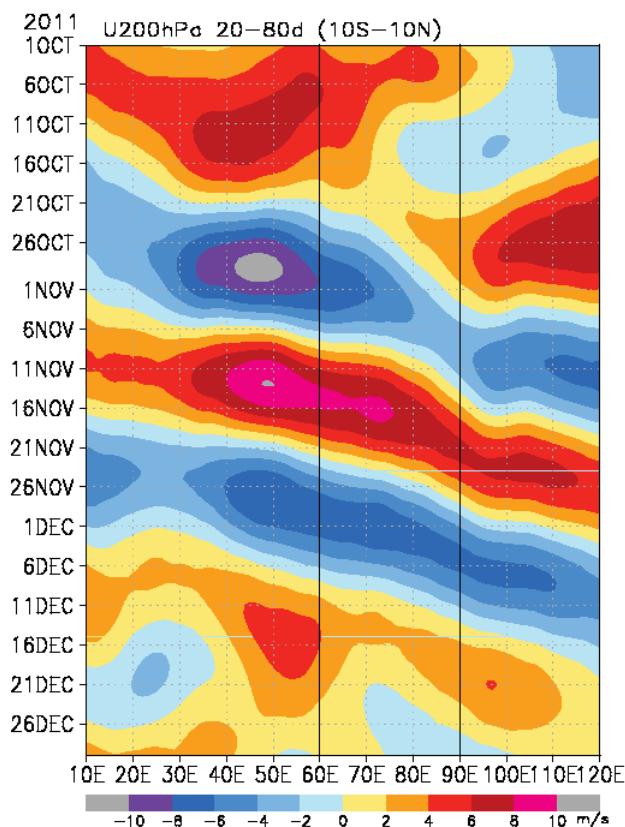
600 hPa
Basic
moisture
(> 80d)
ISO wind
(20-80 d)

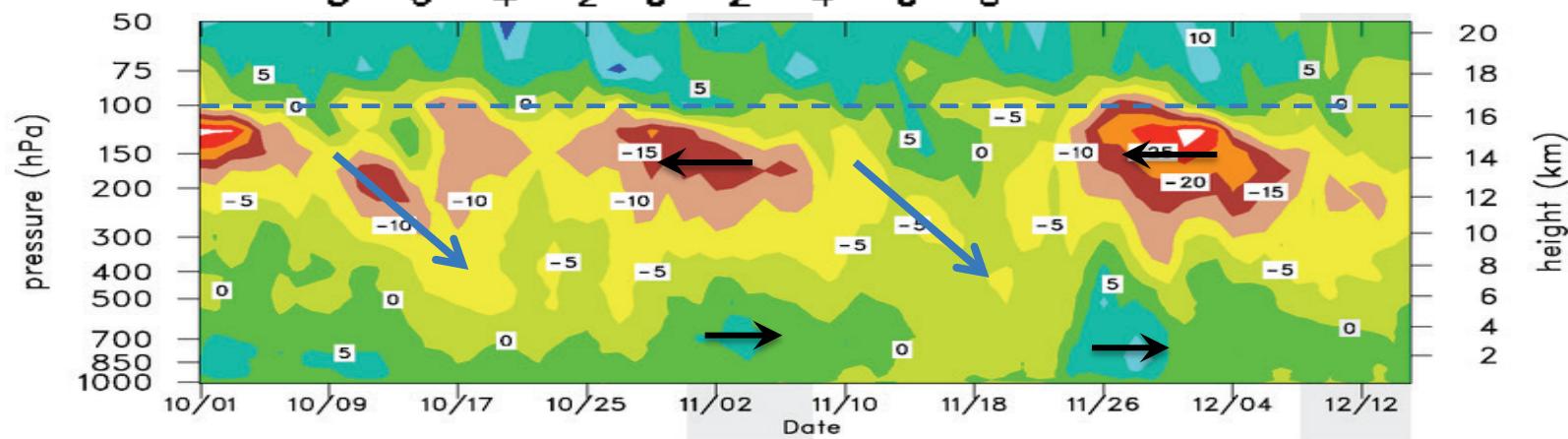
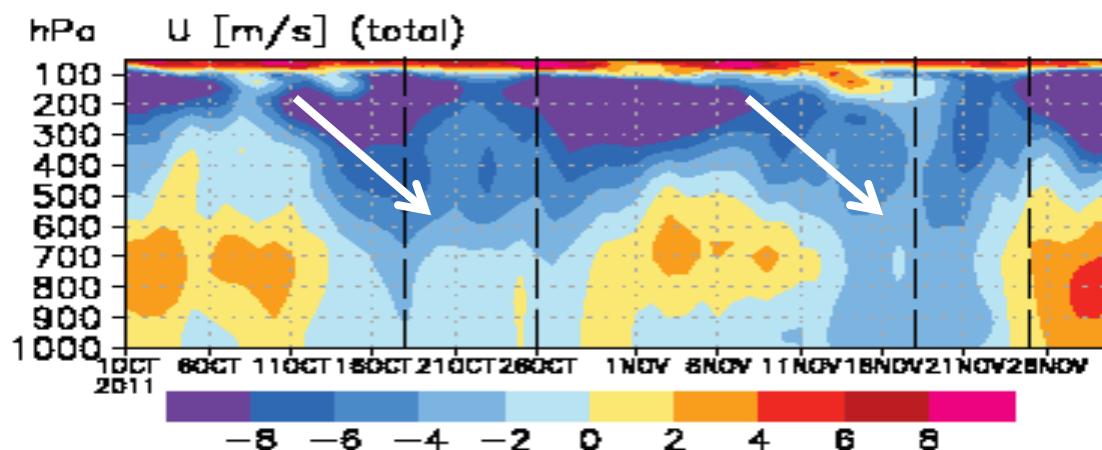
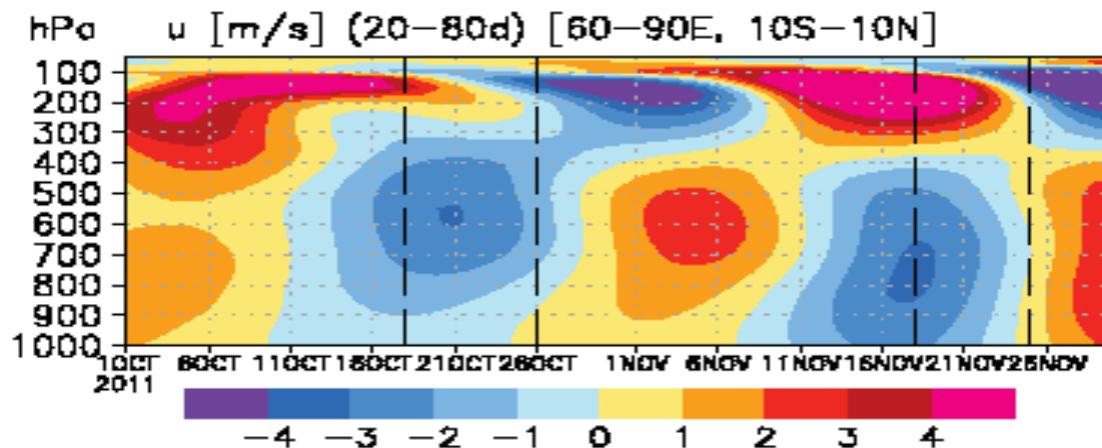
U 500 hPa ERAI 13nov 2011 (20-80d)



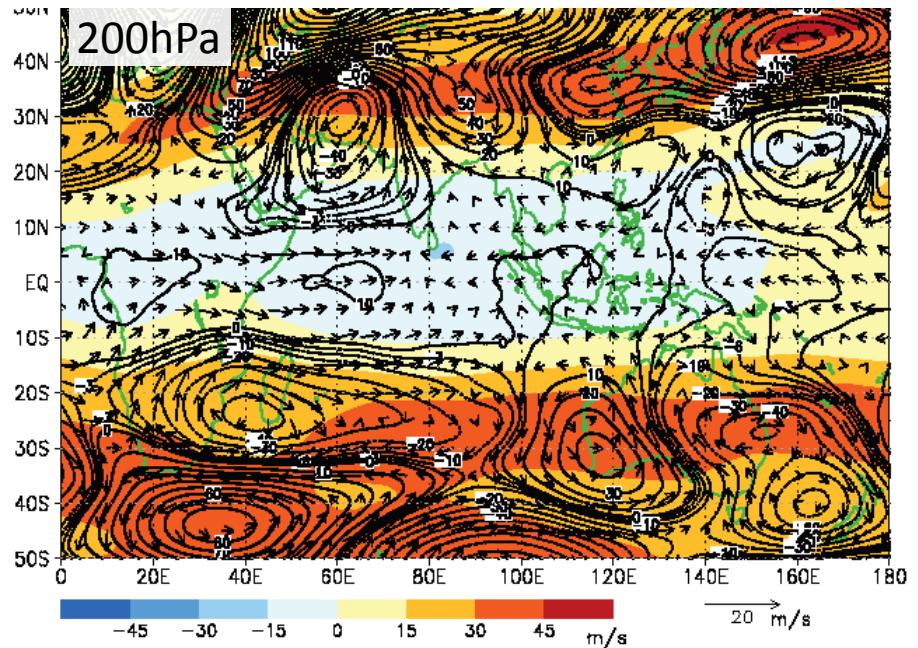
Discussion

20-80 day filtered U (10S-10N)

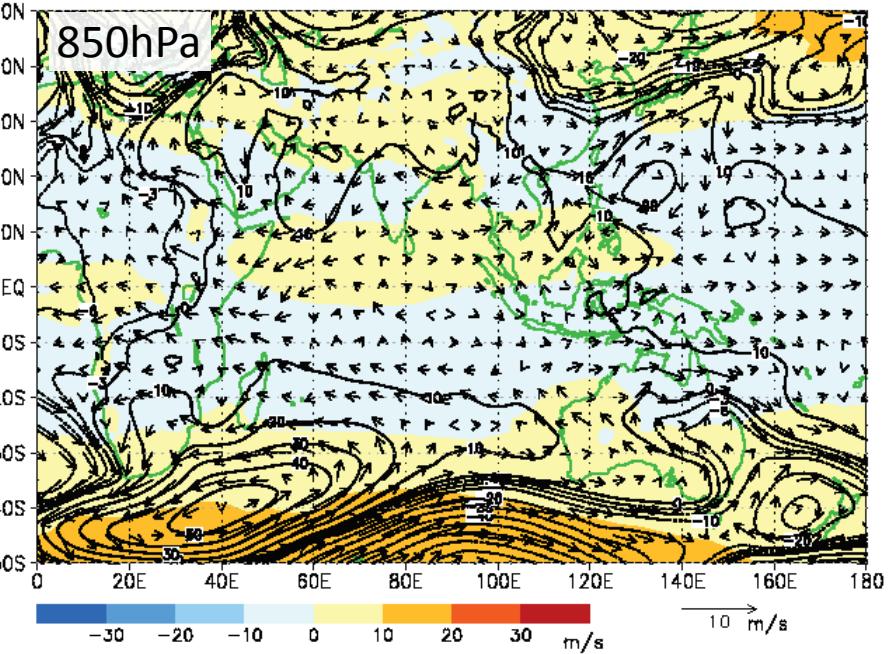
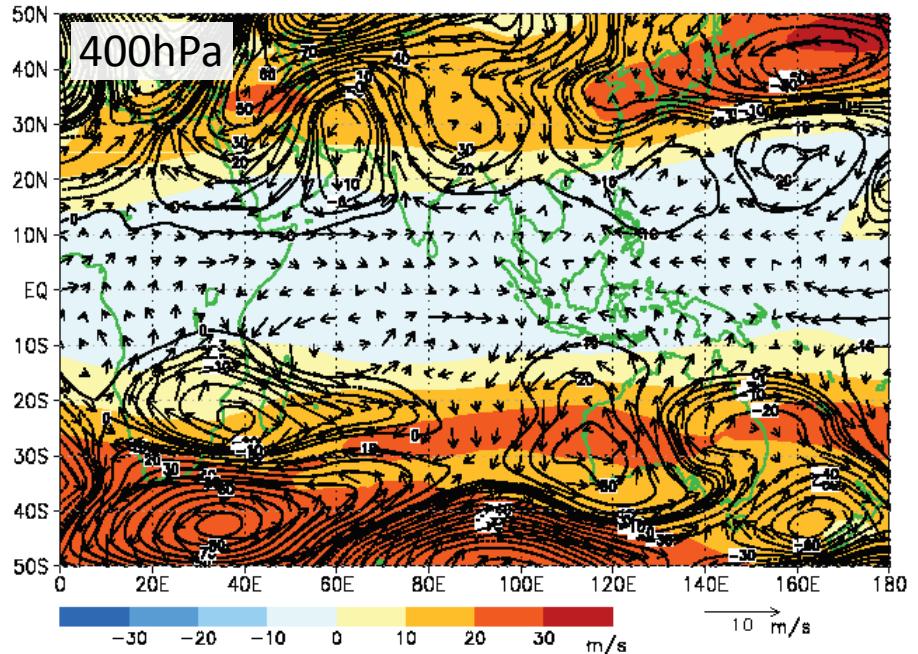
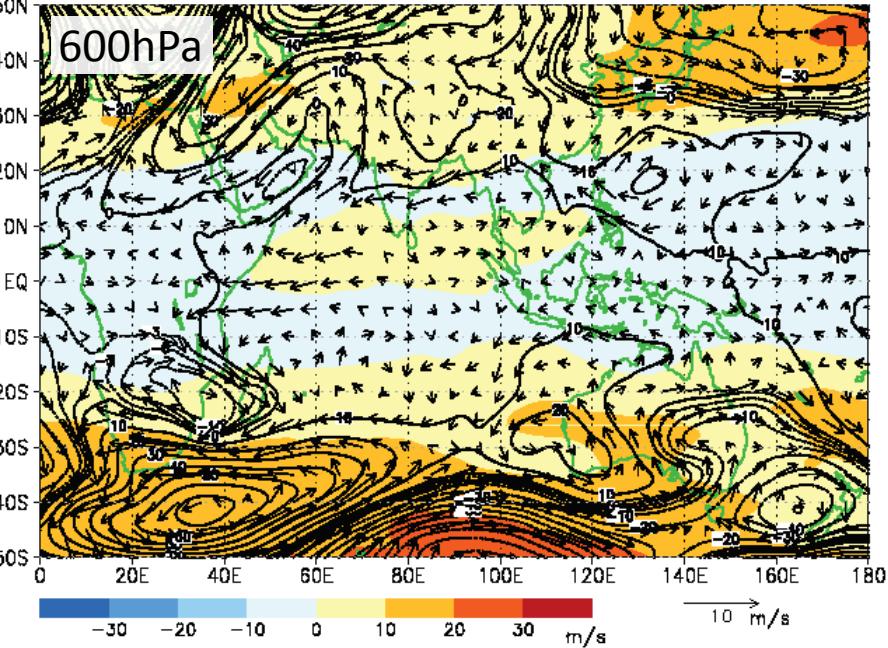




U (>80d), geopotential height & wind (20-80d)

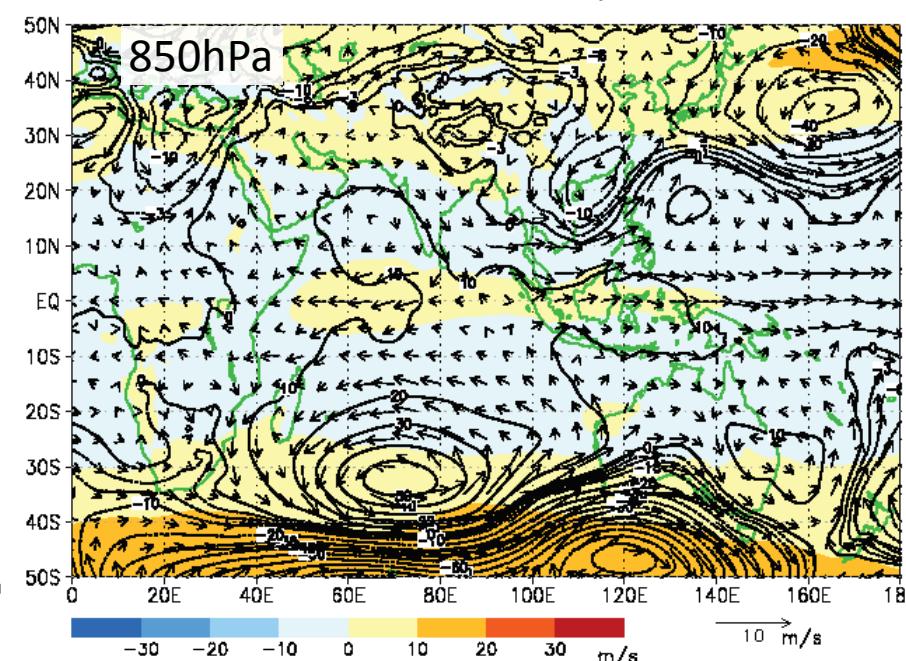
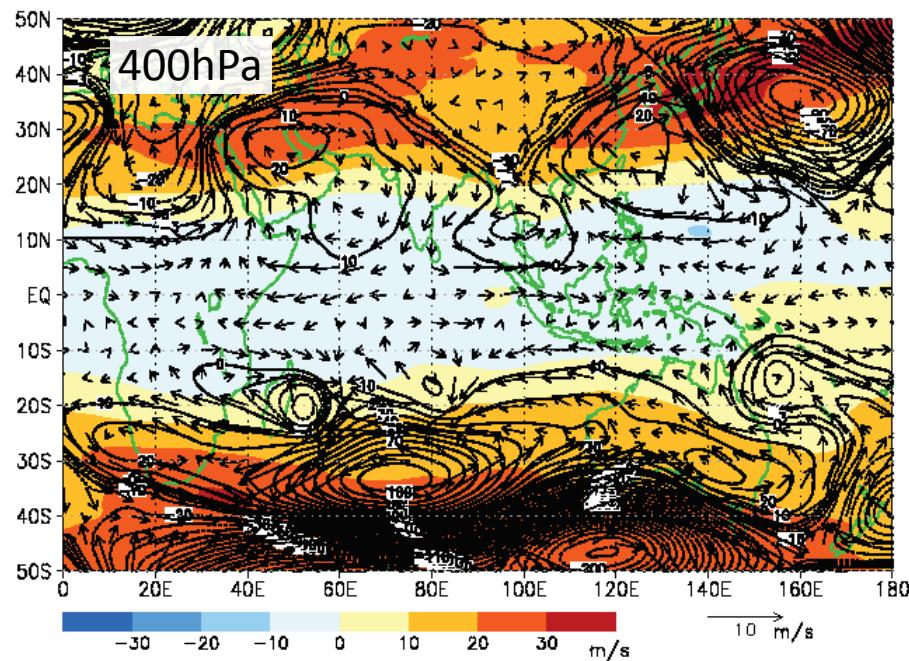
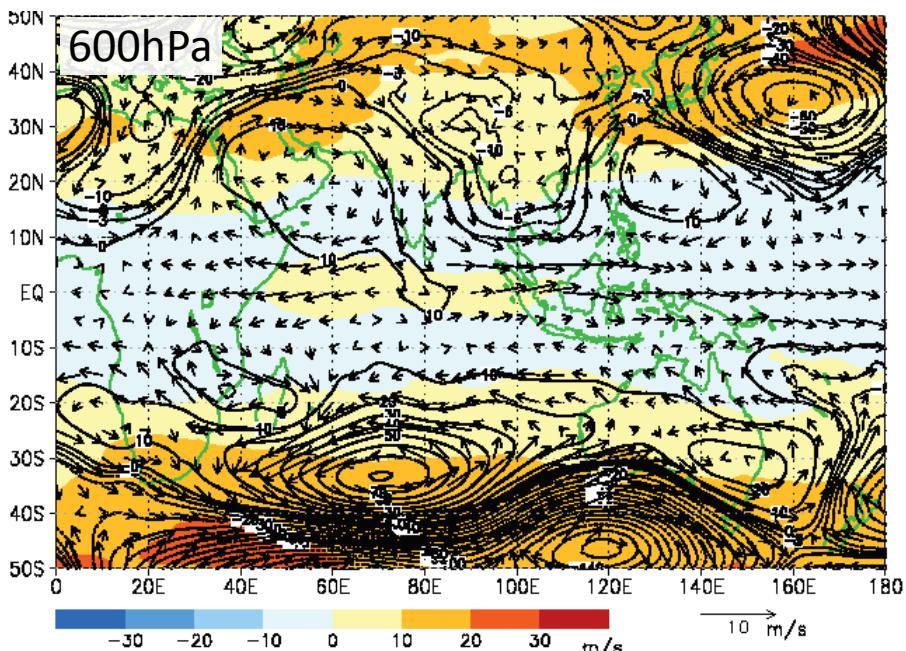
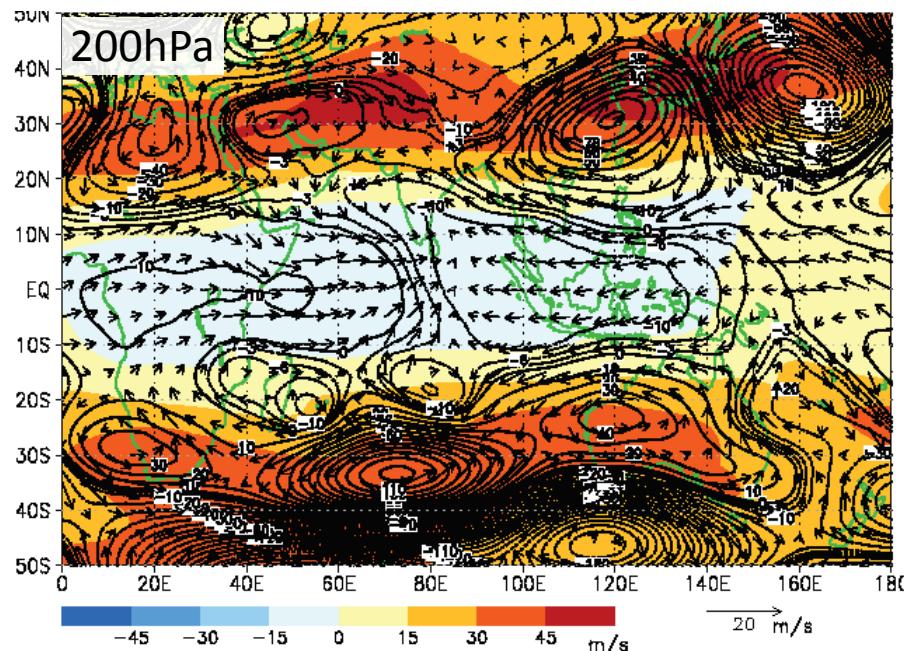


MJO1 (9-14 oct)

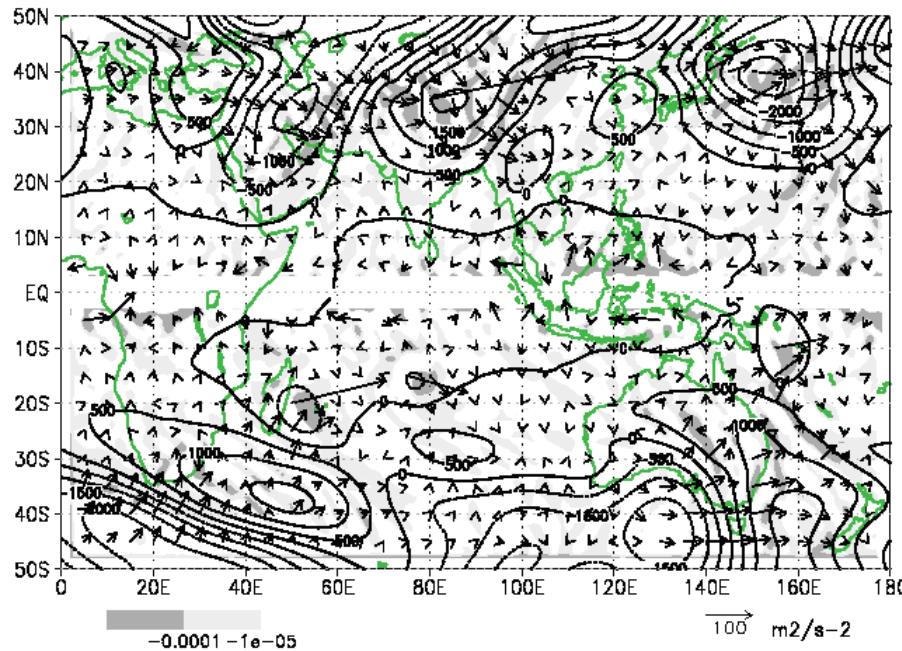


U (>80d), geopotential height & wind (20-80d)

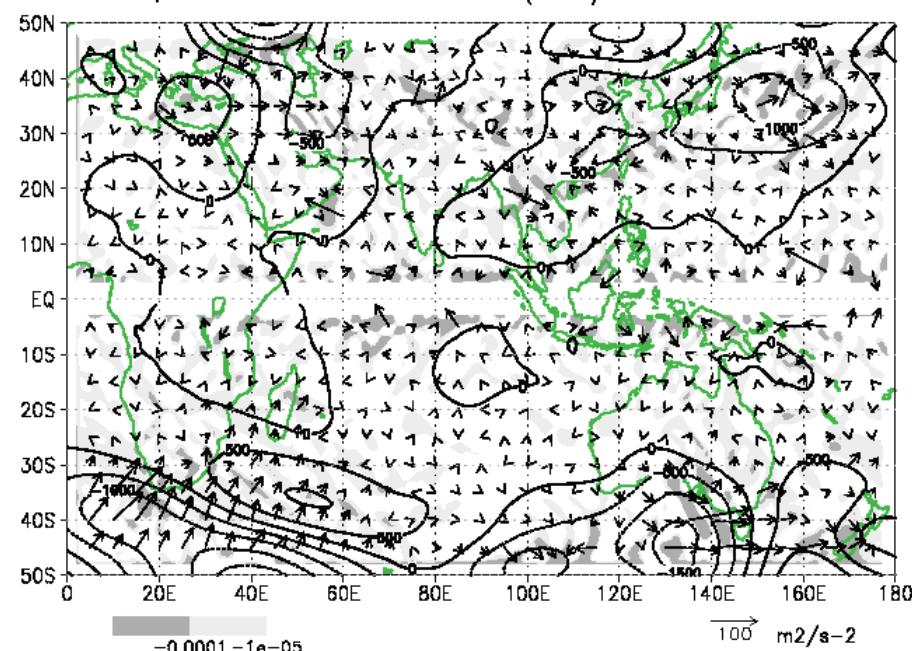
MJO2 (9-14 nov)



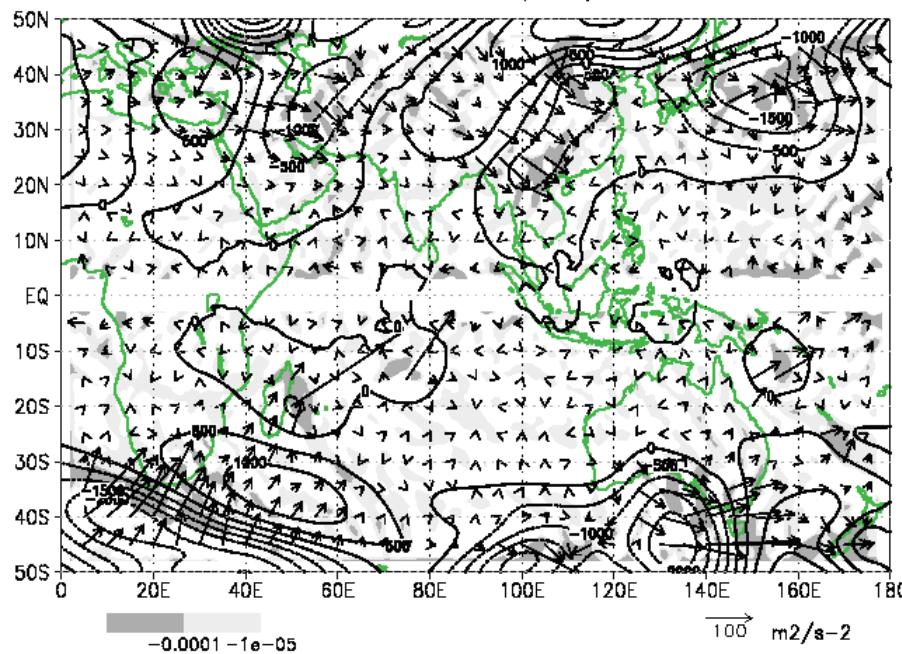
HGT, WAF 200 hPa ERAI 9nov 2011 (<80d)



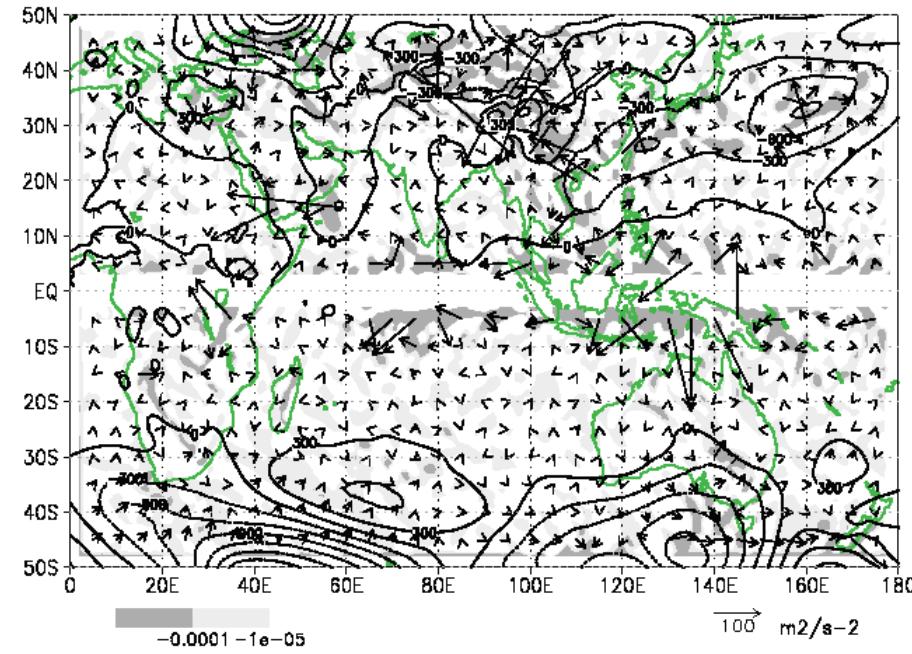
HGT, WAF 600 hPa ERAI 9nov 2011 (<80d)



HGT, WAF 400 hPa ERAI 9nov 2011 (<80d)

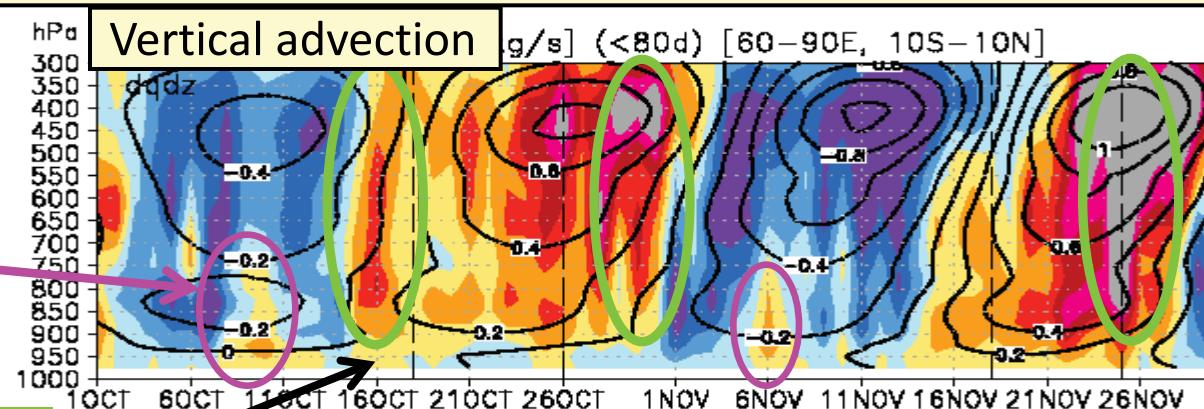


HGT, WAF 850 hPa ERAI 9nov 2011 (<80d)



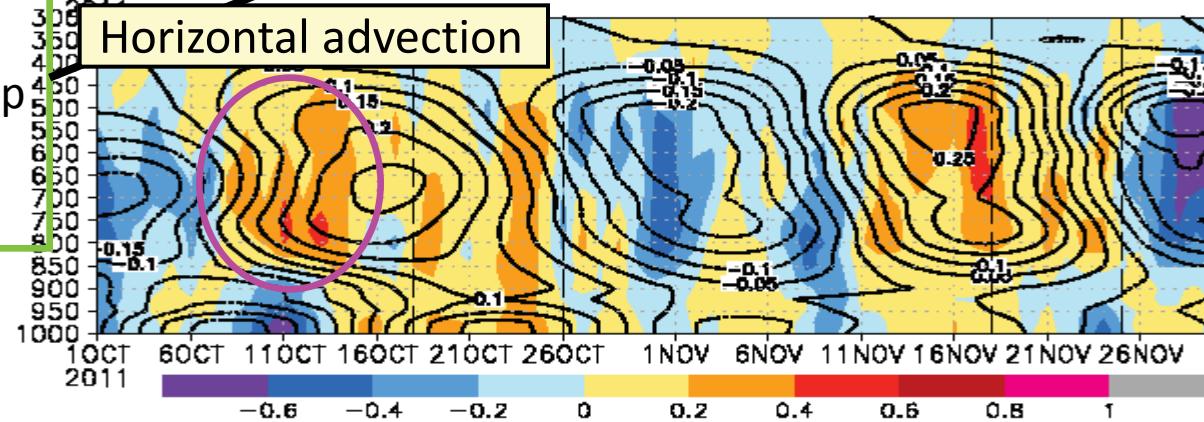
ERA-Interim Moisture advection ISO + high-freq (<80d) 60-90E, 10S-10N

shallow convective events

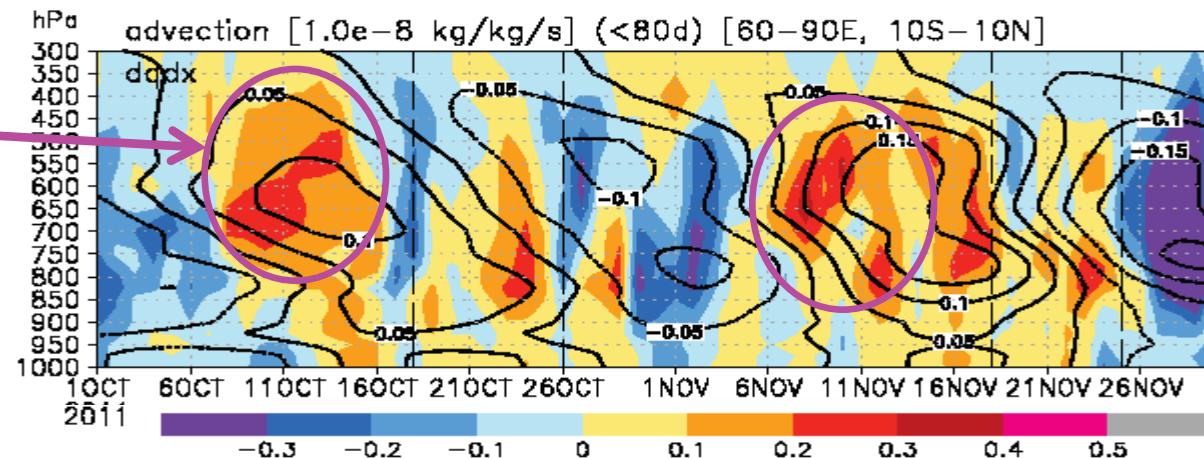


high-freq convection pronounced later stages

Synchronization of high-freq deep convection (Kelvin ?)



mid-tropospheric zonal moisture advection



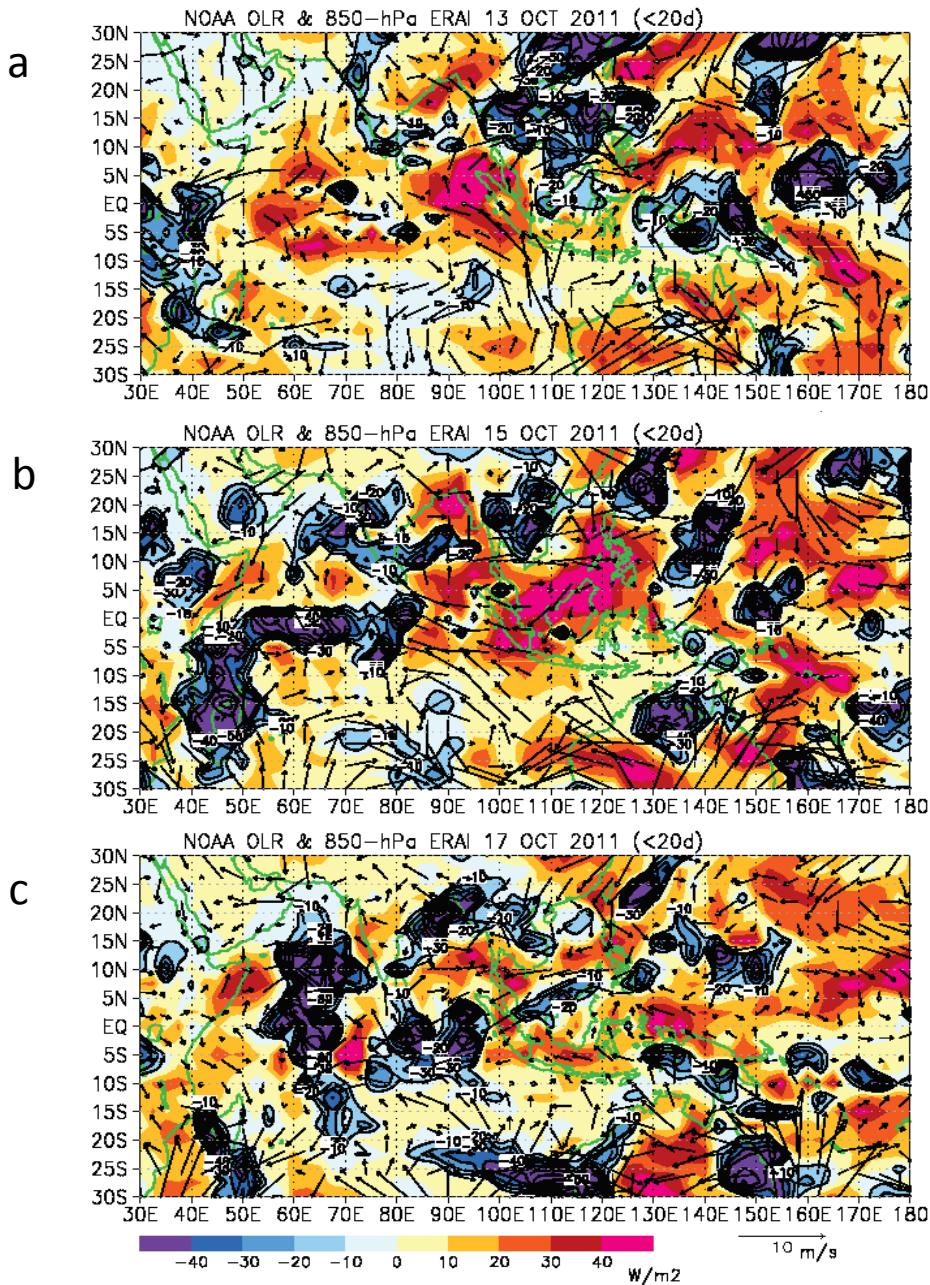
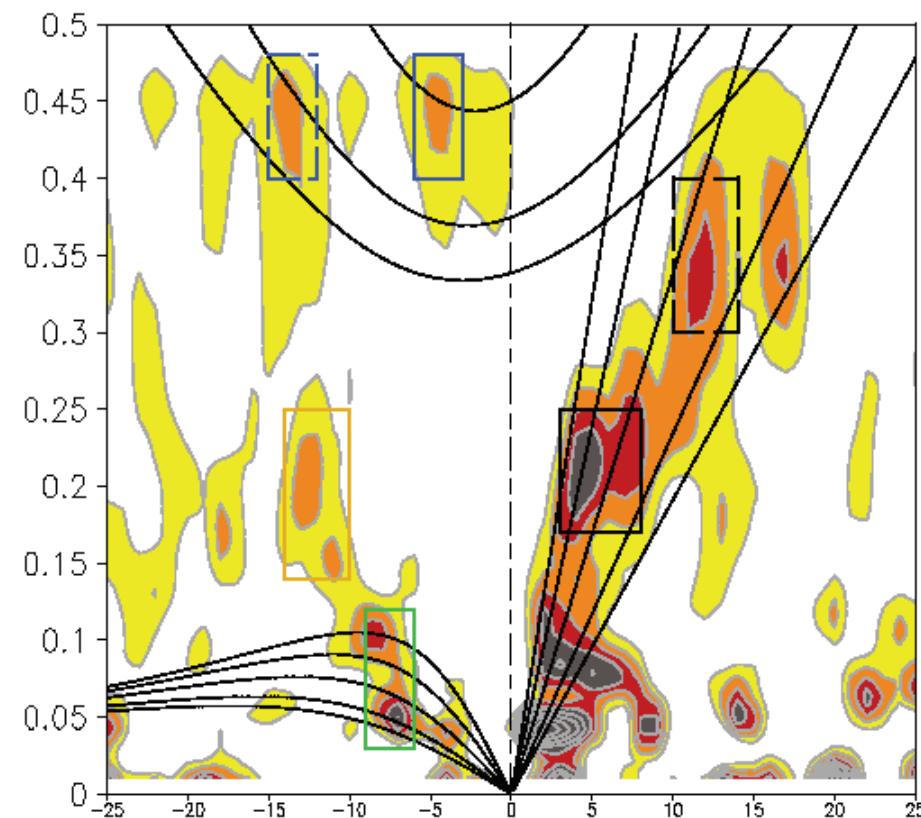


FIG. 18. Same as Fig. 3 except for high-frequency components on (a) 13 (b) 15 and (c) 17 October 2011

Spector analysis of NOAA OLR (15S-15N)

Combined Fourier-Wavelet Transform (CFWT) Kikuchi (2013)

a



b

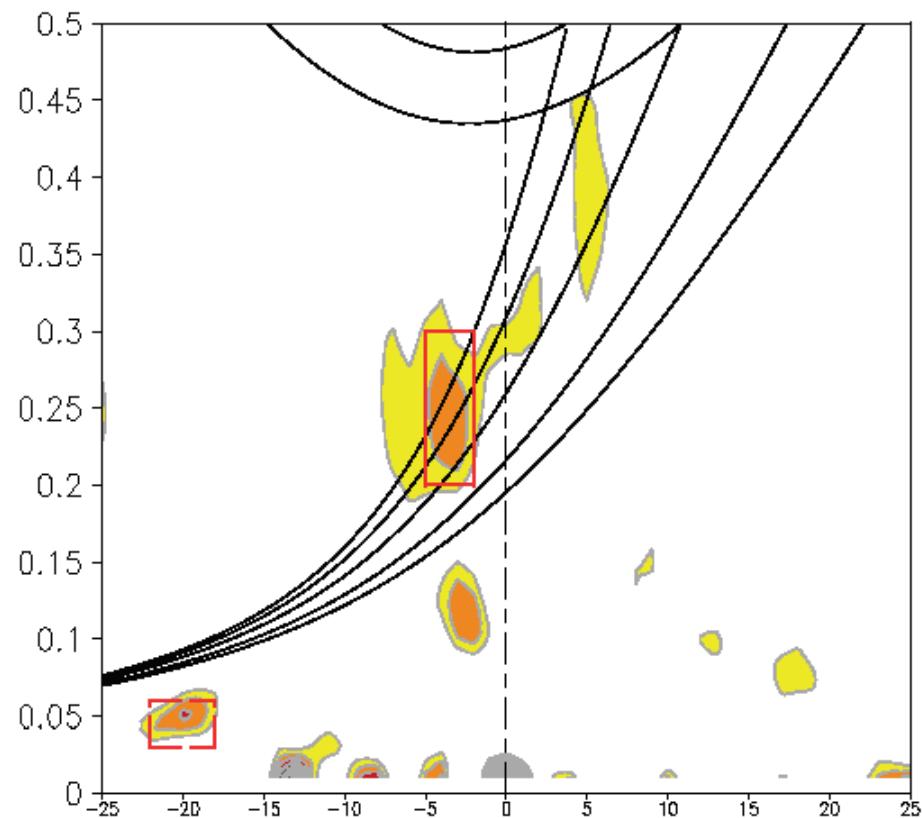


FIG. 19. Zonal wavenumber-frequency power spectrum of NOAA OLR for the equatorially (a) symmetric and (b) asymmetric components based on the CFWT, averaged in 15° S– 15° N and normalized by the background spectrum (Kikuchi 2013)

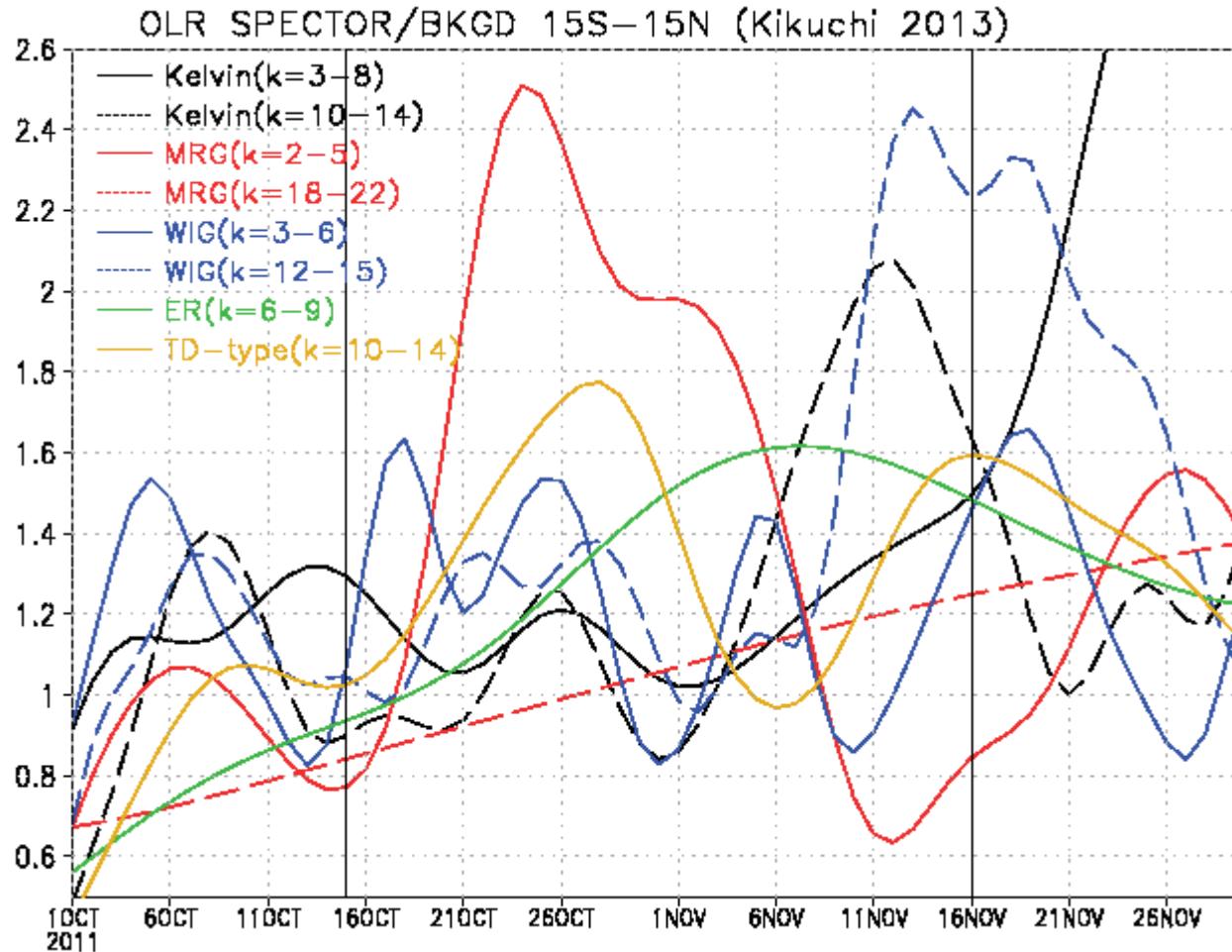


FIG. 20. Time series of the power of NOAA OLR in the specific zonal wavenumber-frequency domains marked in Fig. 19: Equatorial Kelvin waves (black lines), mixed Rossby-gravity wave (red lines), westward inertia-gravity waves (blue lines), equatorial Rossby waves (green line), and tropical relevation type waves (yellow line). Solid (dashed) lines indicate lower (higher) zonal wavenumber components.

Decomposition of horizontal moisture advection

Zhao et al. 2013

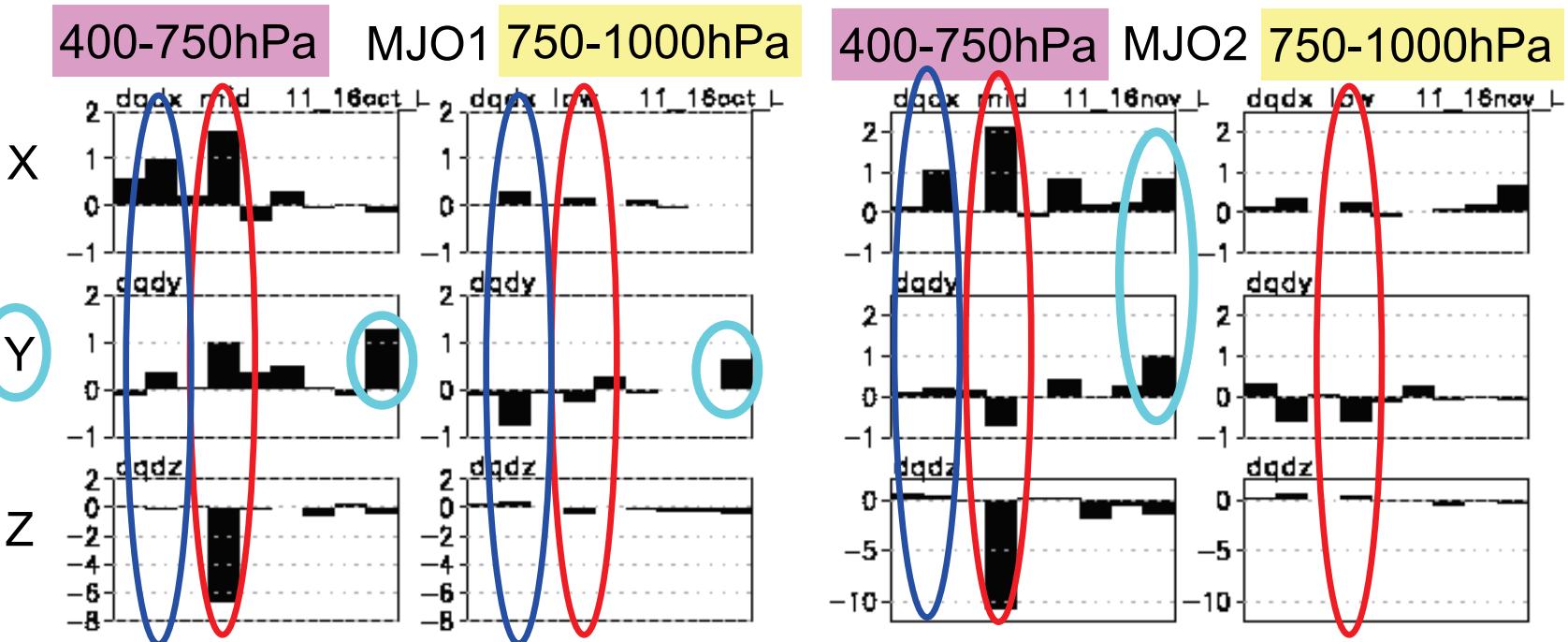
$$\left(\frac{\partial q}{\partial t} \right) \neq (-V \nabla \bar{q})' + -\omega \frac{\partial q}{\partial p} + -\frac{Q_2}{L},$$

60-90E, 10S-10N

$$u = \bar{u} + u' + u^*, \quad v = \bar{v} + v' + v^* \quad q = \bar{q} + q' + q^*$$

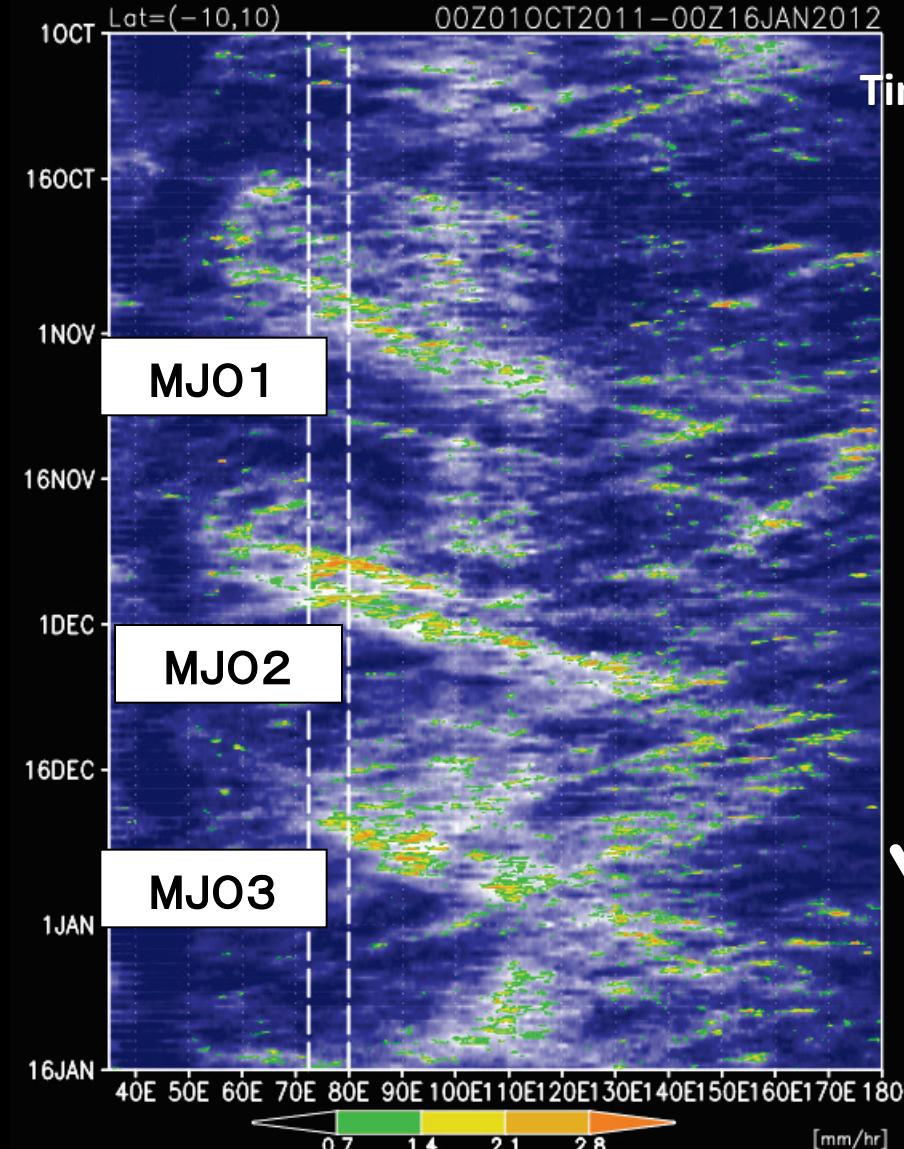
overbar: >80 day (basic state) prime: 20-80 day (ISO) asterisk: < 20 day (high-freq)

$(-\bar{V} \cdot \nabla \bar{q})'$	$(-\bar{V} \cdot \nabla q^*)'$	$(-\bar{V} \cdot \nabla q^*)'$	$(-\bar{V}' \cdot \nabla \bar{q})'$	$(-\bar{V}' \cdot \nabla q^*)'$	$(-\bar{V}' \cdot \nabla q^*)'$	$(-\bar{V}^* \cdot \nabla \bar{q})'$	$(-\bar{V}^* \cdot \nabla q^*)'$	$(-\bar{V}^* \cdot \nabla q^*)'$
1	2	3	4	5	6	7	8	9

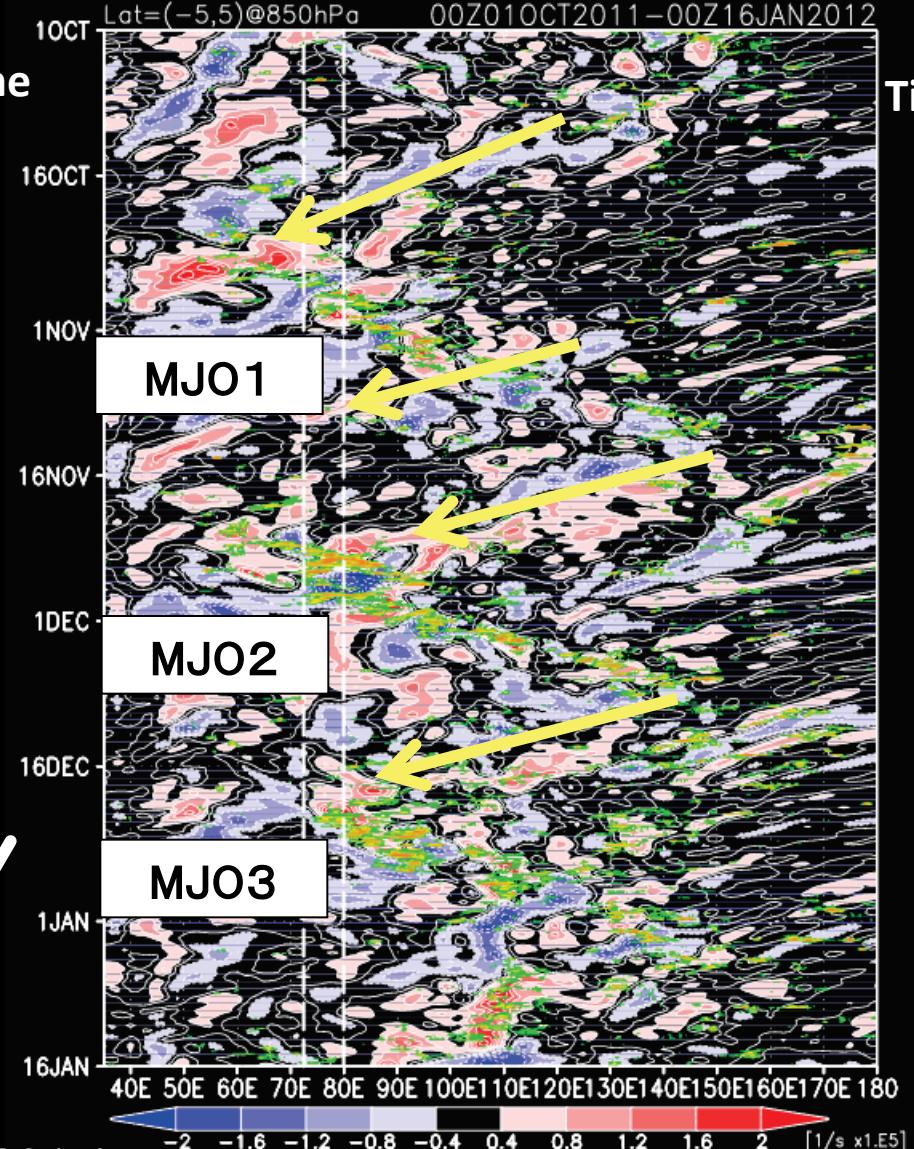


high-freq x high-freq in the meridional component

TRMM, MTSAT(10S–10N)



850 hPa vorticity (5S–5N)



Yasunaga et al. (2014)

Summary and future work

- Horizontal advection of mid-tropospheric moisture by easterly anomalies were detected both in MJO1 and in MJO2.
- The easterly anomalies appear to be associated with extratropical wave trains that intruded into the equatorial region. This suggests contribution of dry dynamics to the moist processes in MJO.
- The different conditions between the two cases (drier condition in MJO1 than in MJO2) may lead to the difference in the vertical evolution (slower in MJO1 than in MJO2) and the contributions from high-frequency disturbances (e.g., equatorial waves, transient eddies).
- Evaluation of high-frequency effects (e.g., diabatic heating, CCEWs) using outputs of global cloud-system-resolving simulations is underway.

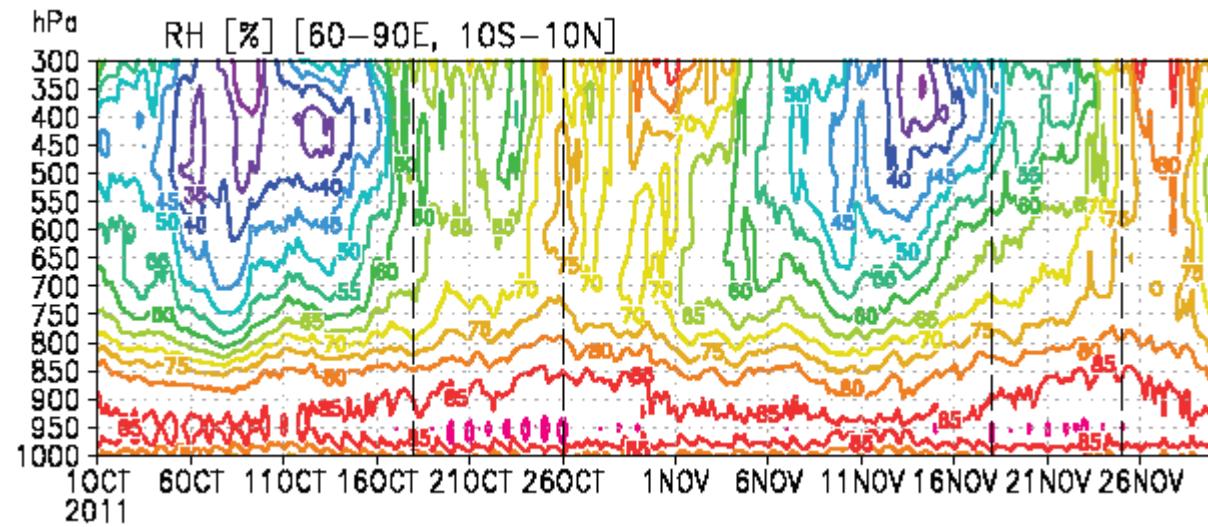
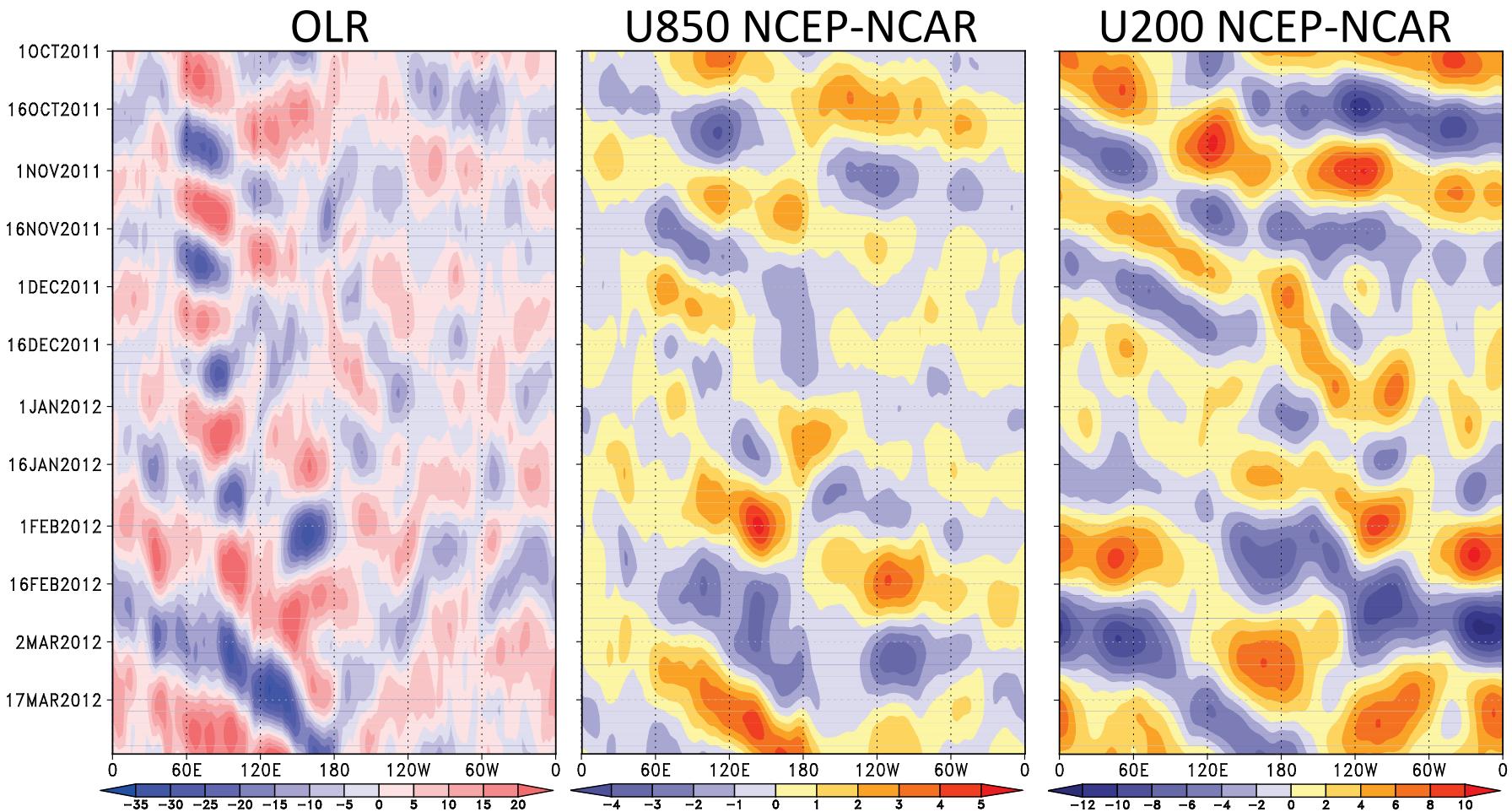


FIG. S2. Time-height sections of relative humidity (ERA-interim) averaged in the 60-90E, 10S-10N domain.

Backup slides

25-90 bandpass filtered fields 15S-15N



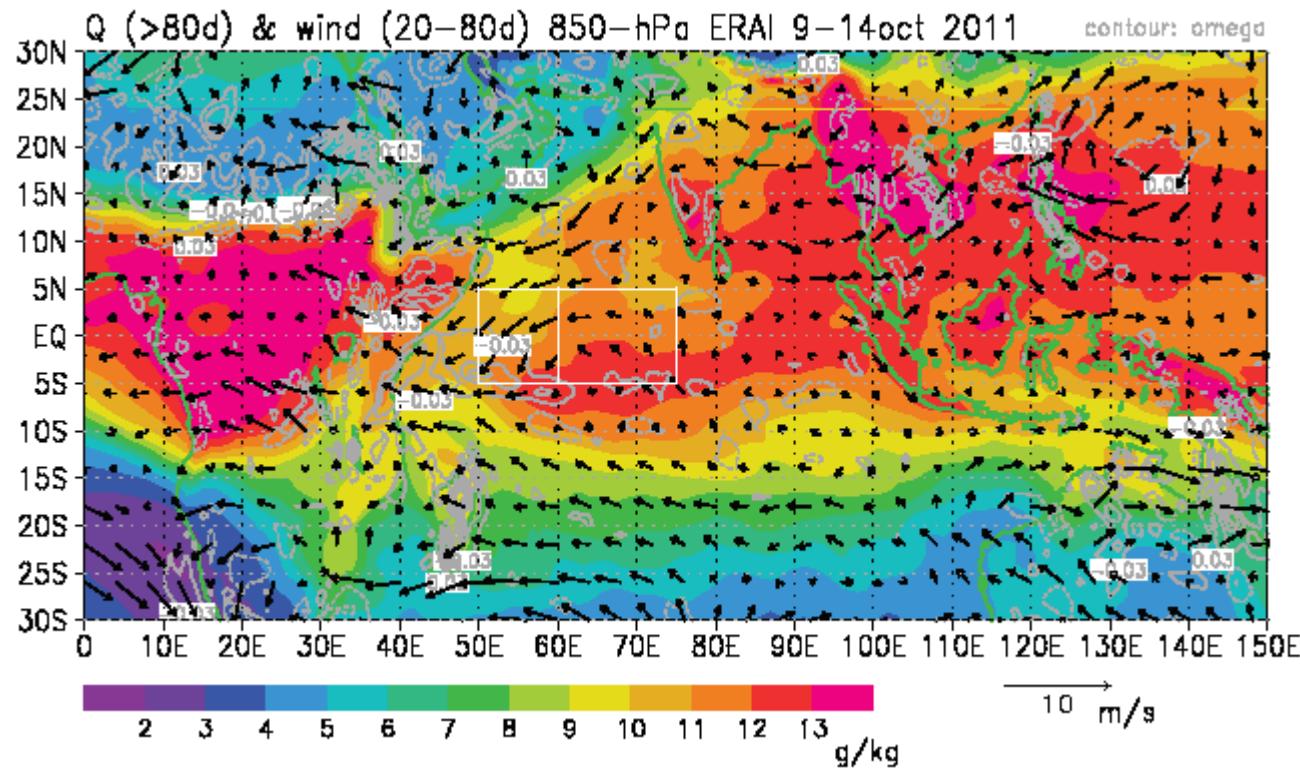


FIG. 12. Horizontal distributions of low-frequency (>80-day) moisture and intraseasonal (20-80 day) wind vectors (a) at 600-hPa in the preconditioning stage and (b) at 850 hPa in the convection initiation period of MJO1. Squares indicate MJO convection initiation region 1.

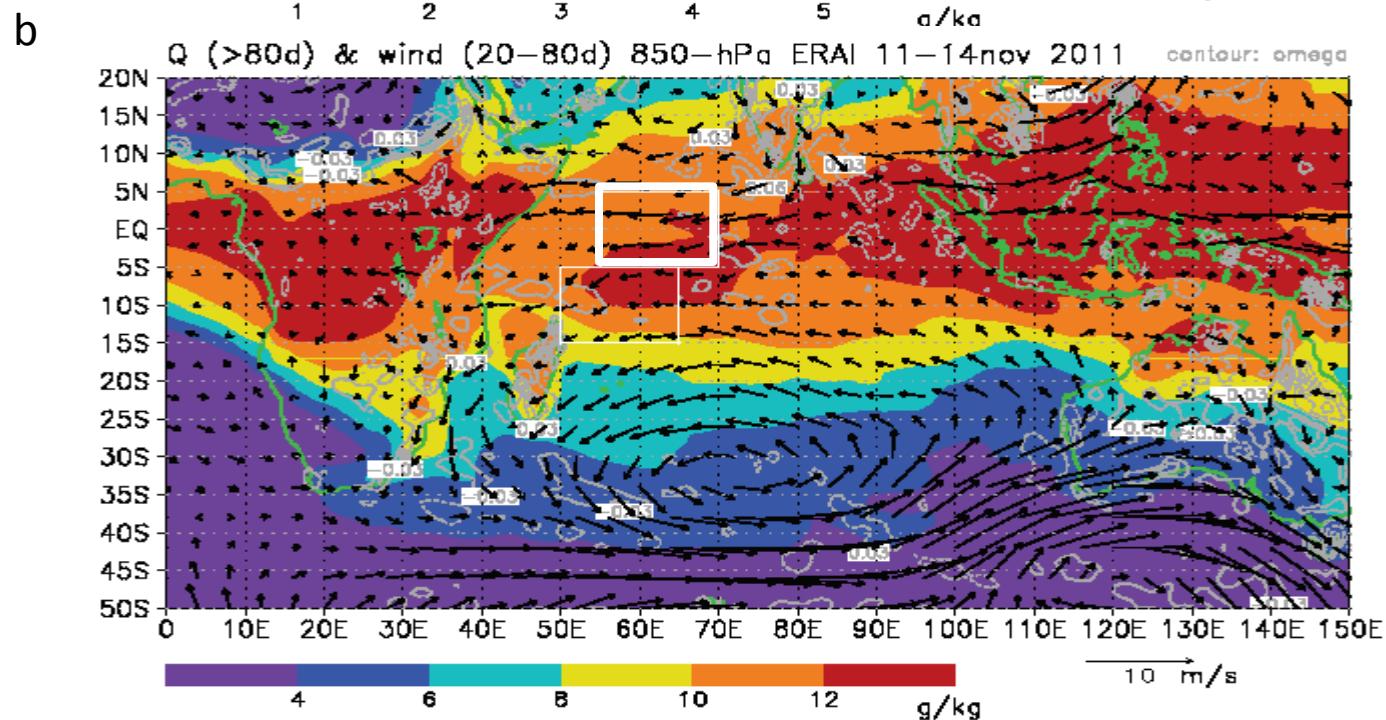
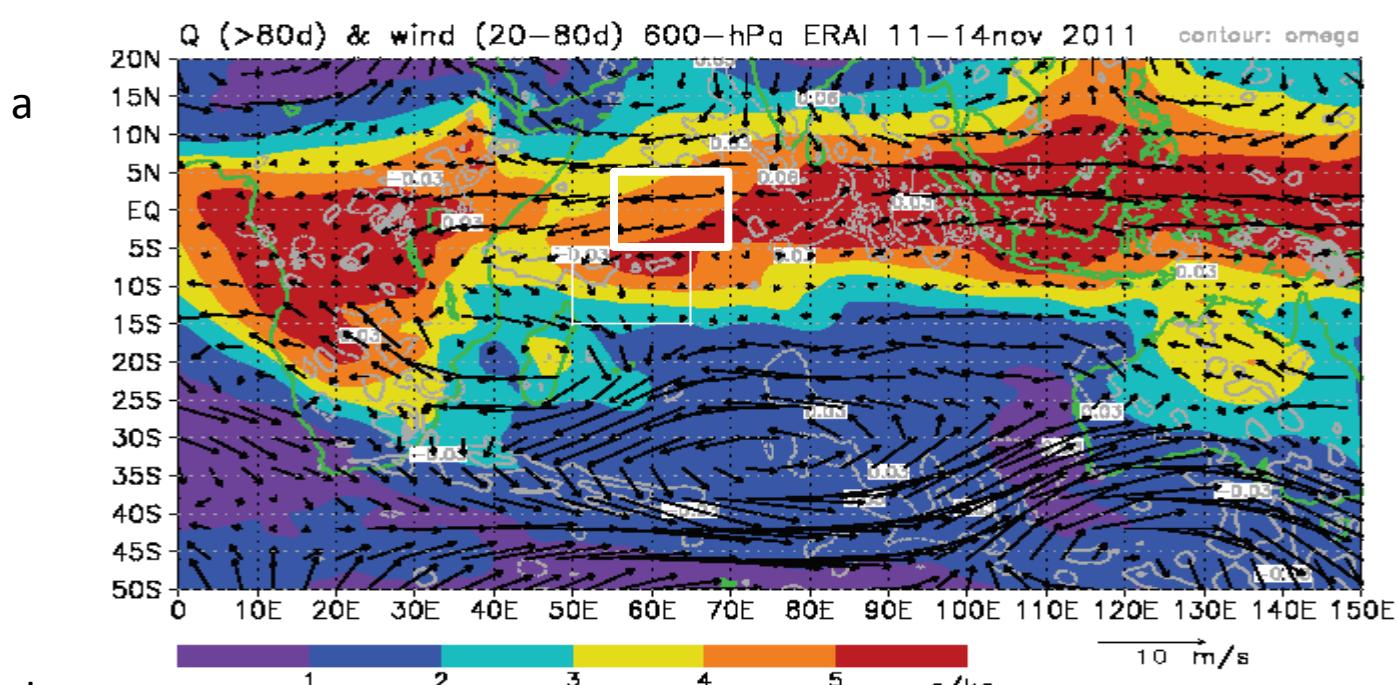


FIG. 16. Same as Fig. 12 except for MJO2. Squares indicate MJO convection initiation region 2.

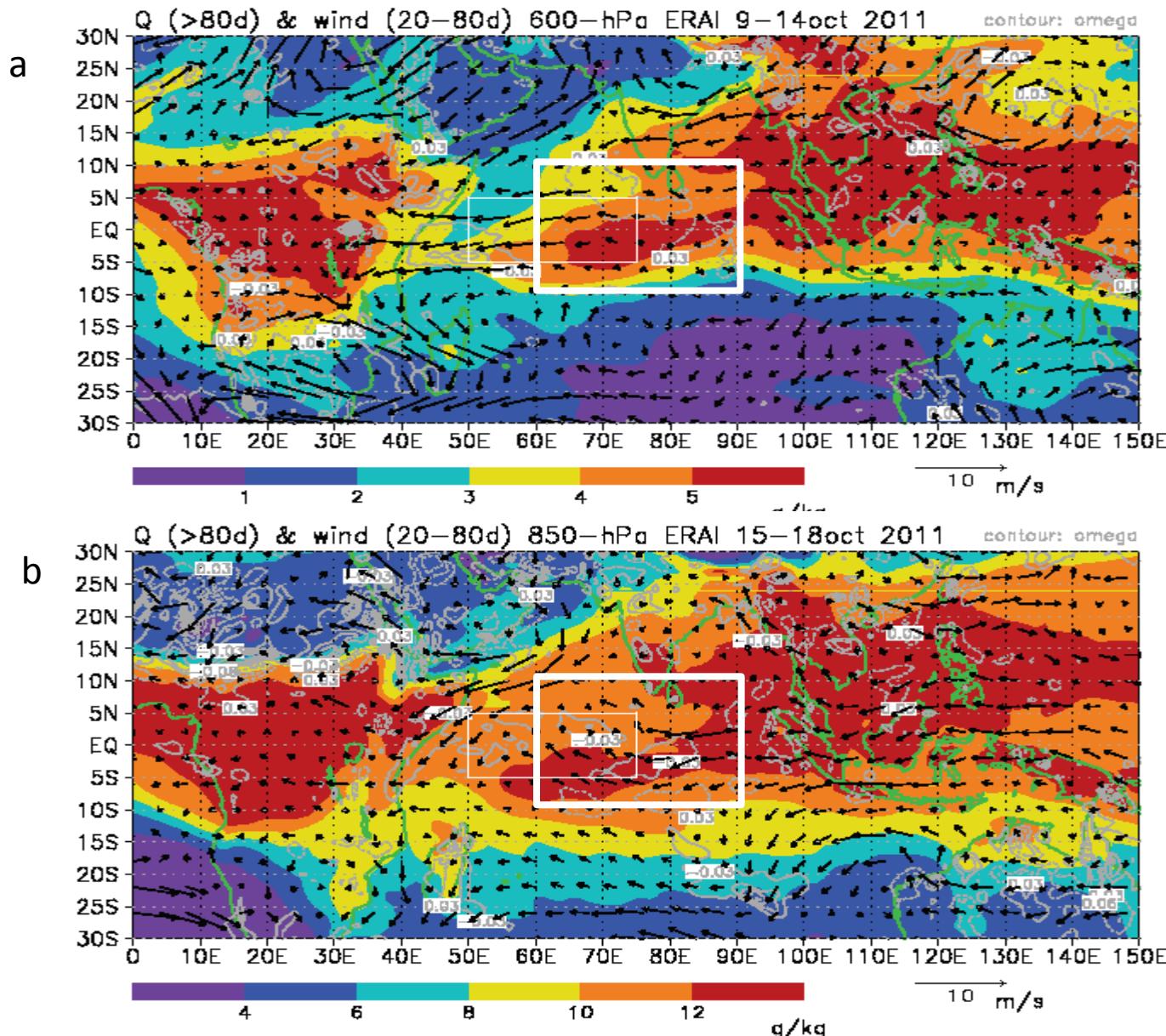


FIG. 12. Horizontal distributions of low-frequency (>80-day) moisture and intraseasonal (20-80 day) wind vectors (a) at 600-hPa in the preconditioning stage and (b) at 850 hPa in the convection initiation period of MJO1. Squares indicate MJO convection initiation region 1.

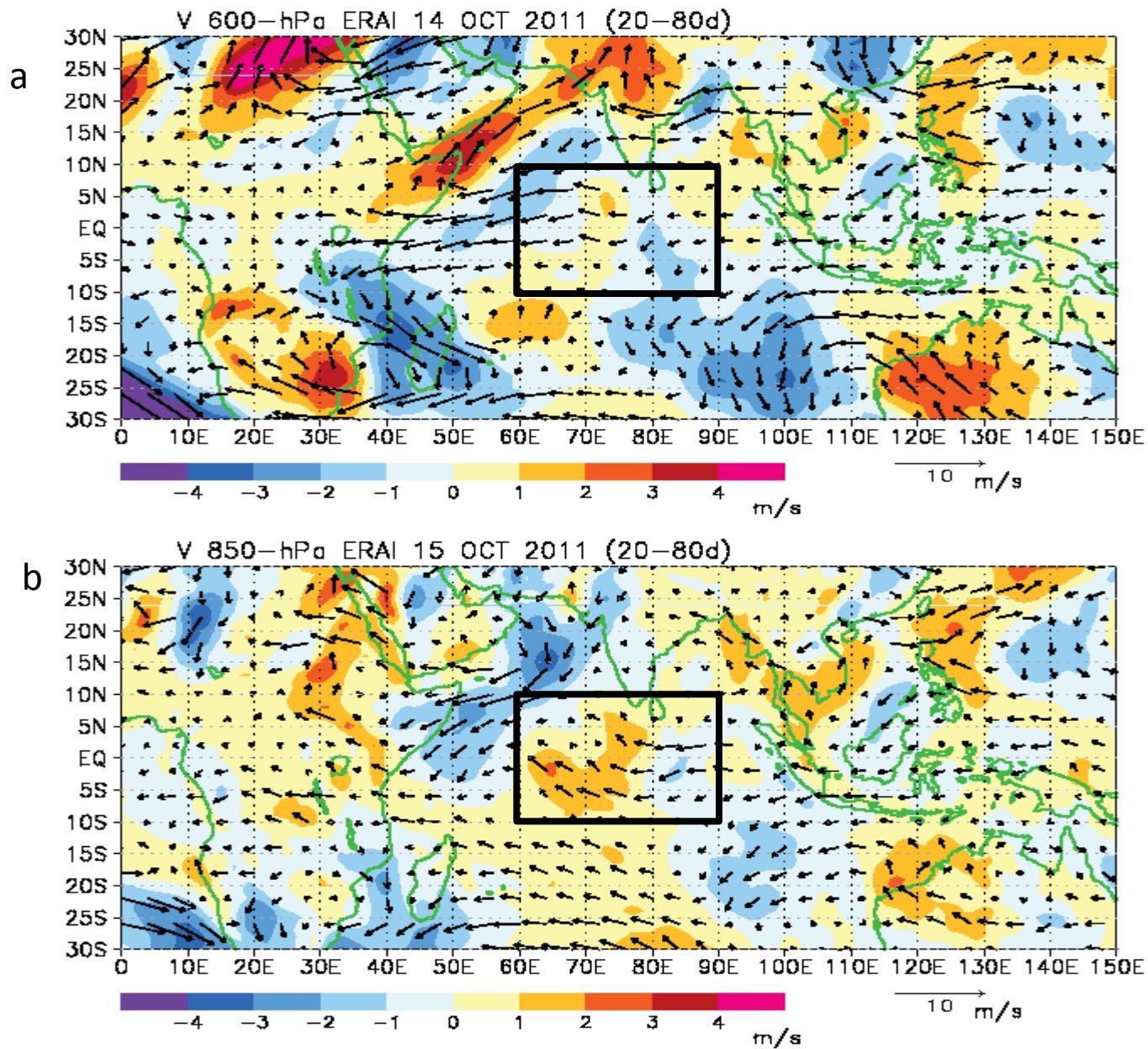
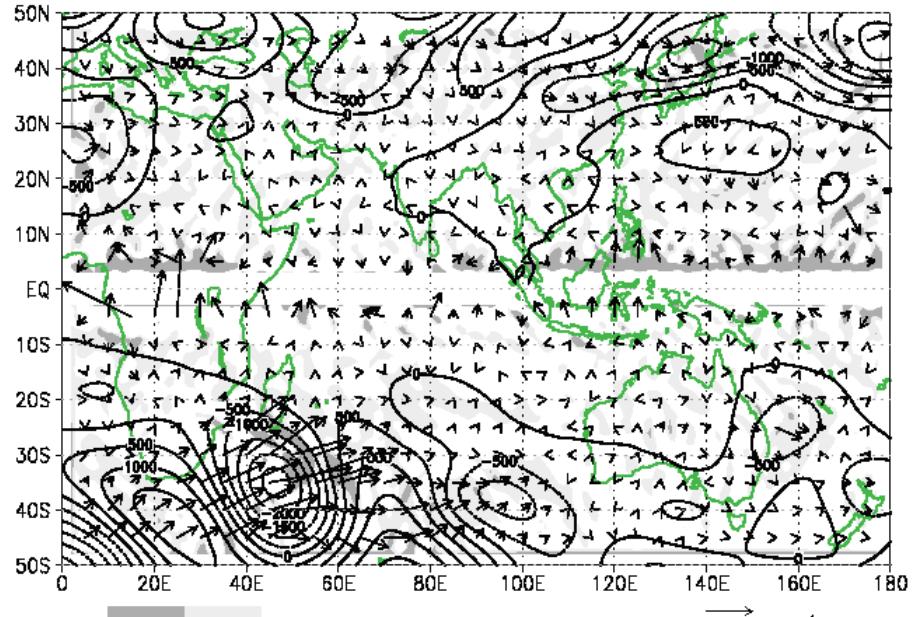
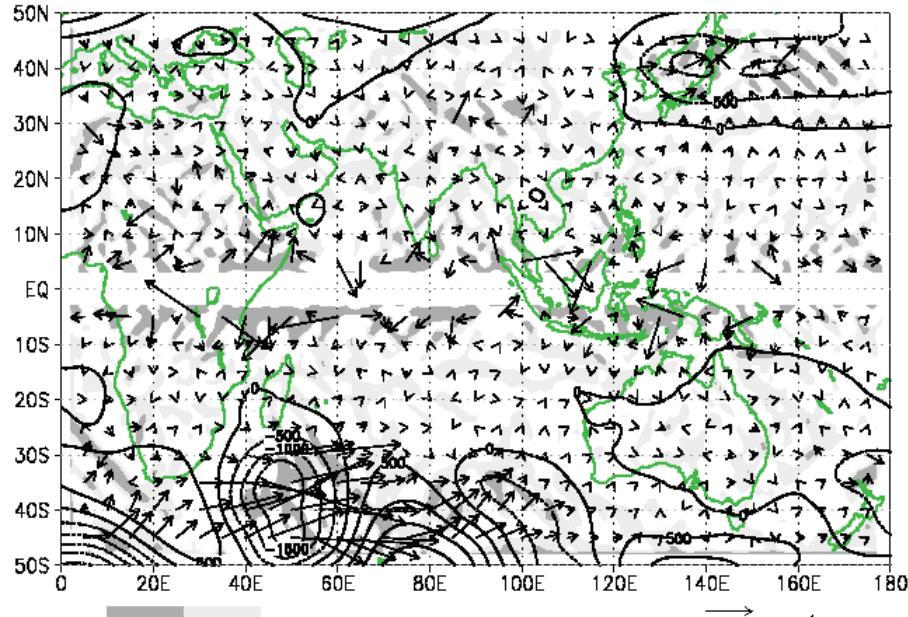


FIG. 17. Same as Fig. 12 except for intraseasonal meridional velocity at (a) 600 hPa on 14 October and (b) 850 hPa on 15 October.

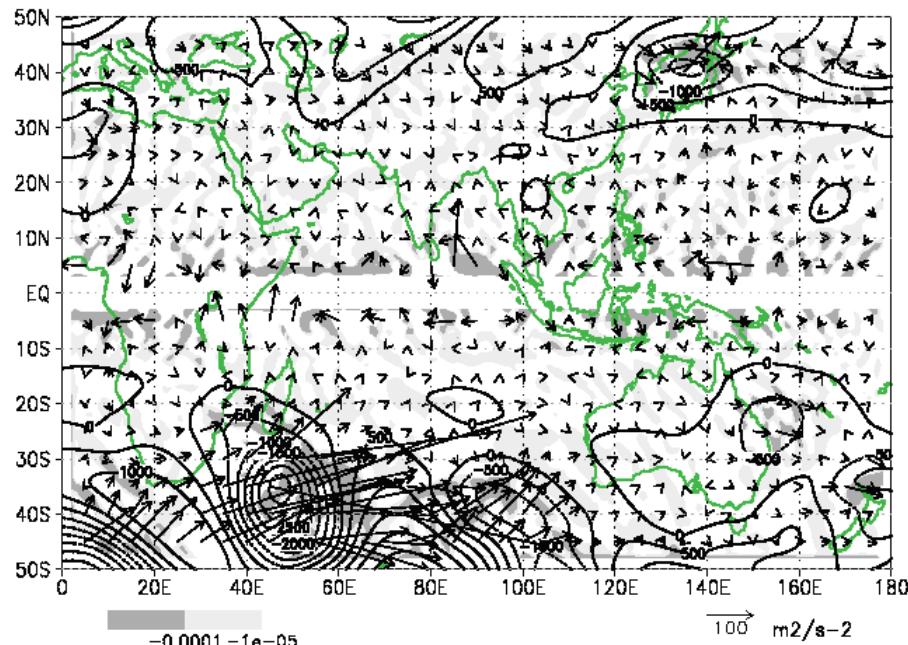
HGT, WAF 200 hPa ERAI Oct 2011 (<80d)



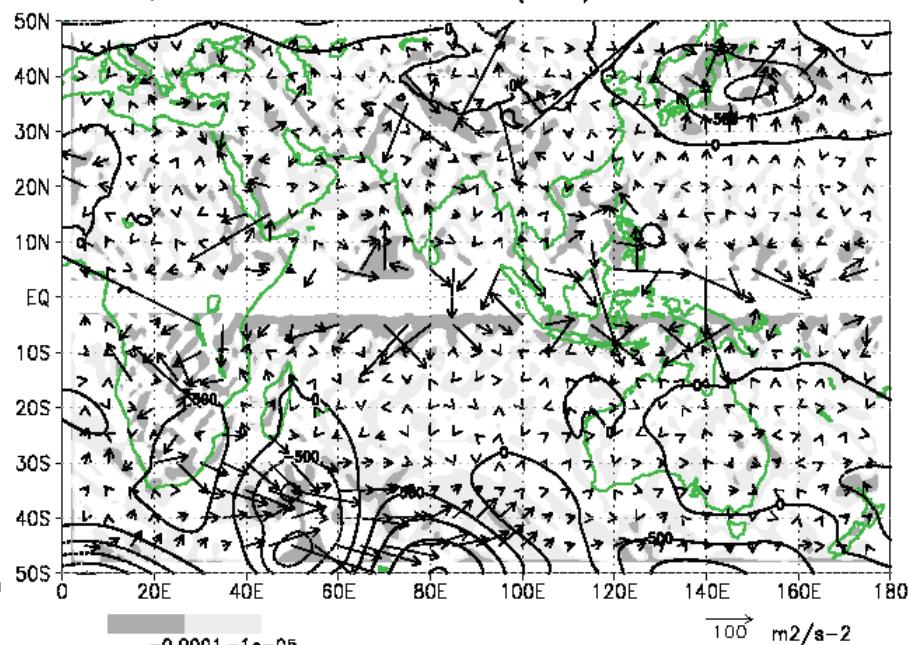
HGT, WAF 600 hPa ERAI Oct 2011 (<80d)



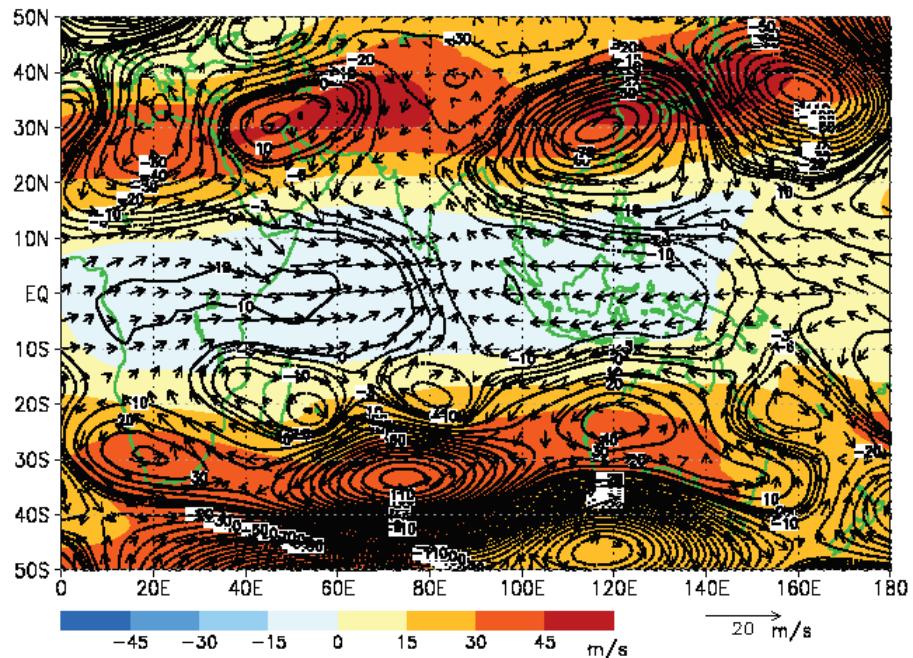
HGT, WAF 400 hPa ERAI Oct 2011 (<80d)



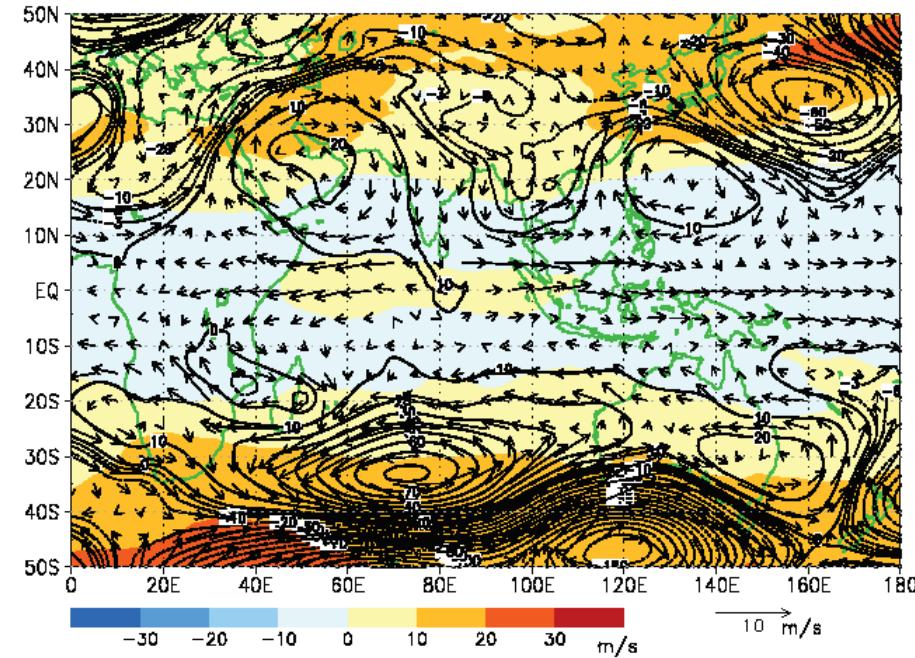
HGT, WAF 850 hPa ERAI Oct 2011 (<80d)



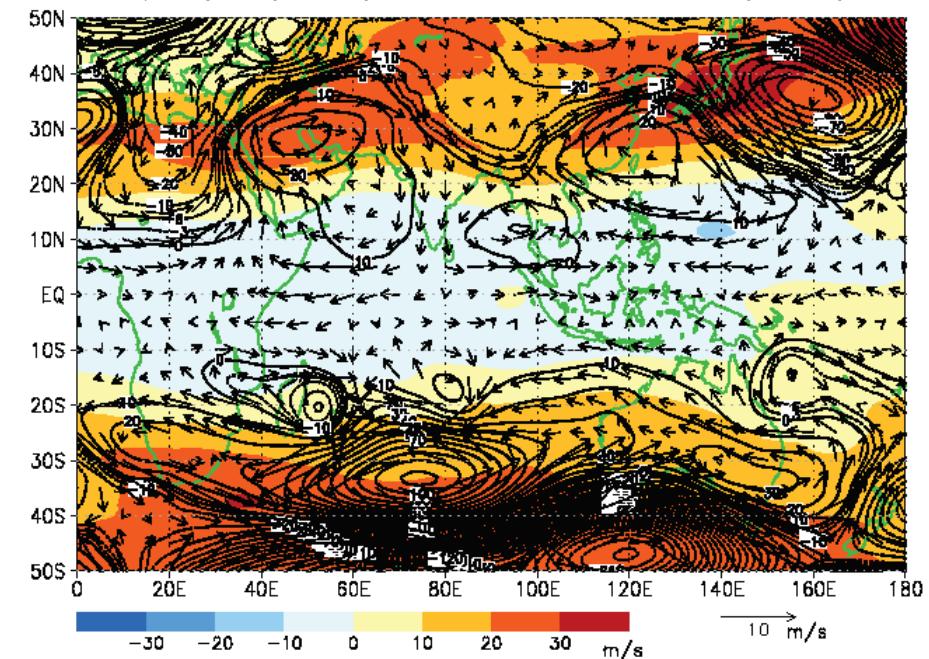
U(>80d) HGT(20–80d) 200 hPa ERAI 11–14nov 2011 (20–80d)



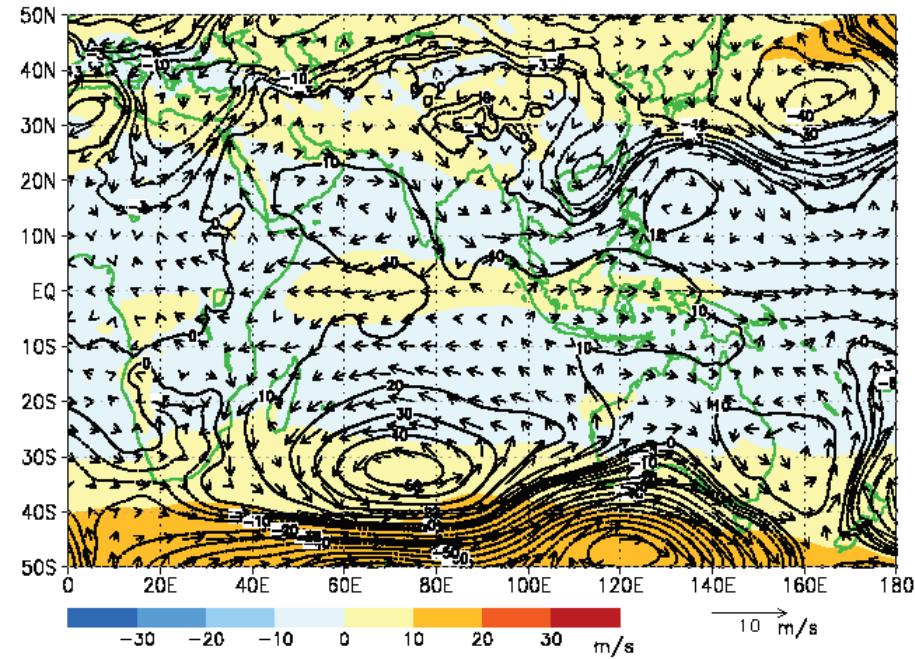
U(>80d) HGT(20–80d) 600 hPa ERAI 11–14nov 2011 (20–80d)



U(>80d) HGT(20–80d) 400 hPa ERAI 11–14nov 2011 (20–80d)



U(>80d) HGT(20–80d) 850 hPa ERAI 11–14nov 2011 (20–80d)



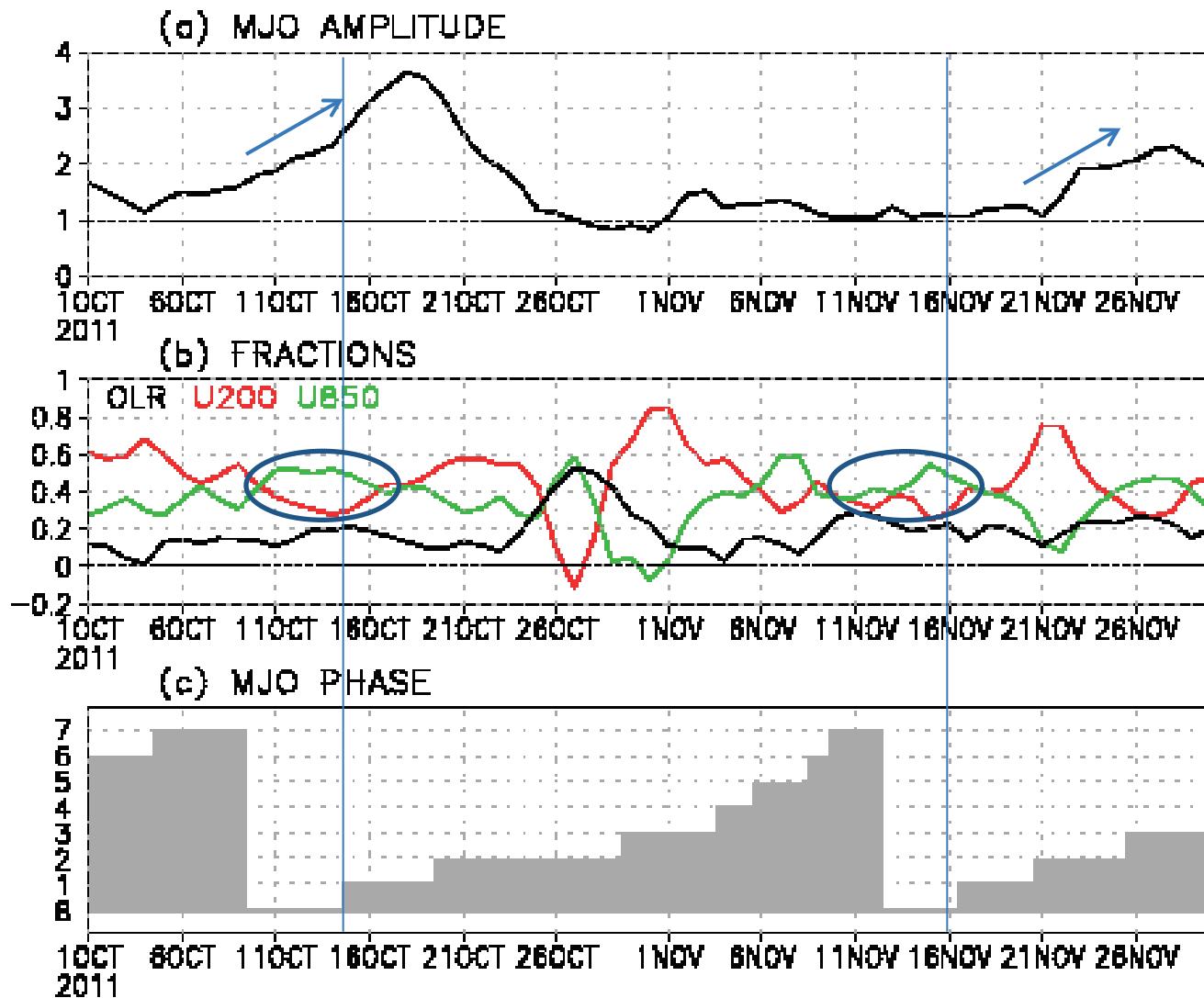


FIG. 2. Time series of (a).MJO amplitude, (b) fractions of OLR (black line), U850 (green line), and U200 (red line), (c) MJO phase based on WH2004. NCEP=NCAR reanalysis data are used. (CAWR web site)

Decomposition of horizontal moisture advection

Zhao et al. 2012

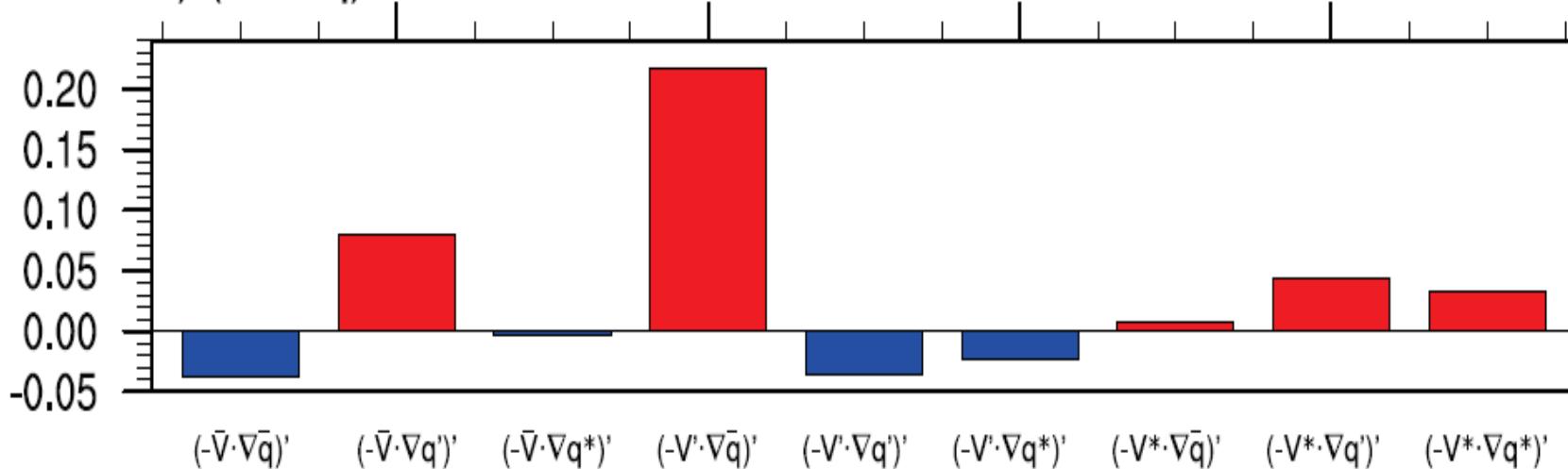
$$u = \bar{u} + u' + u^*, \quad v = \bar{v} + v' + v^* \quad q = \bar{q} + q' + q^*$$

a bar: low-frequency background state (LFBS, > 90 days)

a prime: intraseasonal (20-90-day) component

a star: higher frequency component (< 20 days)

b) $(-\nabla \cdot \nabla q')$



→ The largest terms are the advection of **the mean moisture by the MJO flow** ($-V' \cdot \nabla \bar{q}$) and the advection of **MJO perturbation moisture by the background flow** ($-\bar{V} \cdot \nabla q'$)

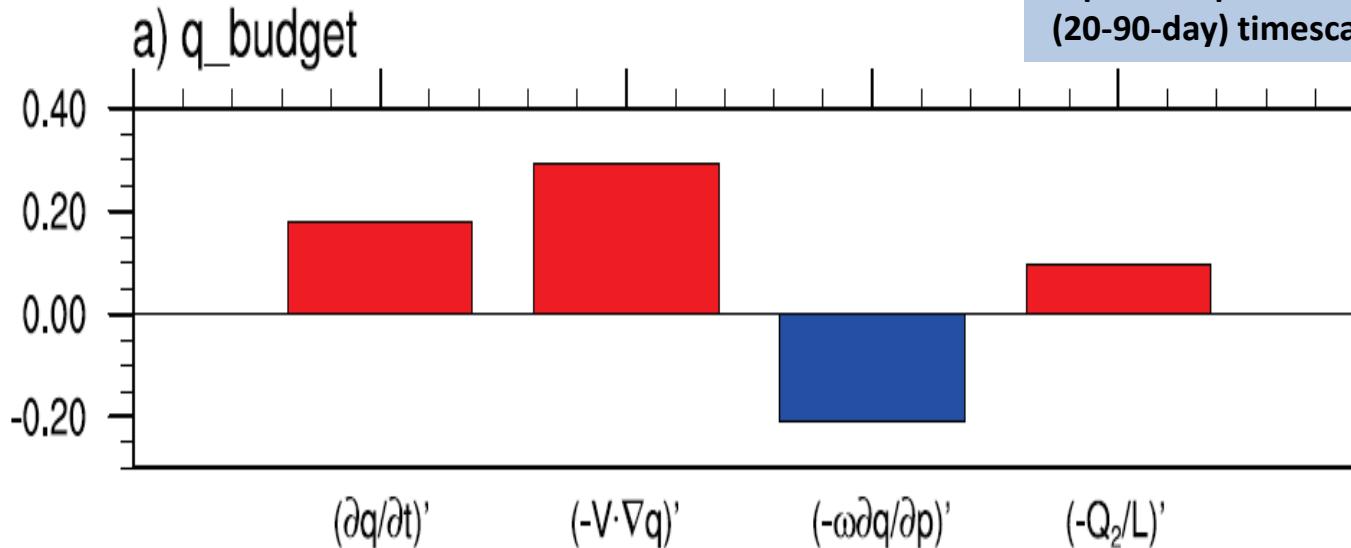
Vertically integrated (1000-700hPa) moisture budget analysis

Zhao et al. 2012

$$(\frac{\partial q}{\partial t})' \neq (-V \cdot \nabla q)' + -\omega \frac{\partial q}{\partial p}' + -\frac{Q_2}{L}',$$

Integrated from 1000 to 700 hPa
during the initiation period from
day -25 to -15 over 50-70E, 20S-EQ

A prima represents intraseasonal
(20-90-day) timescale



From left to right: ISO humidity tendency, horizontal advection, vertical advection, apparent moisture source

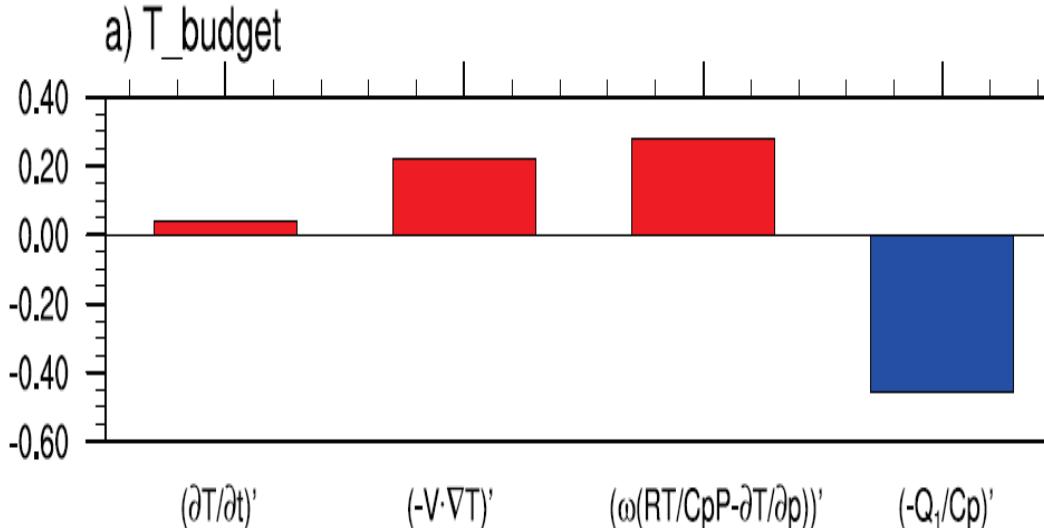
anomalous descending motion => reduces the mean precipitation

=> less condensational heating => more moisture retained in the atmosphere

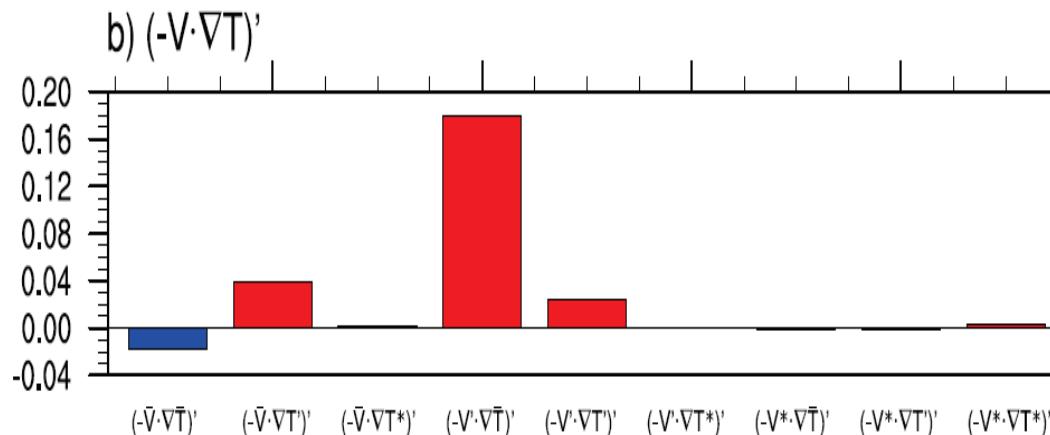
Vertically-integrated heat budget analysis

Zhao et al. 2012

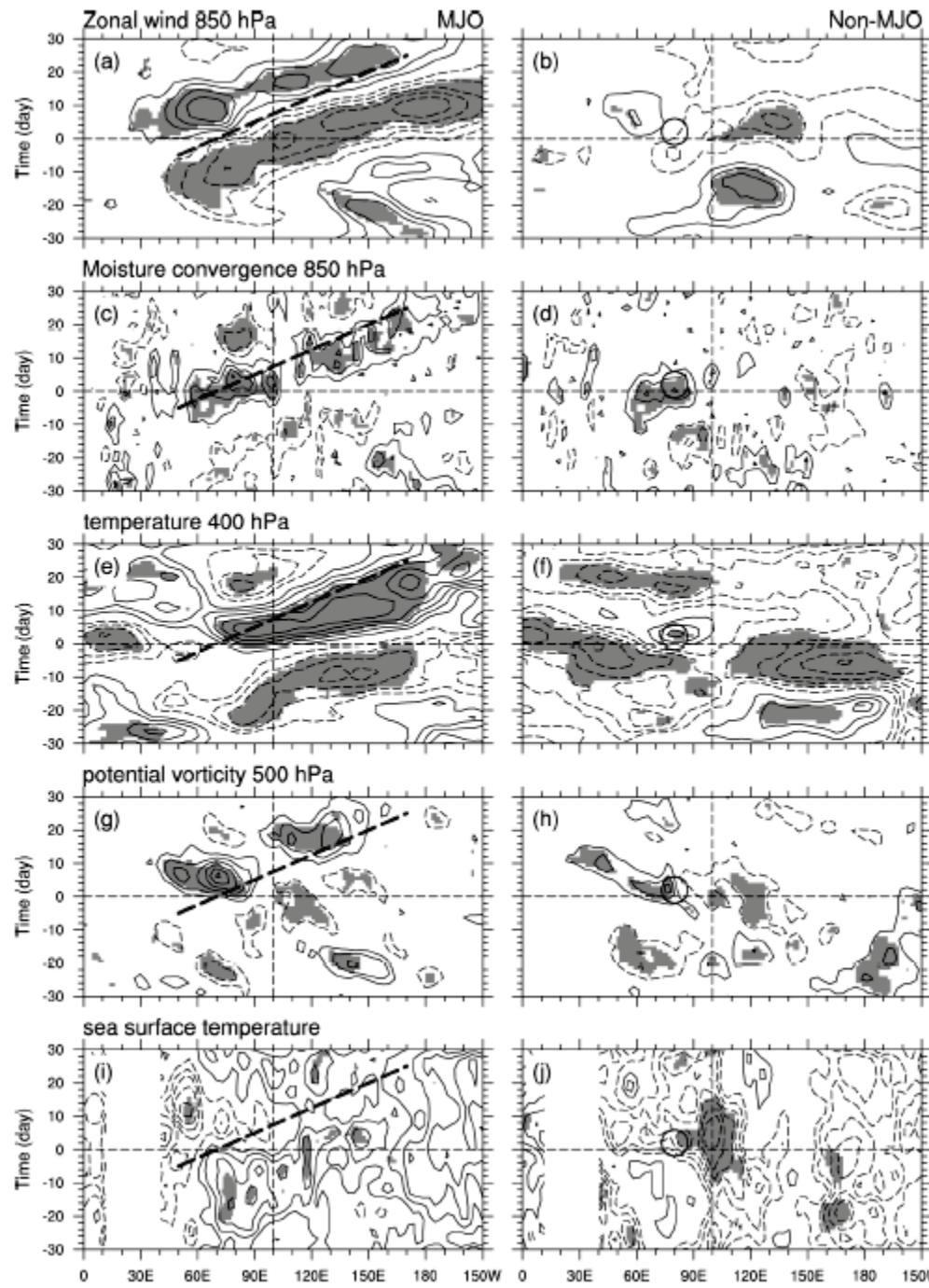
$$(\frac{\partial T}{\partial t})' = (-V \nabla T)' + \omega \frac{RT}{c_p p} - \omega \frac{\partial T}{\partial p}' + (Q_l / c_p)'$$



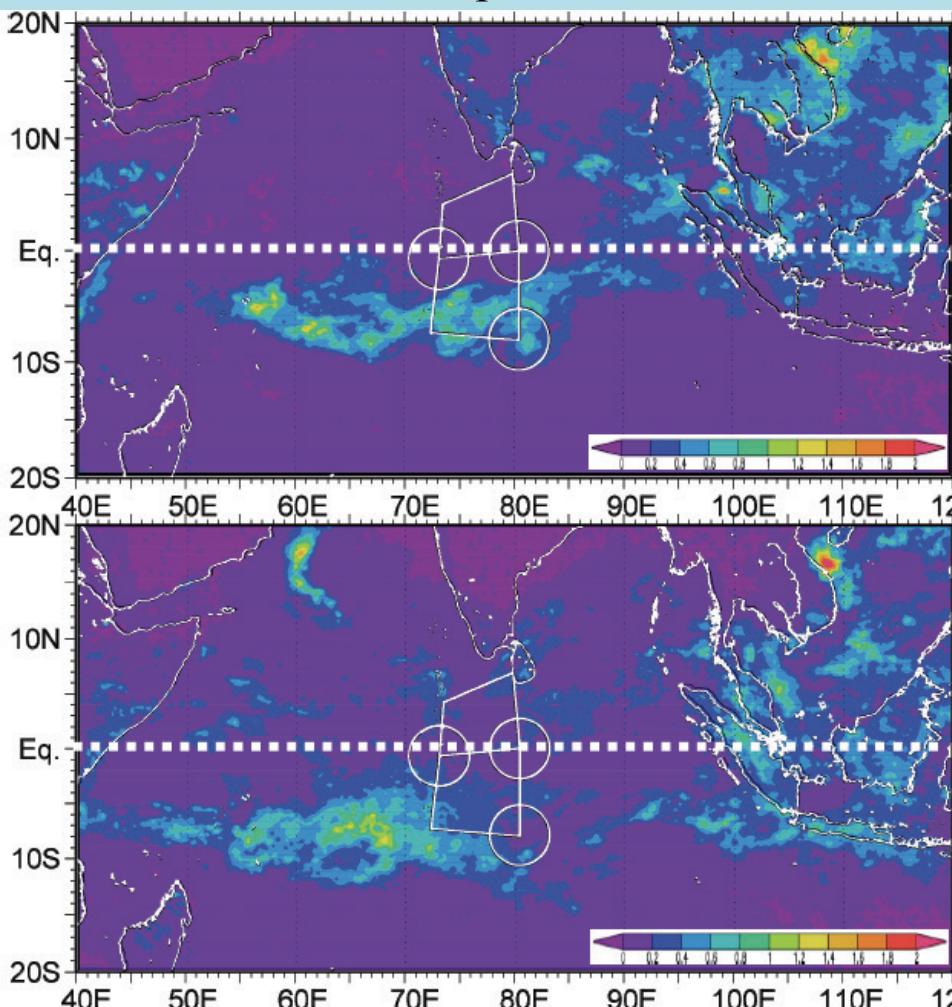
← Descending-induced
adiabatic warming is largely
offset by diabatic cooling.



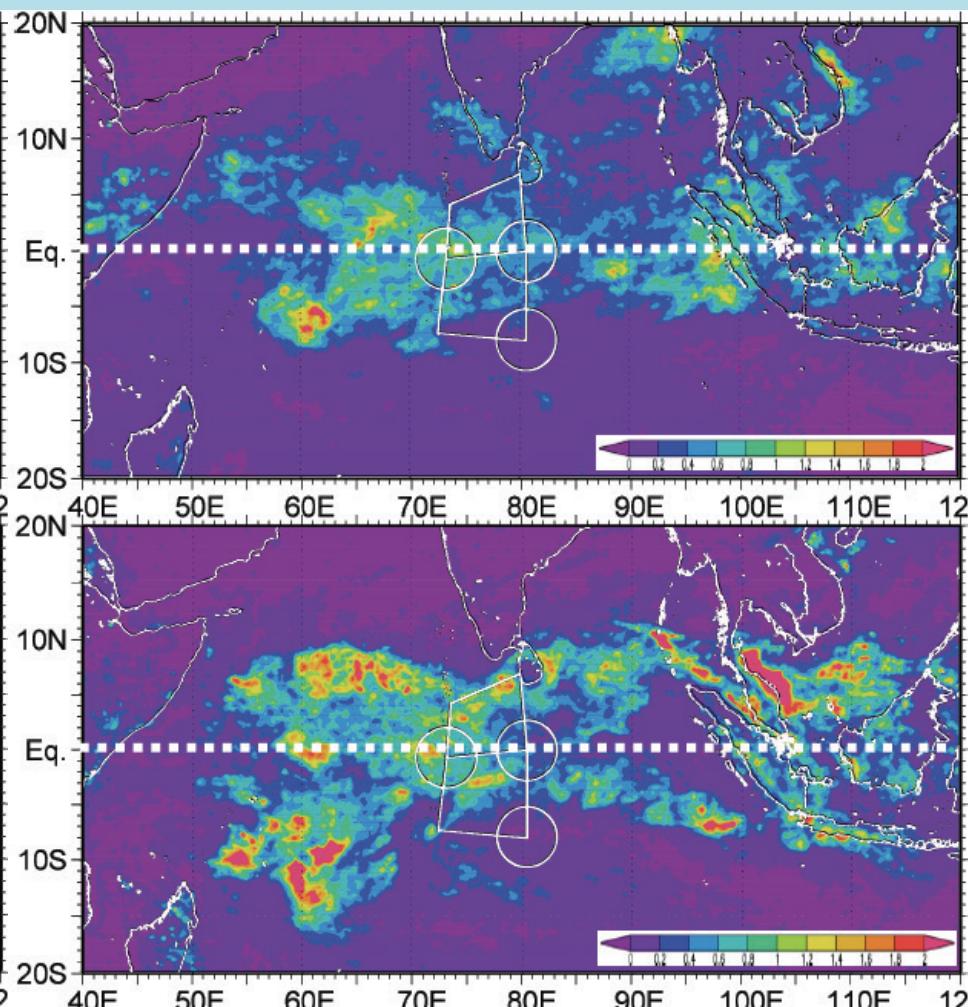
Horizontal advection is
primarily attributed to
advection of the mean
temperature by MJO flow.



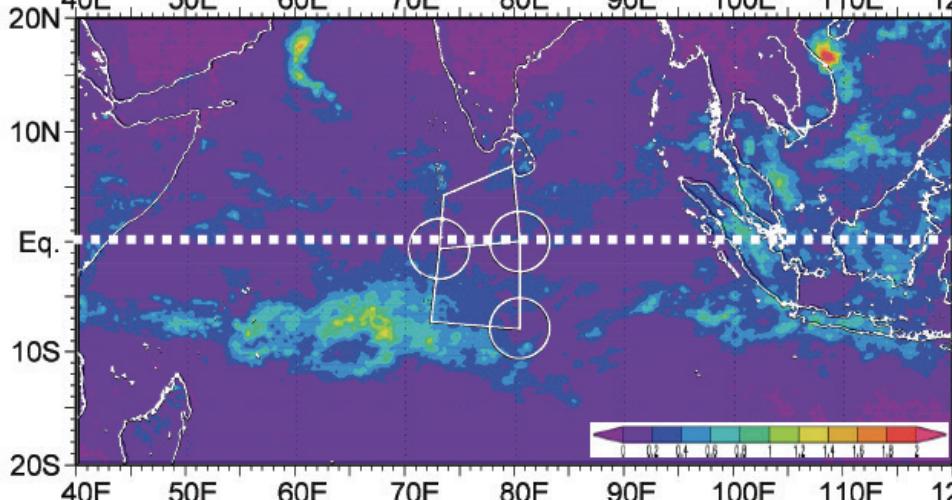
MJO-1 Sept 30 – Oct 15



MJO-1 Oct 15 – Oct 25



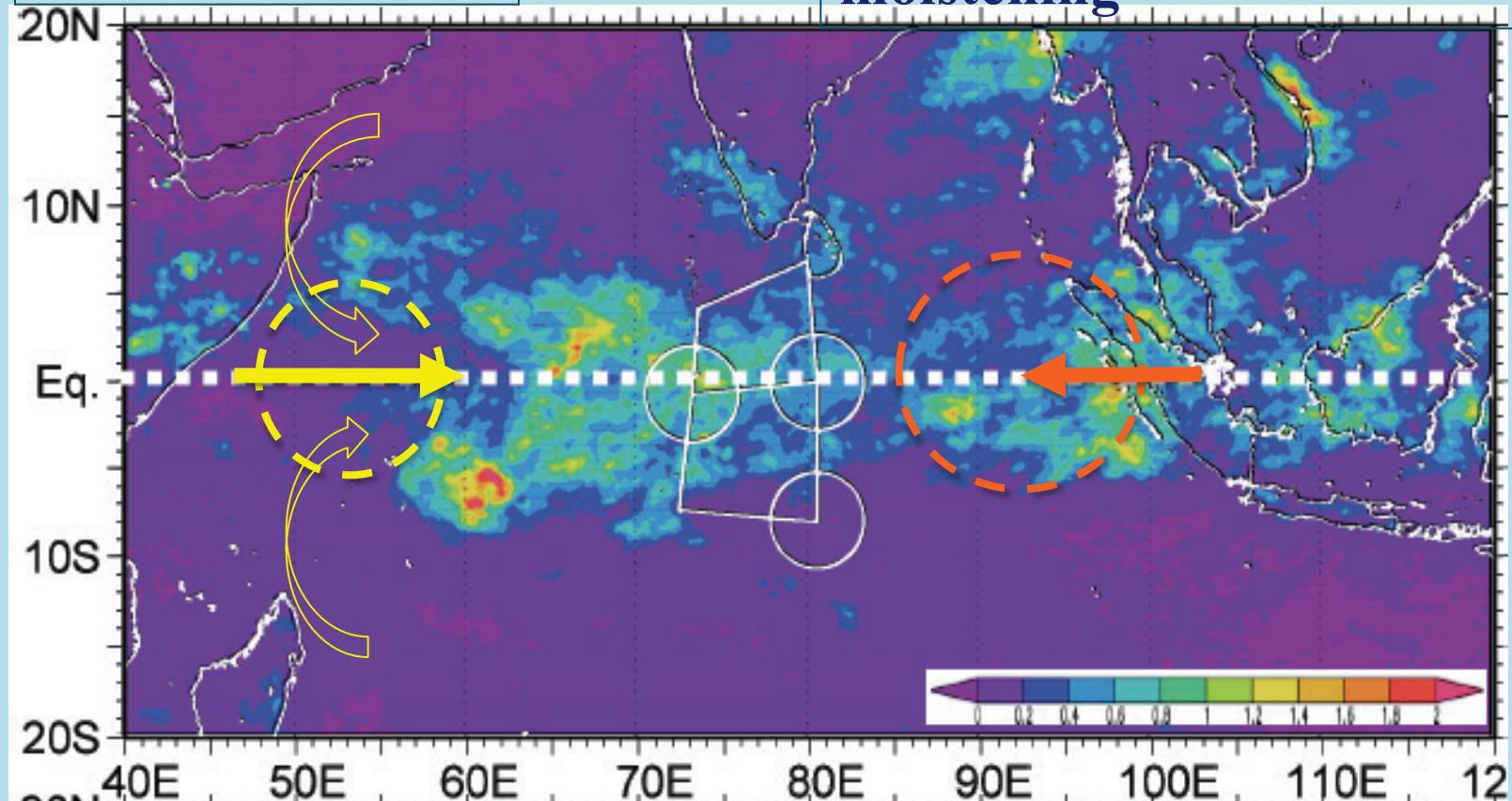
MJO-2 Nov 06 – Nov 20



MJO-2 Nov 20 – Nov 24

Dynamical drying

Dynamical vs. thermodynamical moistening



Decomposition of horizontal moisture advection

Zhao et al. (2013)

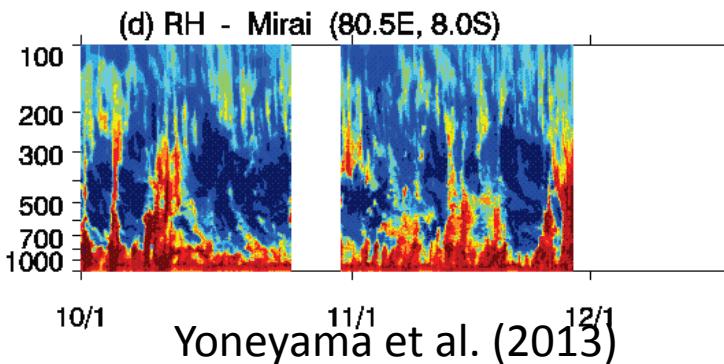
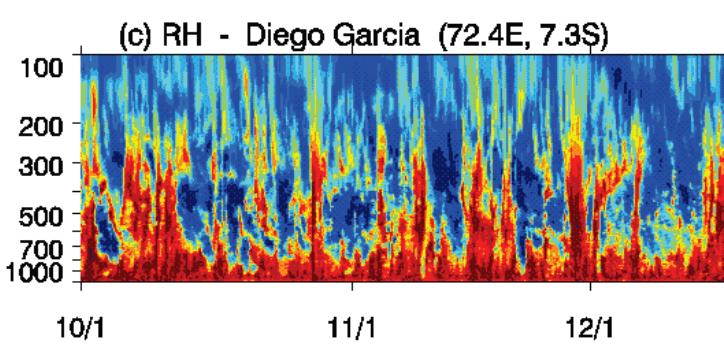
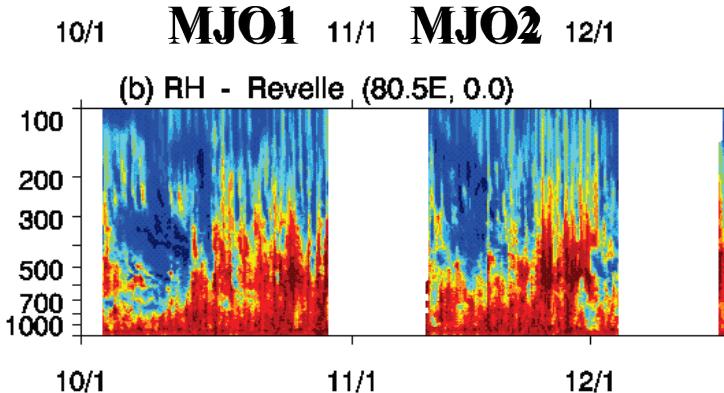
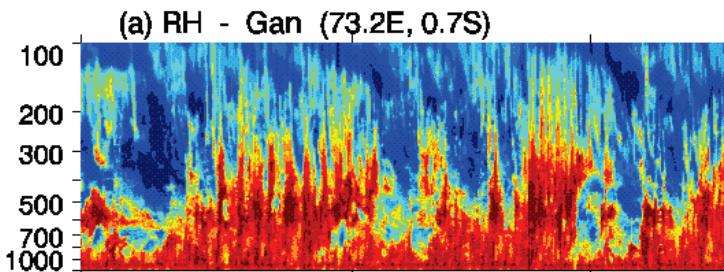
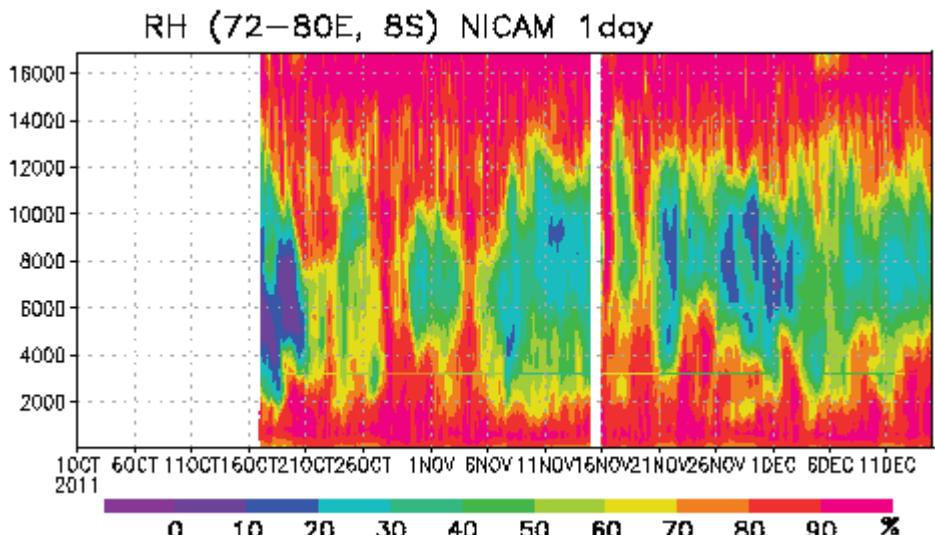
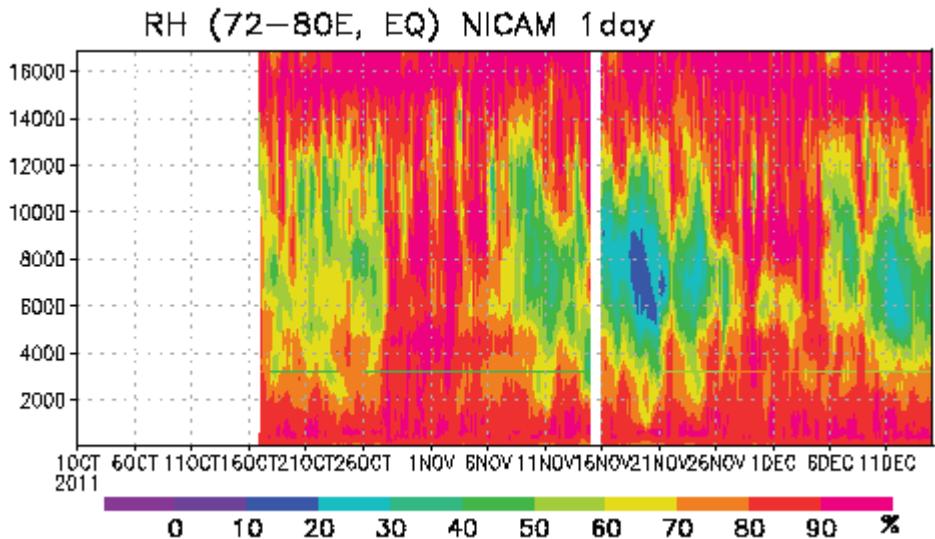
$$\left(\frac{\partial q}{\partial t} \right) = (-V \nabla \bar{q})' + -\omega \frac{\partial q}{\partial p}' + -\frac{Q_2}{L}'$$

$$u = \bar{u} + u' + u^*, \quad v = \bar{v} + v' + v^* \quad q = \bar{q} + q' + q^*$$

overbar: >80 day (**basic state**) prime: 20-80 day (ISO) asterisk: < 20 day

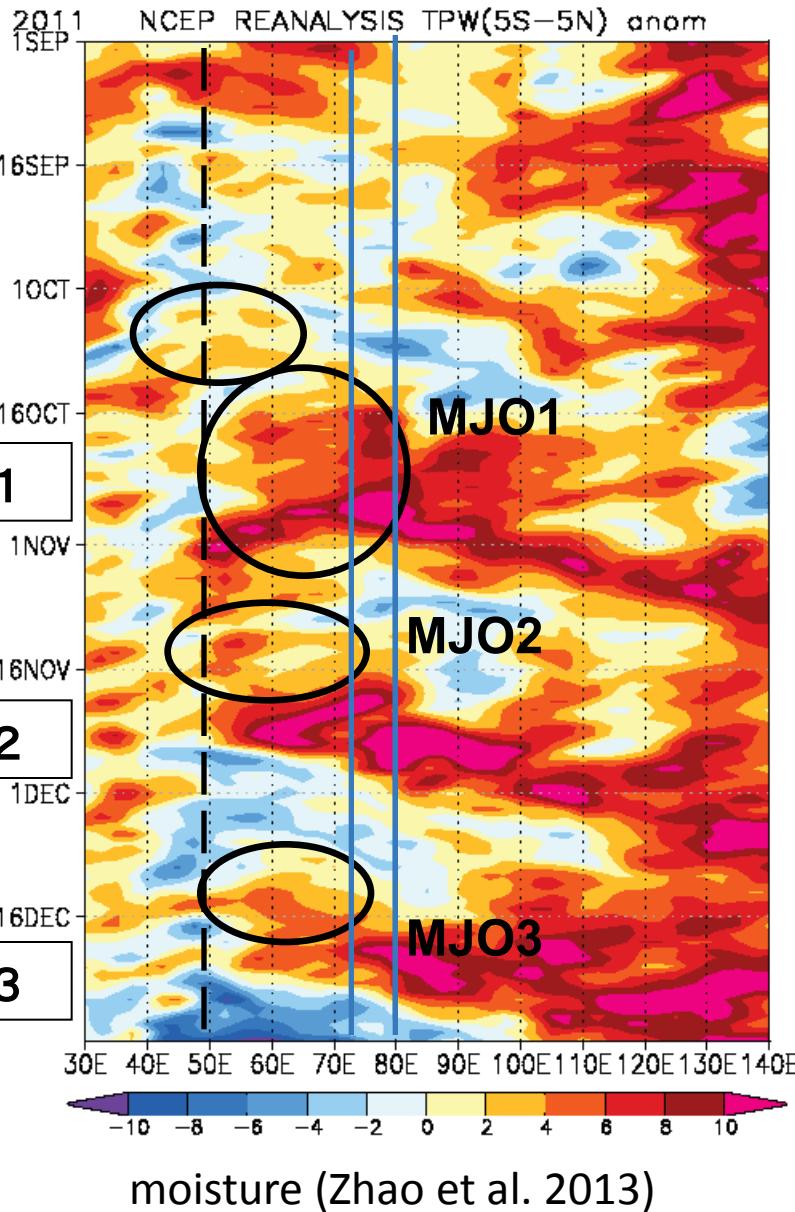
$(-\bar{V} \cdot \nabla \bar{q})'$	$(-\bar{V} \cdot \nabla q')'$ (blue circle)	$(-\bar{V} \cdot \nabla q^*)'$	$(-\bar{V}' \cdot \nabla \bar{q})'$ (red circle)	$(-V' \cdot \nabla q')'$	$(-V' \cdot \nabla q^*)'$	$(-V^* \cdot \nabla \bar{q})'$	$(-V^* \cdot \nabla q')'$	$(-V^* \cdot \nabla q^*)'$
1	2	3	4	5	6	7	8	9

Moisture buildup in CINDY2011/DYNAMO Array and 7-km mesh NICAM hindcasts

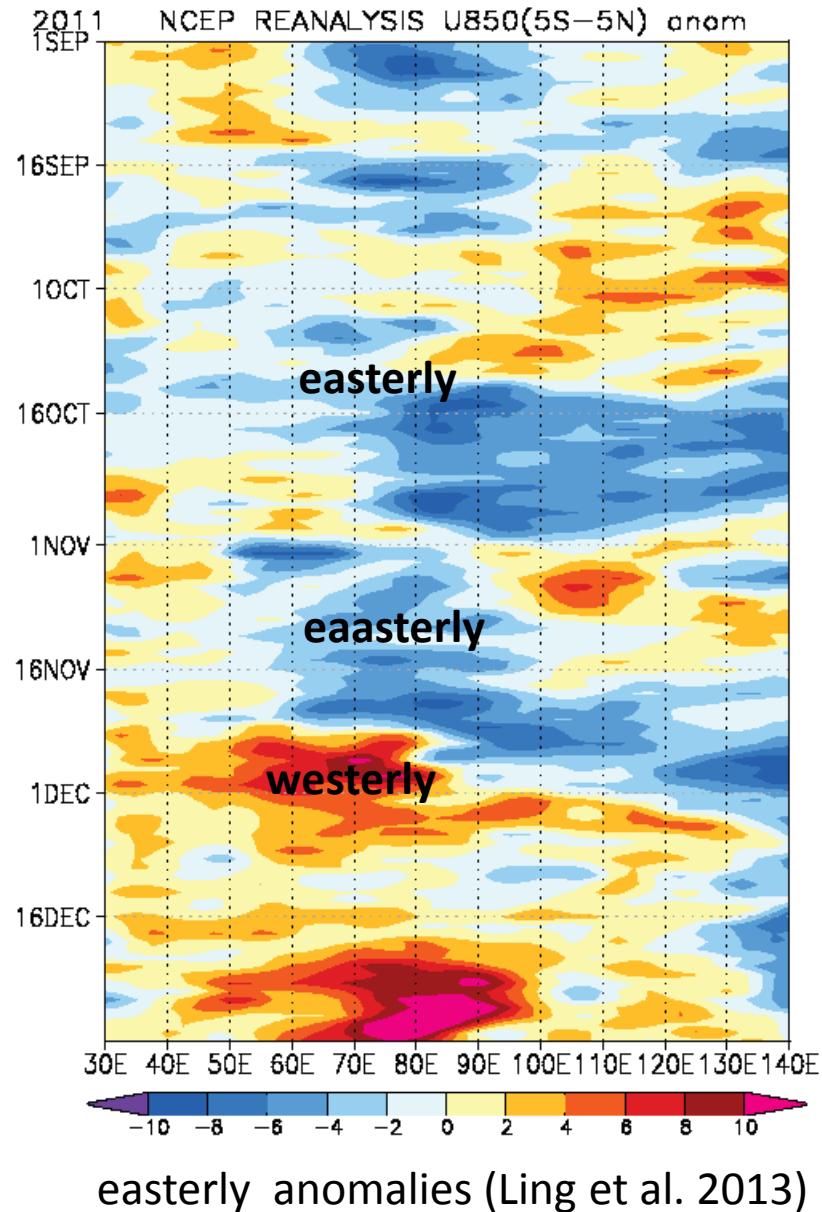


Precursor of MJO convective initiation

Precipitable water (5S-5N)

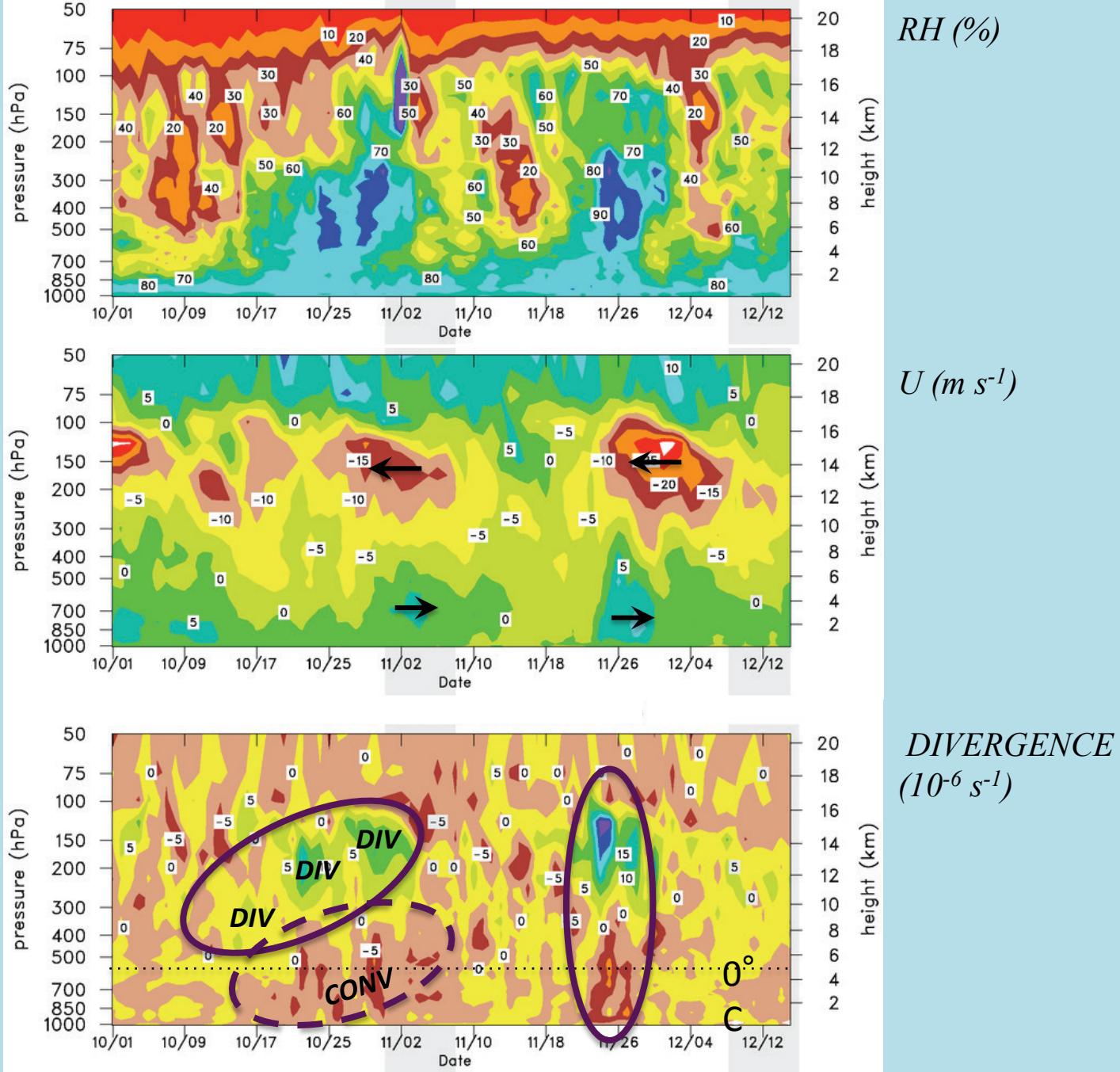


U850 (5S-5N)





Averages of the Northern Array



Courtesy of R. Johnson

Hypothesis on MJO initiation

- **Recharge-discharge (Moist Static Energy, moisture)**

Blade and Hartmann (1993); Kemball-Cook and Weare (2001); Maloney and Hartmann (1998); Maloney (2009); Hu and Randall (1994)

- ★ **moisture buildup (observational evidences)**

Kikuchi and Takayabu (2004); Kiladis et al. (2005); Yoneyama et al. (2008, 2013)

“Moisture mode” Raymond (2000); Sobel and Maloney (2013)

- **Circumnavigation (tropical dynamics)**

Knutson and Weickmann 1987; Wang and Li 2004; Matthews 2008; Straub 2013

- **Extratropical forcing (Rossby wave activity; upper troposphere)**

Lau and Peng 1987; Hsu et al. 1990; Ray et al. 2009; Ray and Li 2013; Zhao et al. 2013

Precursor of MJO initiation

- Increase in moisture and temperature at low-levels; destabilization (5-10 days before convective initiation) (Zhao et al. 2013)
 - Easterly anomalies (low-mid levels), SLP ($k=1$), and negative T (mid levels) appear 20-day before convective initiation and move eastward (Ling et al. 2013)
- ★ Importance of “local “processes to convection initiation (Straub 2013)

Motivation

Convection initiation of MJO over the western **Indian Ocean (IO)** is one of most difficult tasks in understanding and prediction of MJO. The **CINDY2011/DYNAMO** field program was conducted to tackle this tough problem through October 2011-March 2012 (Yoneyama et al. 2013). Existing theories on MJO dynamics addressed the importance of **buildup of moisture/moist static energy (MSE)** [e.g., Blade and Hartmann 1993; Kemball-Cook and Weare 2001; Maloney 2009] . Zhao et al. (2013) demonstrated that accumulation of moisture in the lower troposphere (-5~10 day) as a precursor of MJO convection initiation by composite analysis of 20 year boreal winter cases. Here, we discuss the moistening processes in two MJO cases that initialized during the CINDY/DYNAMO period. Our major concern is the roles of dynamical precursors (e.g., extratropical forcing, circumnavigating signals) and high frequency disturbances on the initiation of convection that broke out at a local scale (Straub 2013) in the western IO region.