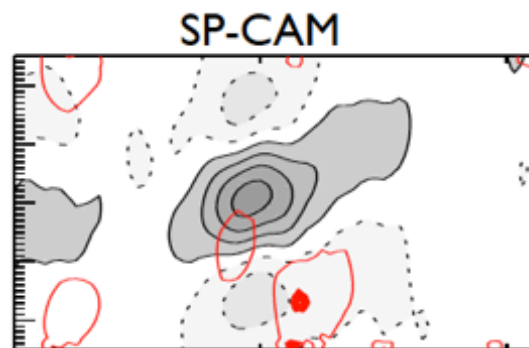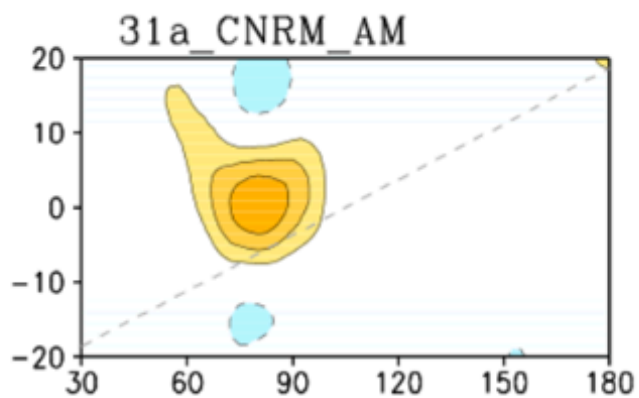# Why does ocean coupling improve the MJO?

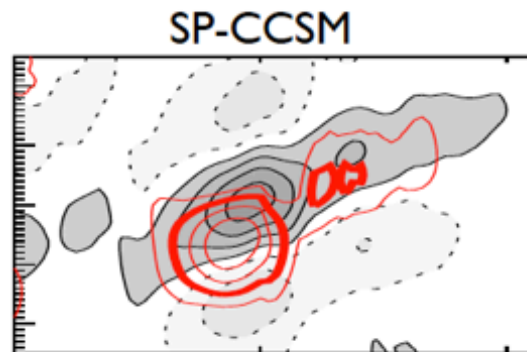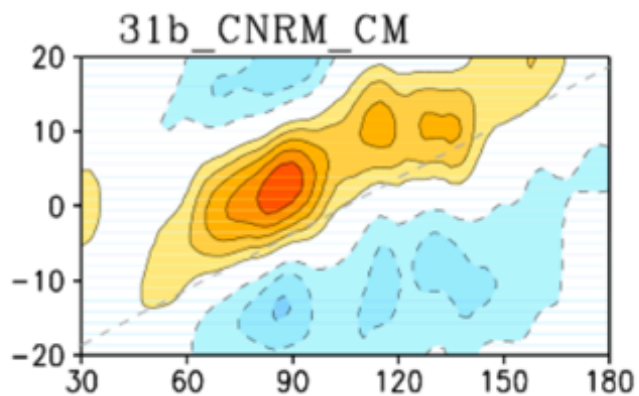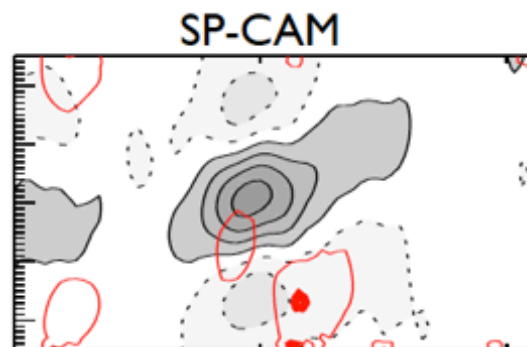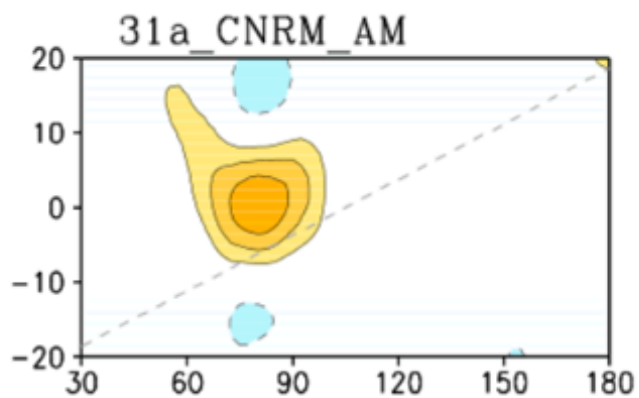Charlotte A. DeMott[1], Cristiana Stan[2], David A. Randall[1], and Mark D. Branson[1]

[1]Department of Atmospheric Science, Colorado State University
[2]Department of Atmospheric, Oceanic and Earth Sciences, George Mason University
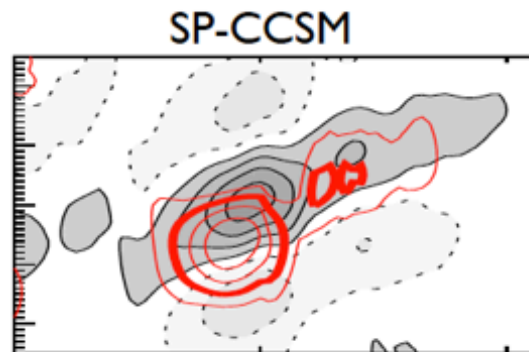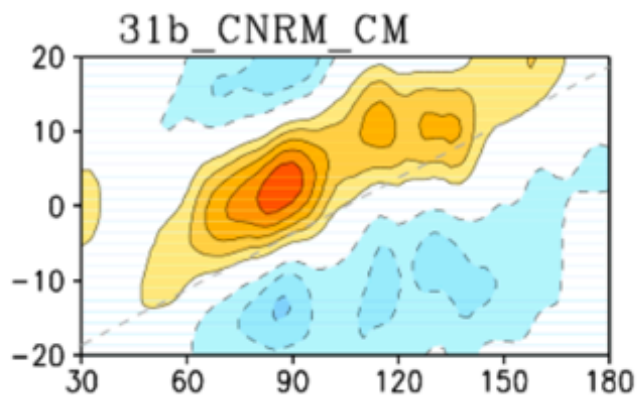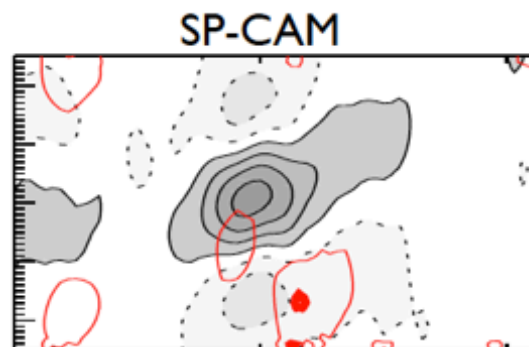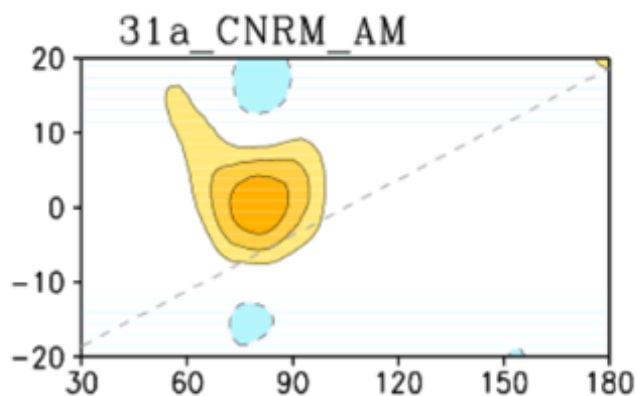
how can we understand the role of air-sea interaction on the simulation of the ISO?

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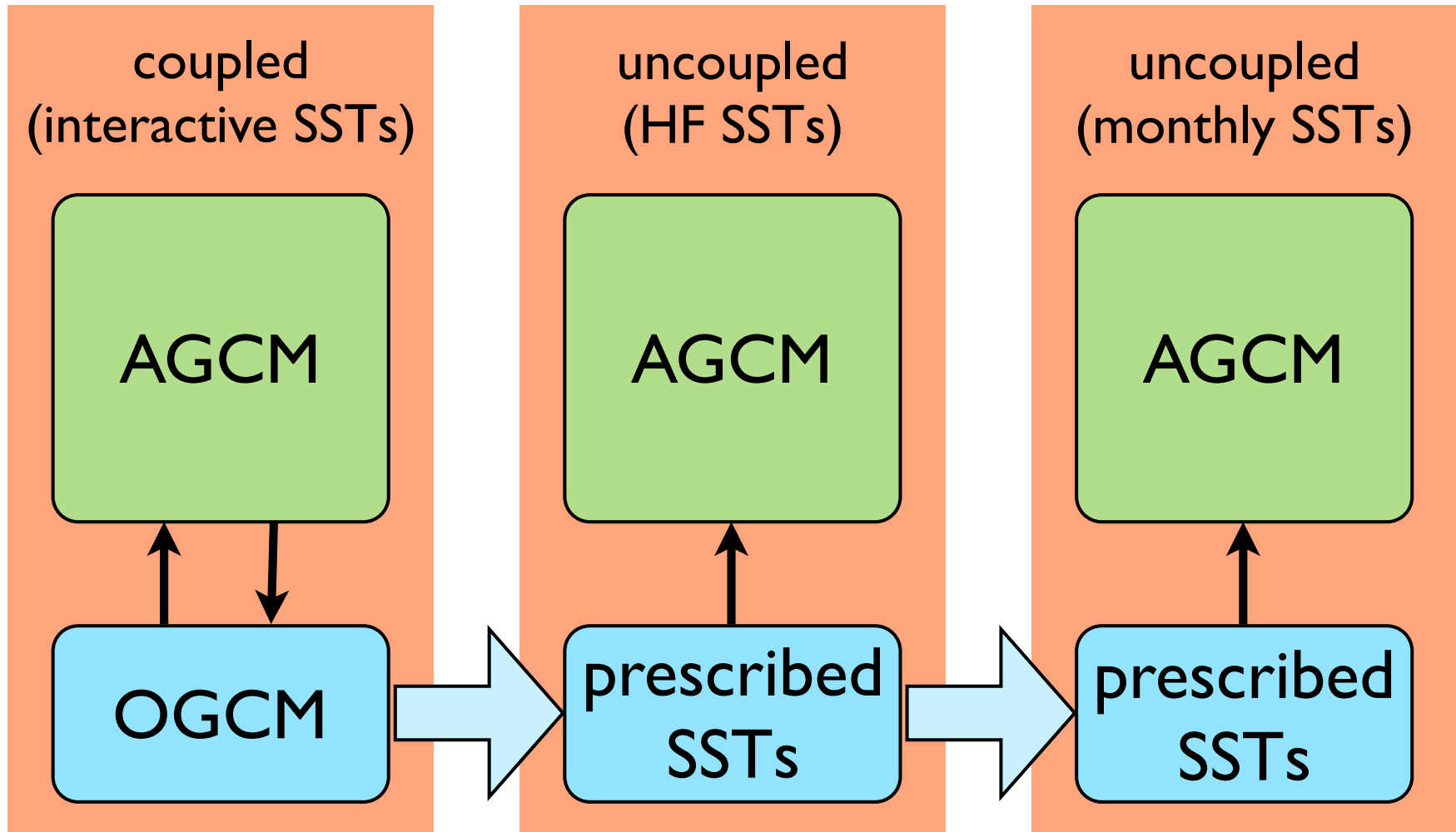
how can we understand the role of air-sea interaction on the simulation of the ISO?
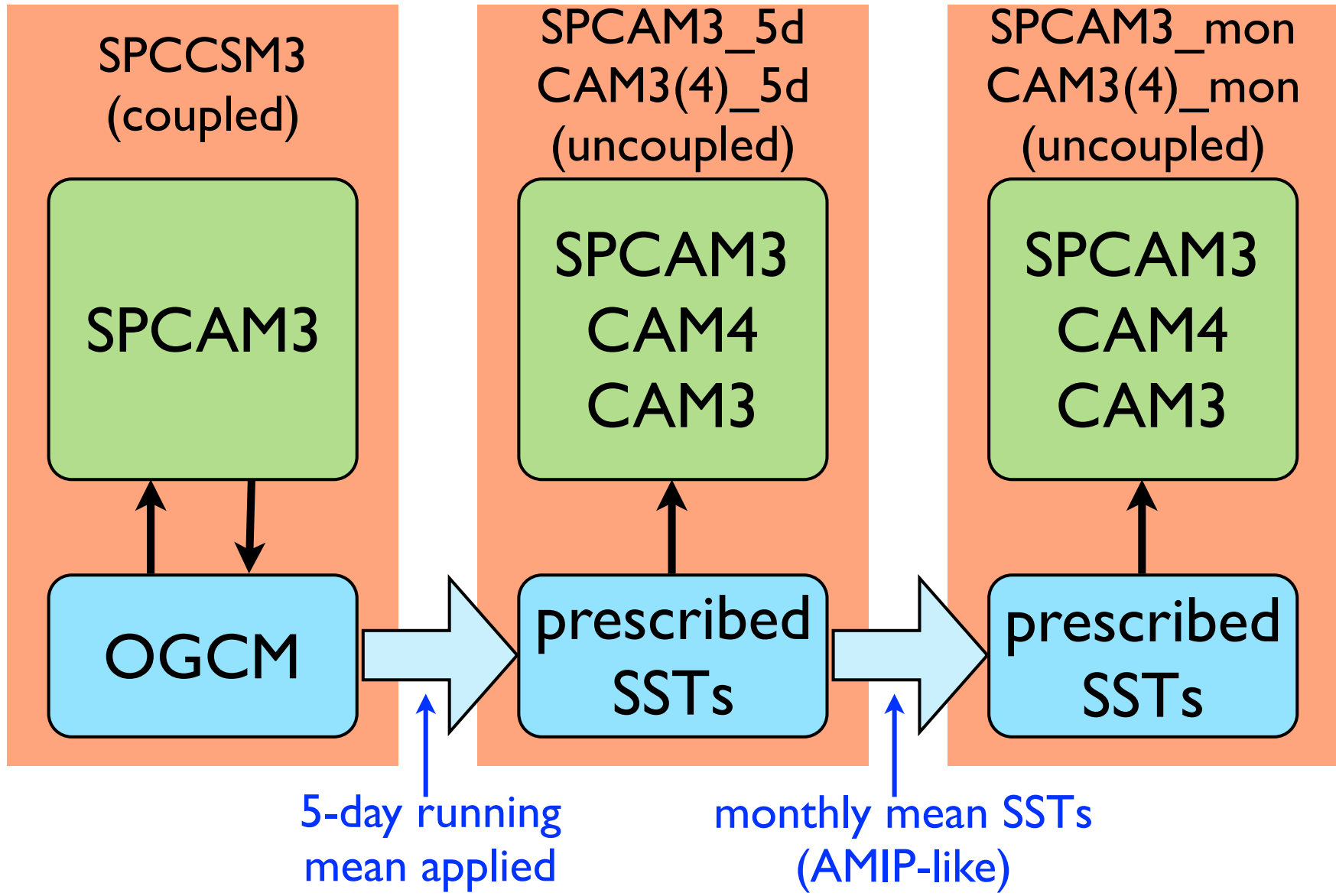

CNRM analysis
by Xianan Jiang
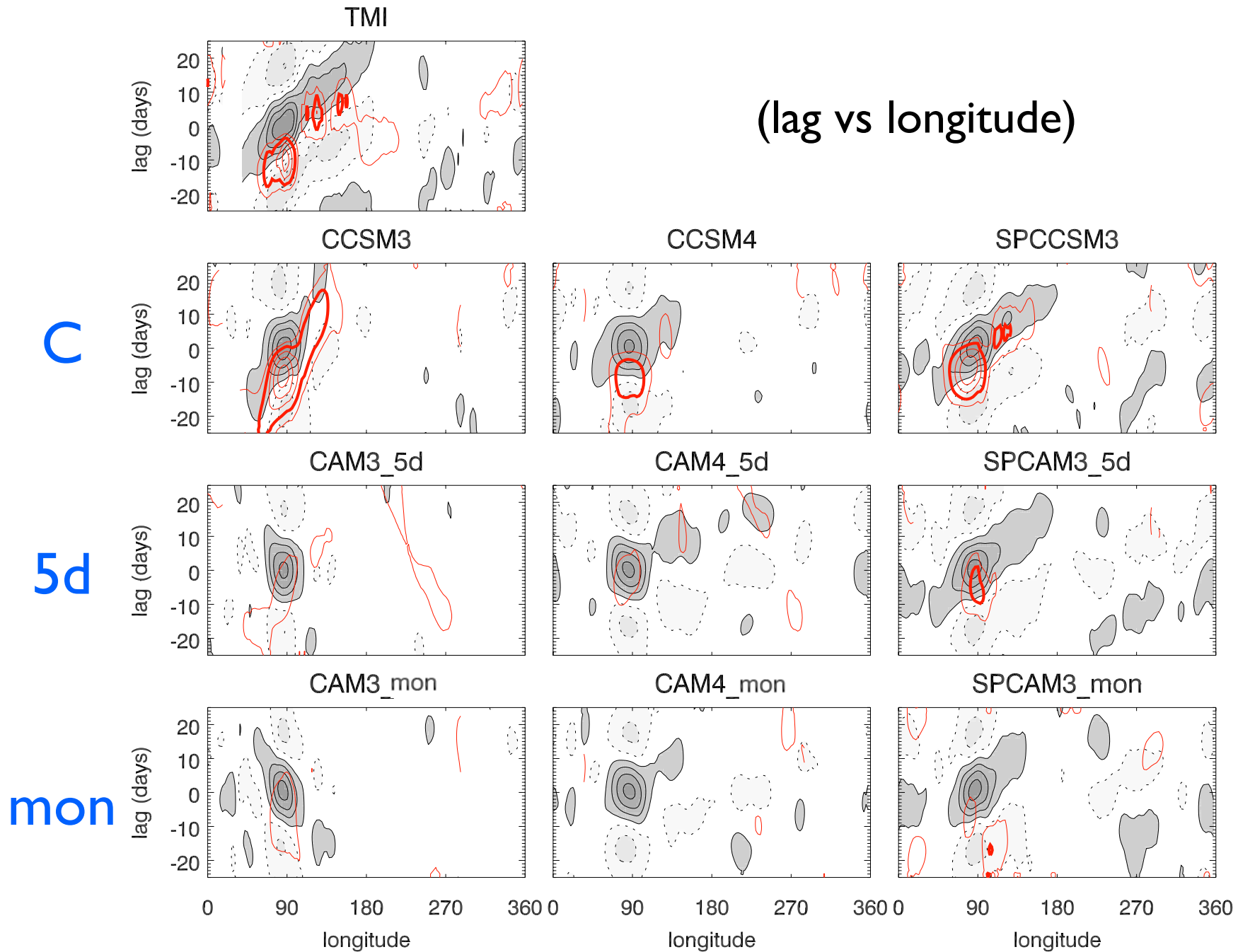
model physics differences cloud the issue
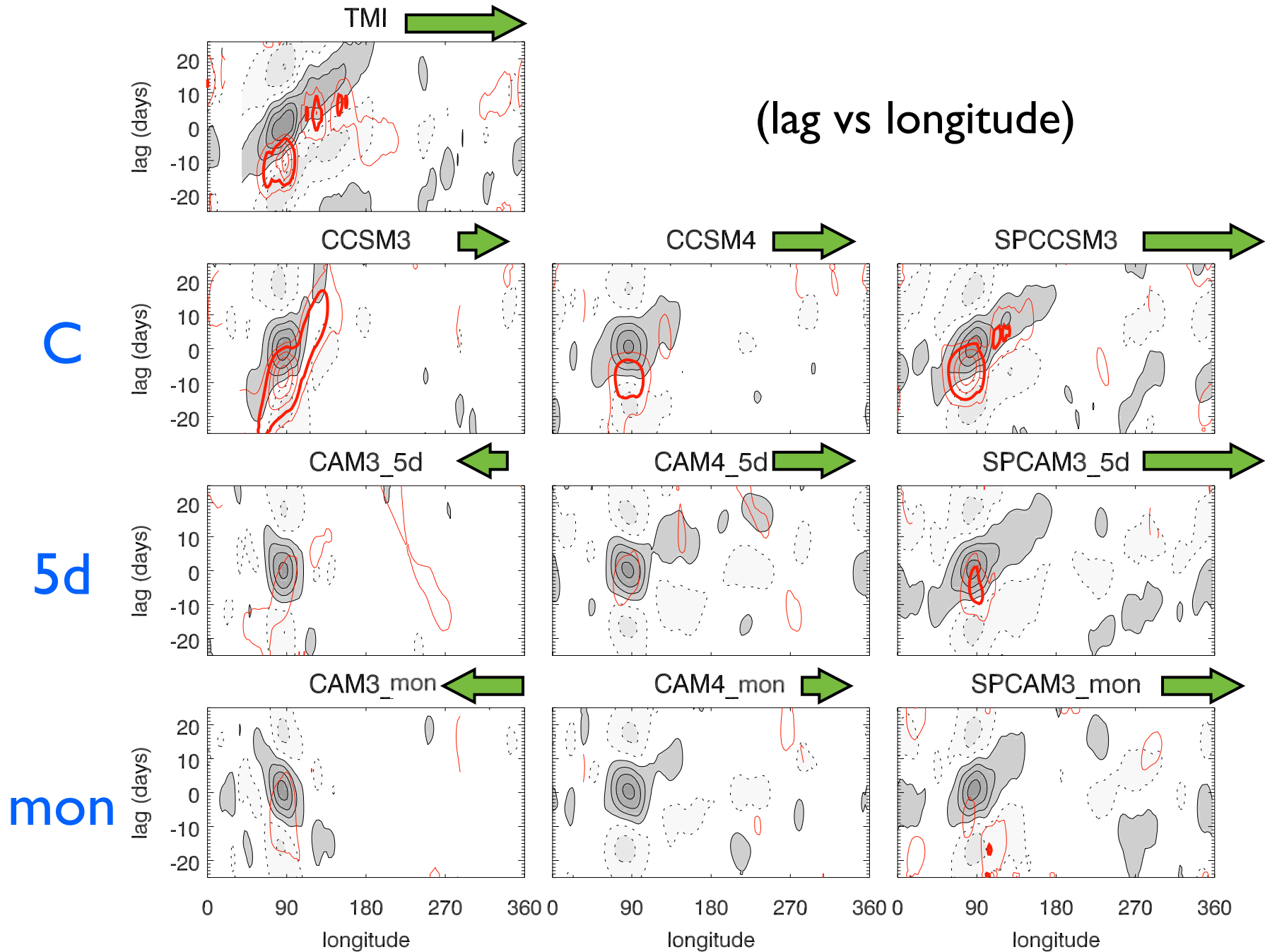
experimental design

experimental design

rainfall and SST anomaly lag-correlation (all seasons)

(lag vs longitude)

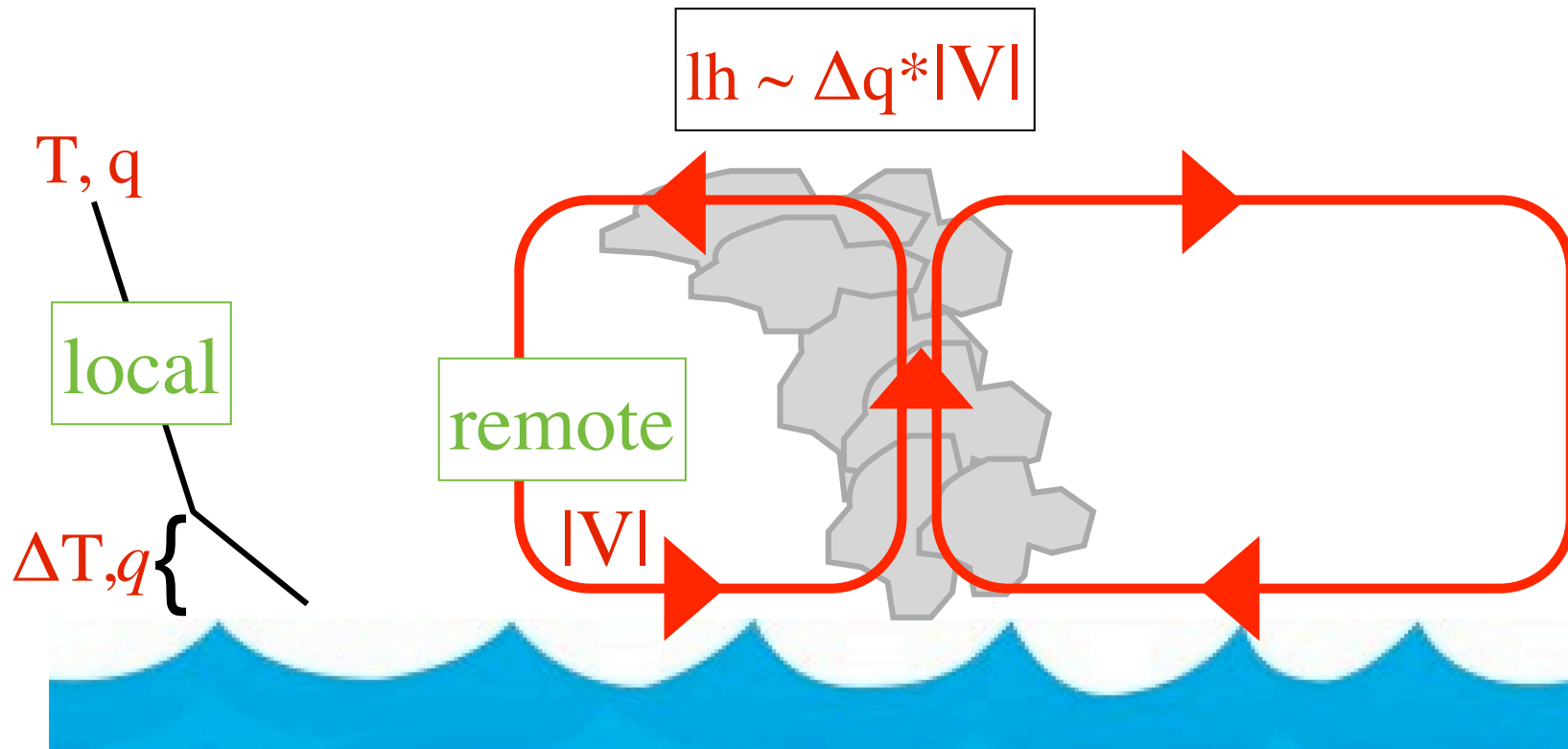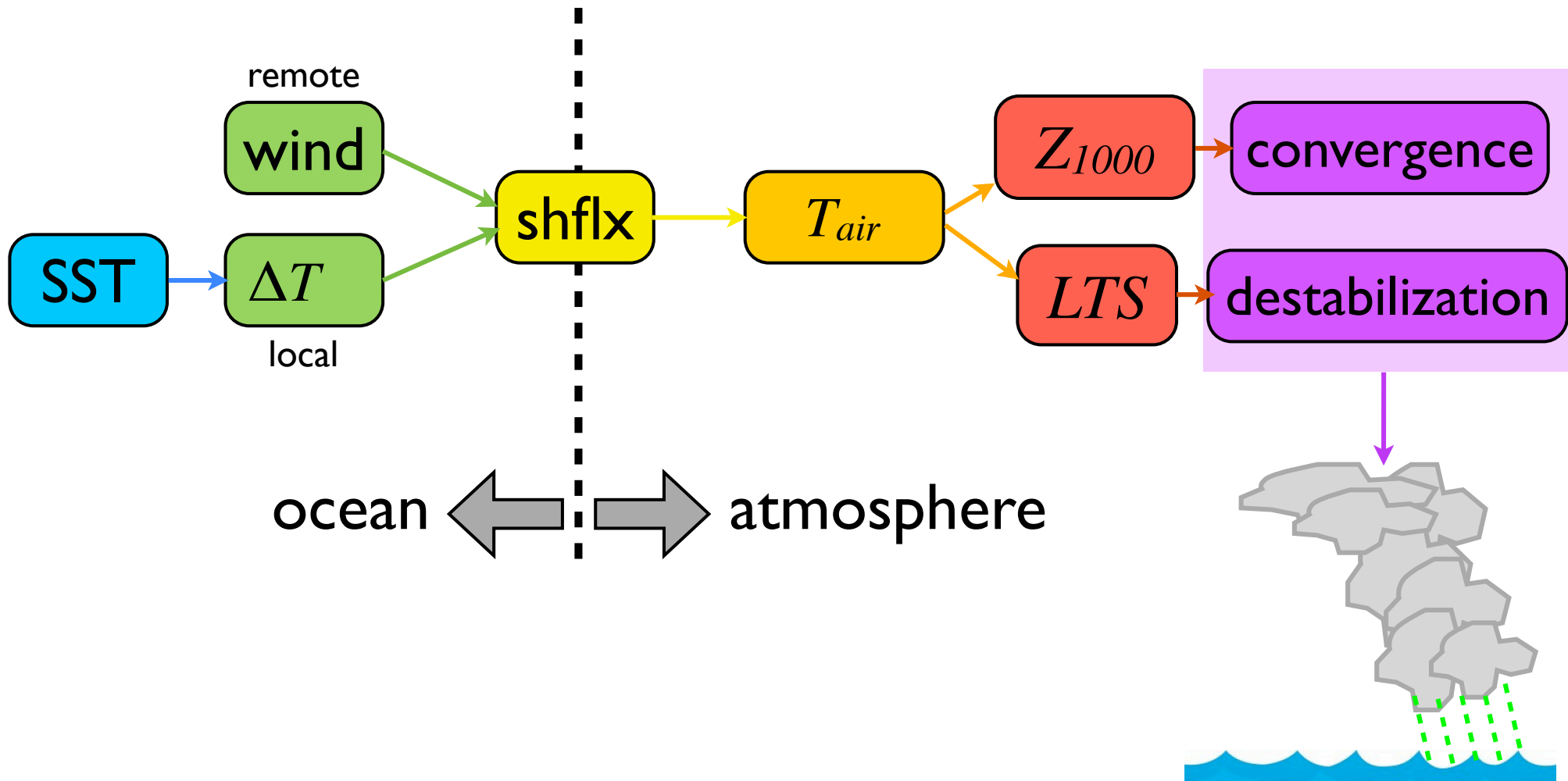# rainfall and SST anomaly lag-correlation (all seasons)

(lag vs longitude)

the atmosphere senses SST anomalies
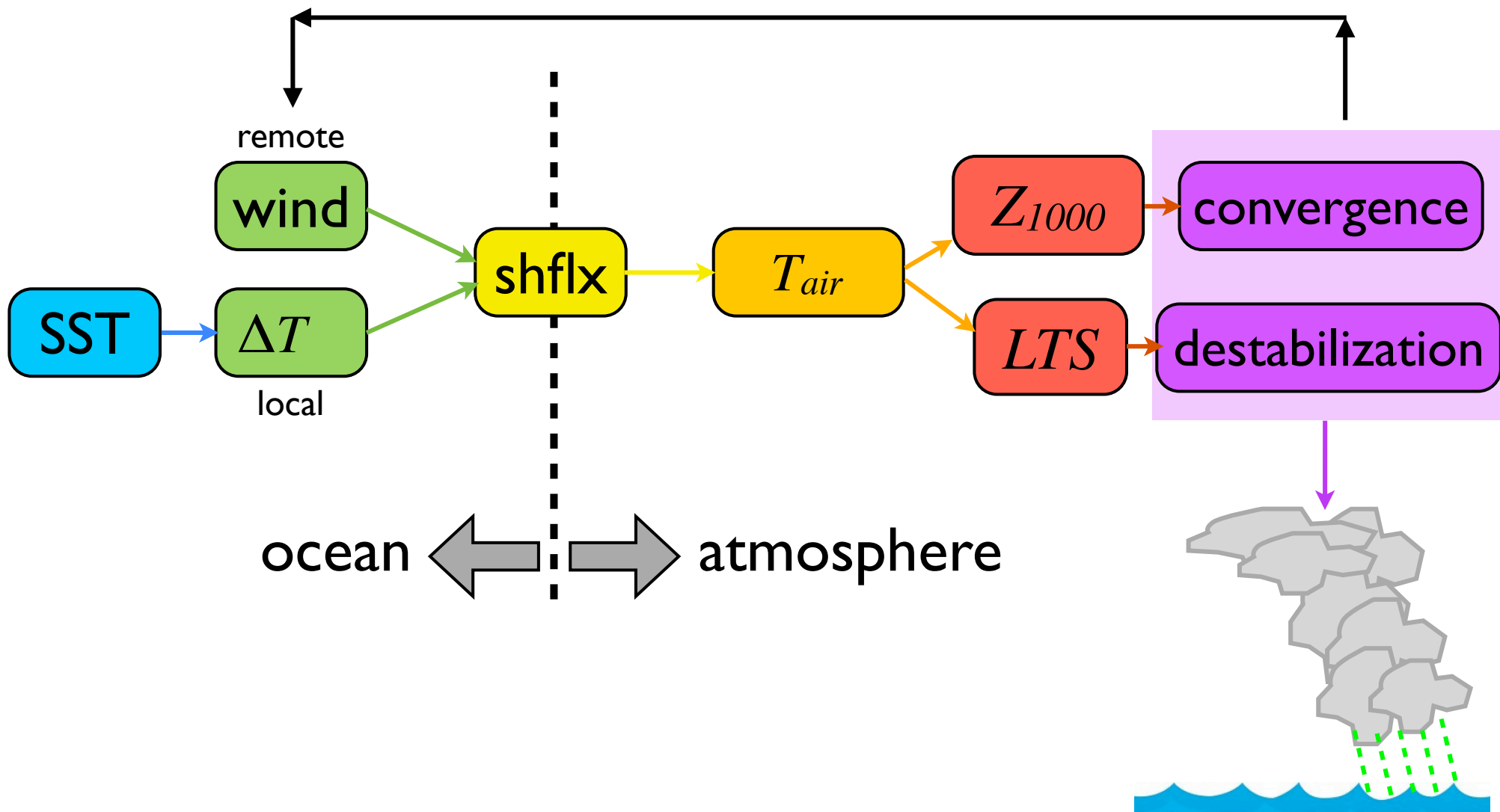via surface fluxes.

$$lh \sim \Delta q * |V|$$
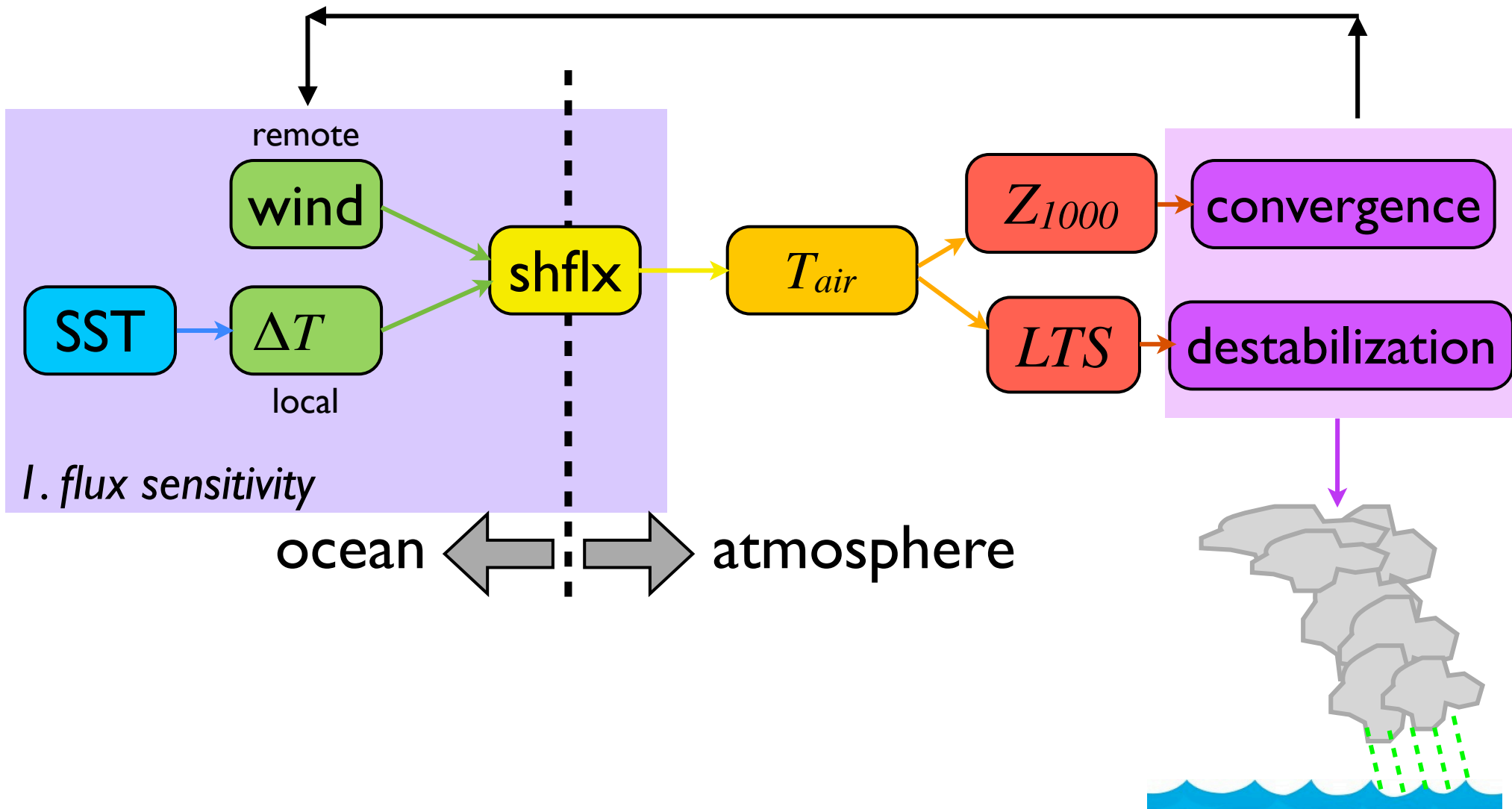
the atmosphere senses SST anomalies
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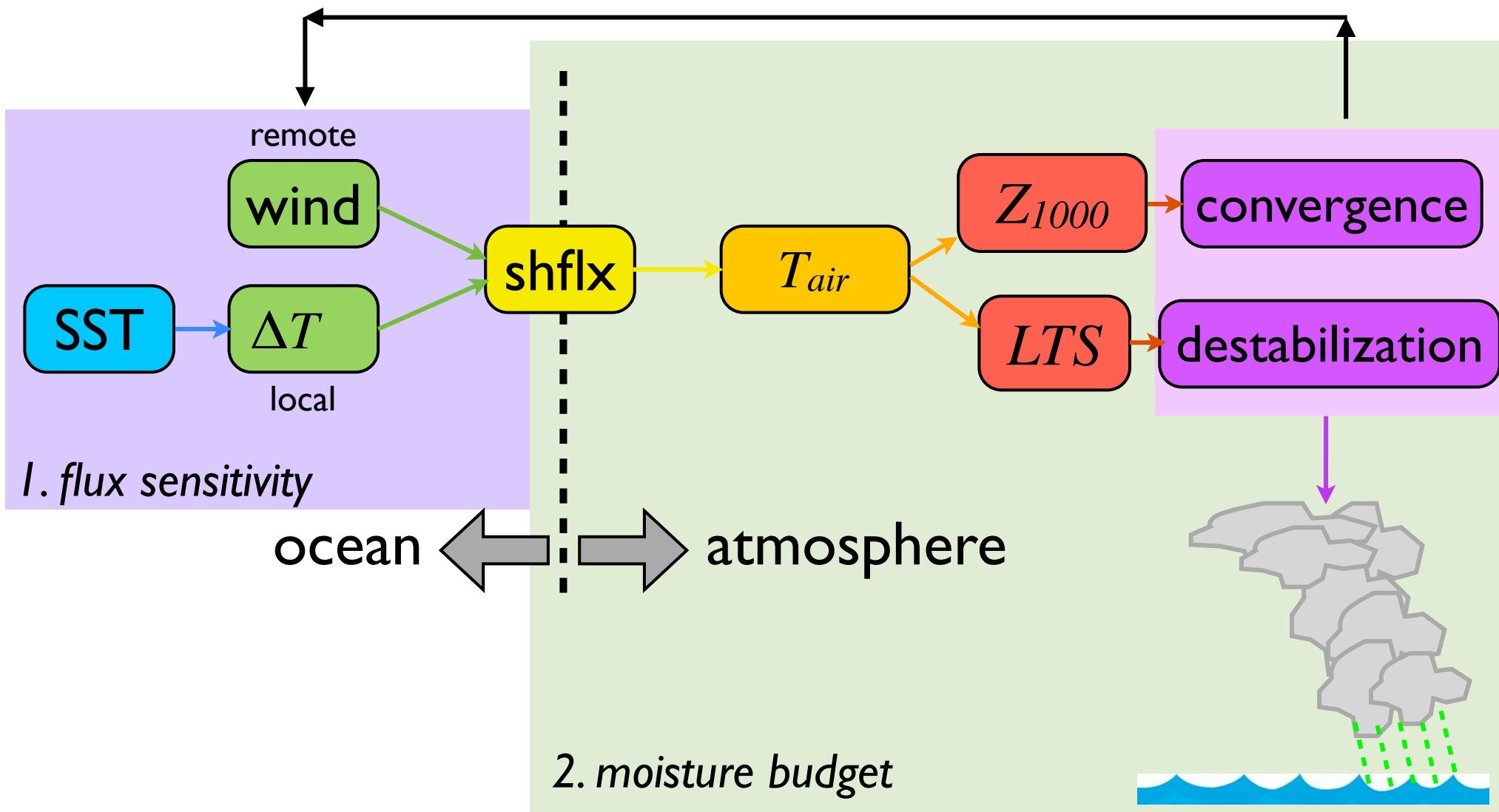$$lh \sim \Delta q * |V|$$

T, q

local

remote

ΔT, q

$|V|$

tropical air-sea interaction
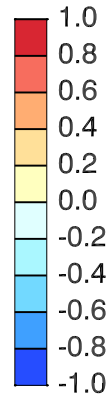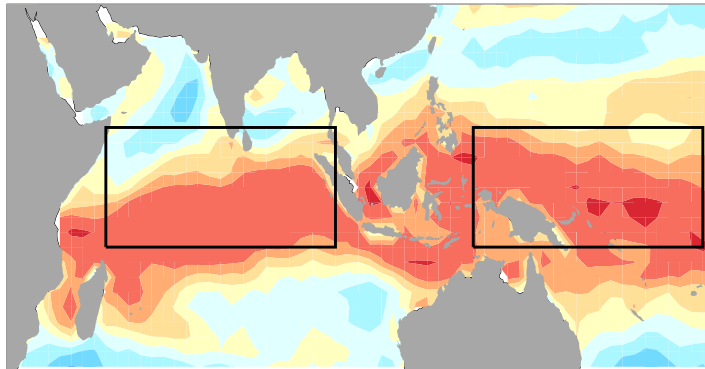
tropical air-sea interaction

tropical air-sea interaction

1. surface flux sensitivity

local vs remote control of surface fluxes

ERAI DJF

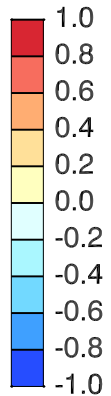*LH* ratio difference: |V|' - (Δq')

wind dominates

$\Delta T, q$ **dominates**

SH ratio difference: |V|' - (ΔT')
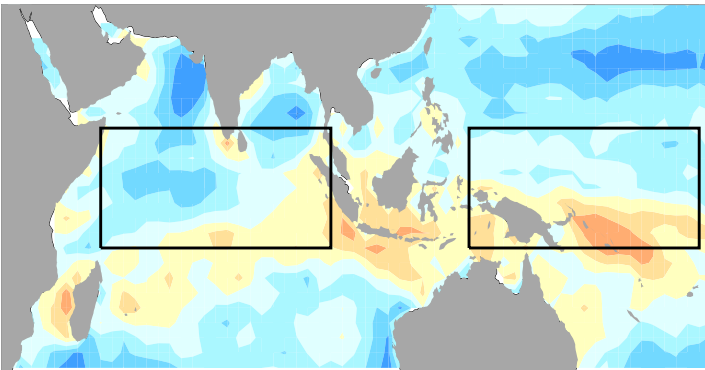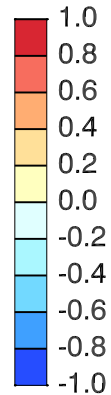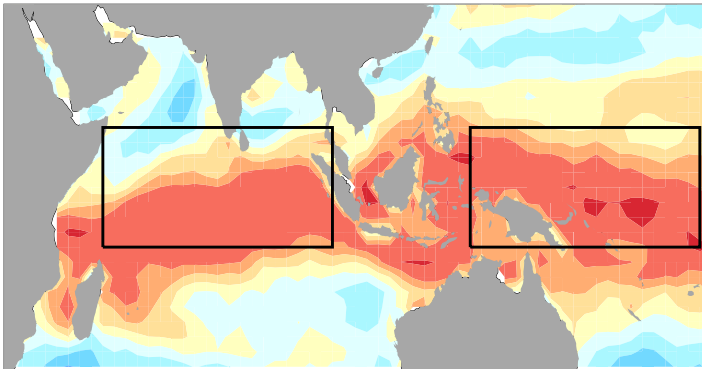
local vs remote control of surface fluxes

ERAI DJF

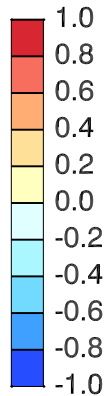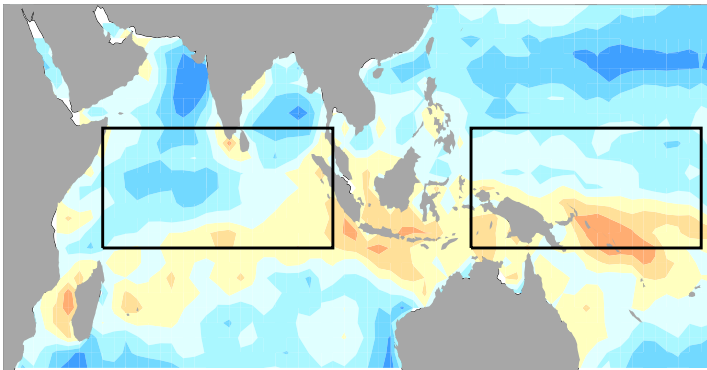**LH** ratio difference: |V|' - (Δq')

} **wind dominates**

} $\Delta T, q$ **dominates**
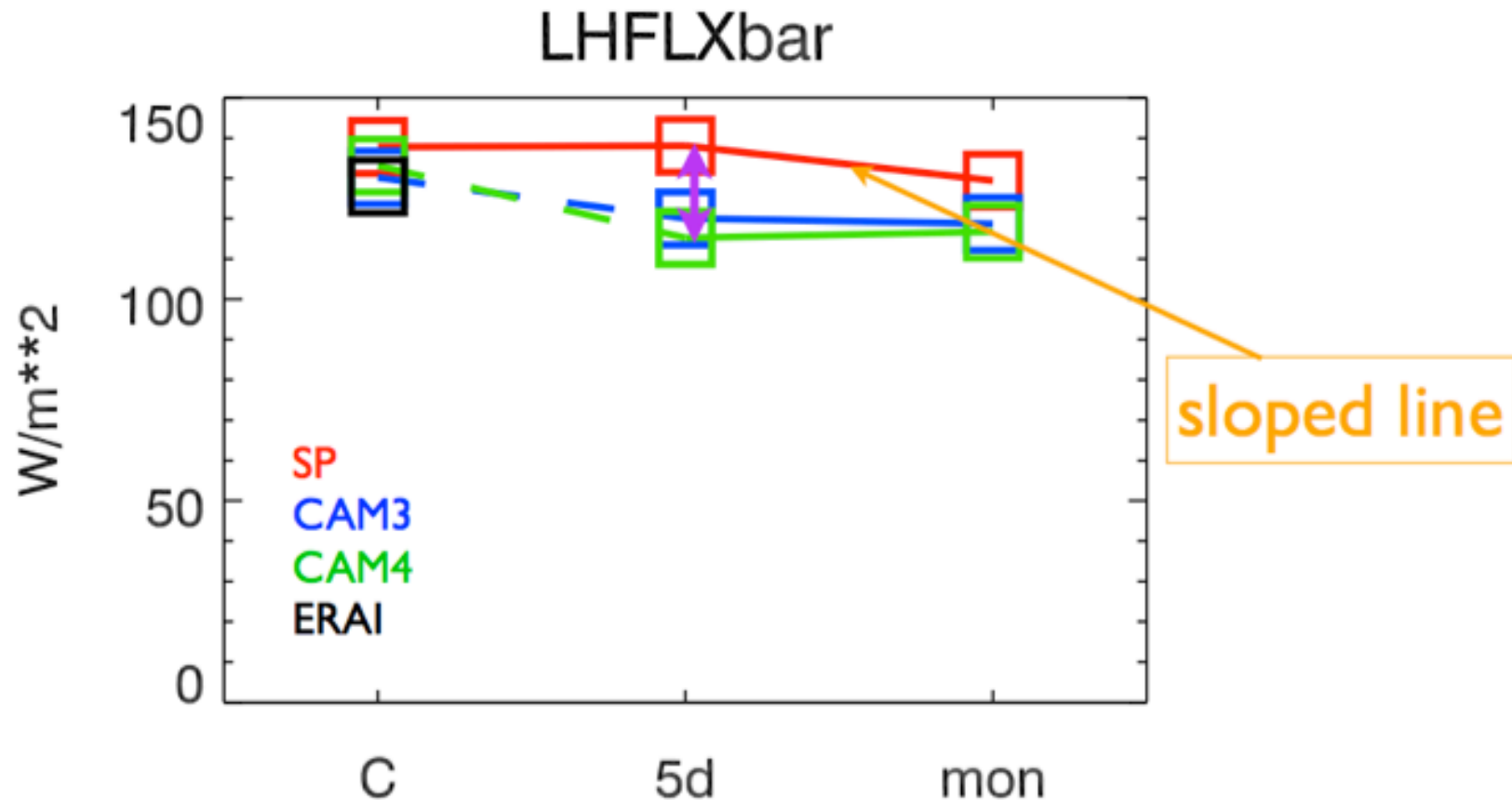
SH ratio difference: |V|' - (ΔT')

- how sensitive are local and remote effects to SST?

- for observations and models?

we will focus on the average over the **Indian Ocean**

"box" diagrams
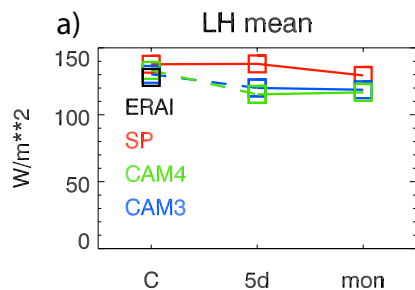

vertical offsets: arise from model physics
slopes of solid lines: arise from ocean treatment
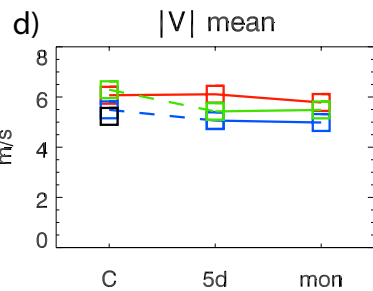slopes of dashed lines: may arise from both
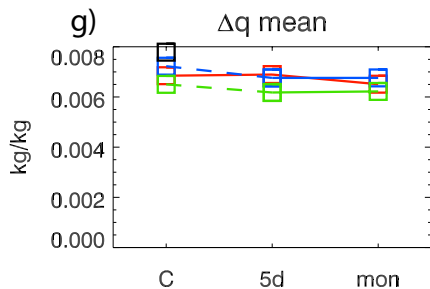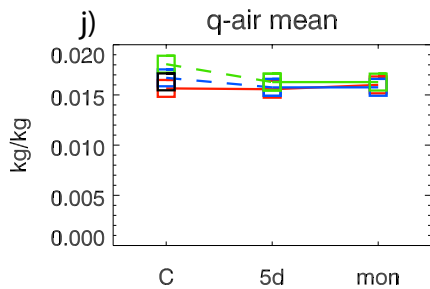
Indian Ocean latent heat flux relationships

Indian Ocean latent heat flux relationships

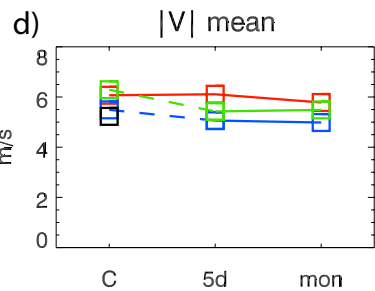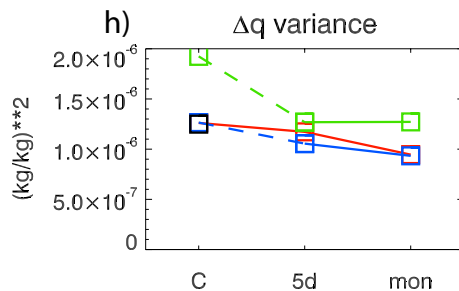# Indian Ocean latent heat flux relationships

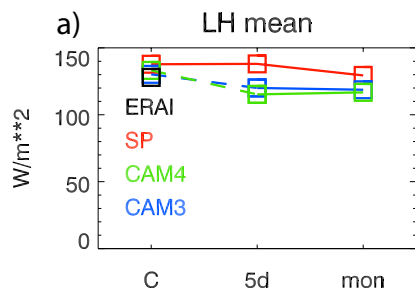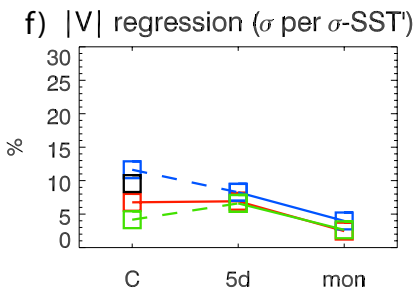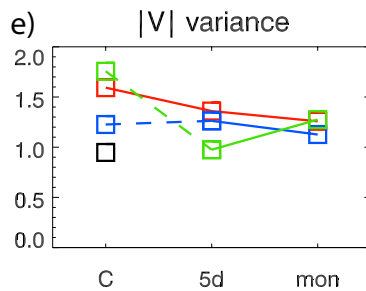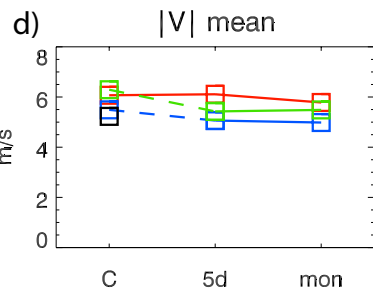# Indian Ocean latent heat flux relationships

- "local" or "direct" SST control accounts for <30% of local dynamic range (<10% of total variance).

- latent heat flux variance is primarily driven by wind speed variance, which is a "remote" or "indirect" effect of coupling.

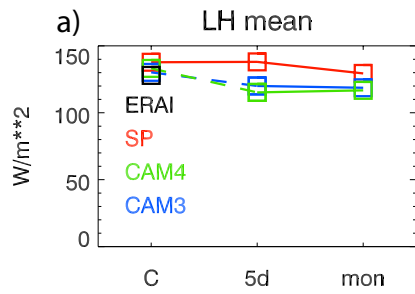2. moisture budget

vertically integrated moisture budgets

$$dq/dt = (\text{vert. adv.} - \text{precip.}) + \text{horiz. adv.} + \text{lhflx} + \text{res.}$$

moisture budget cross section composites
all contours [1x10^6 g/kg/s] / [mm/day]

mean q[g/kg] (departure from zonal mean)

moisture budget cross section composites
all contours [$1{\times}10^6$ g/kg/s] / [mm/day]

mean q[g/kg] (departure from zonal mean)

interpretation

interpretation

- mid-level moistening determines propagation direction.
- horizontal advection by convectively-driven wind anomalies dominates moisture advection.
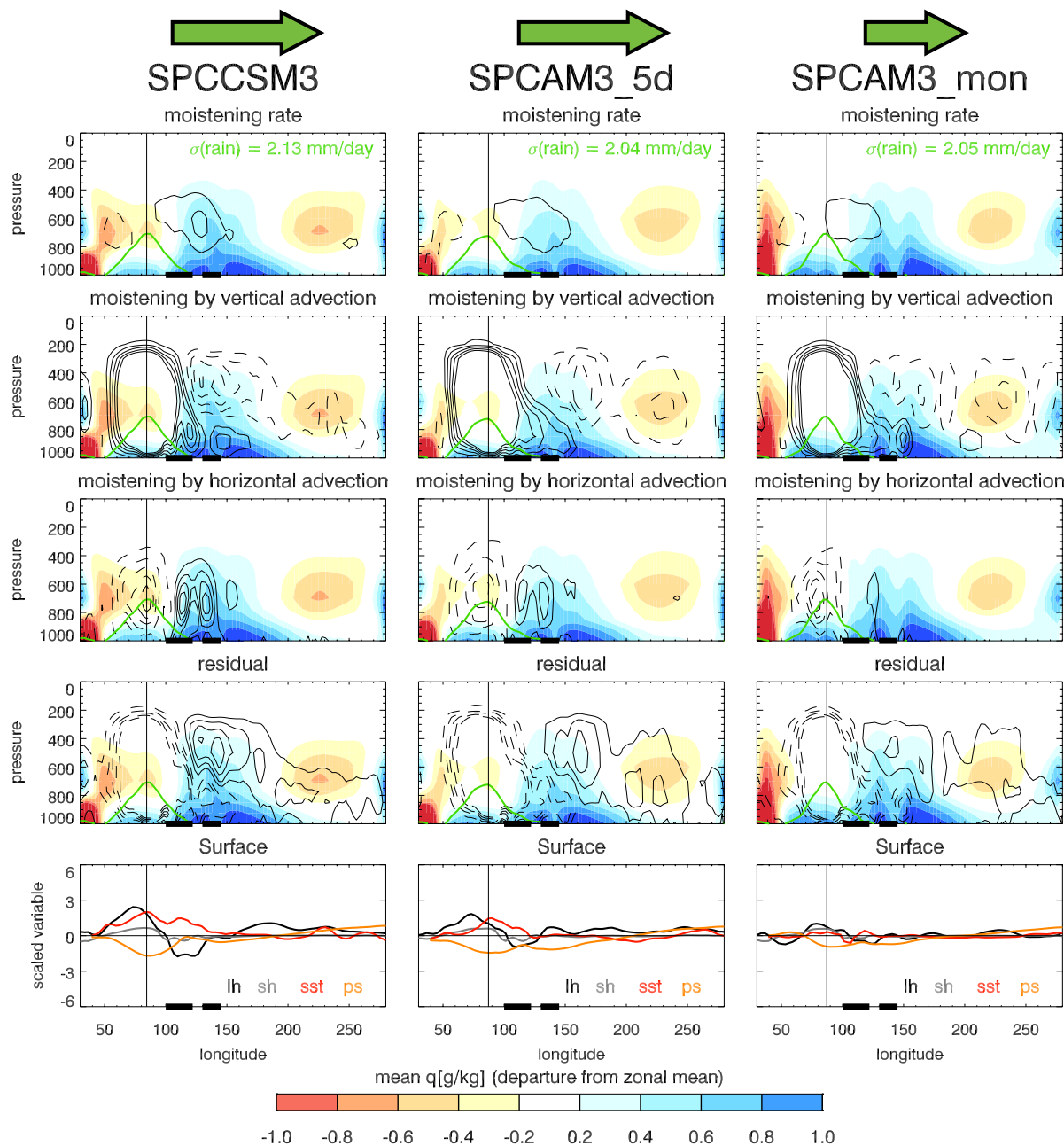- local air-sea interaction during IO developing phase produces more robust convection capable of downstream moistening.
- model physics strongly influences the "processing" of surface fluxes.

conclusions

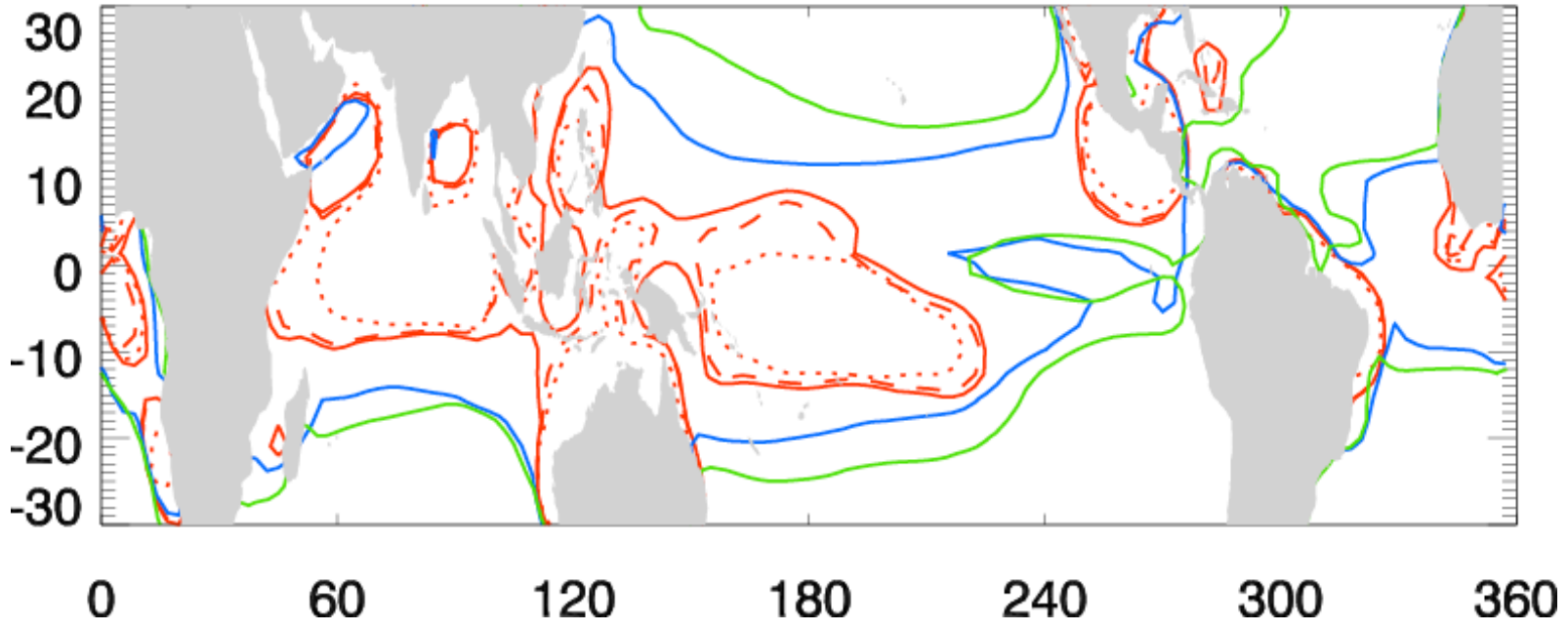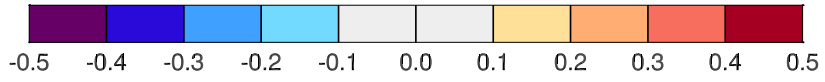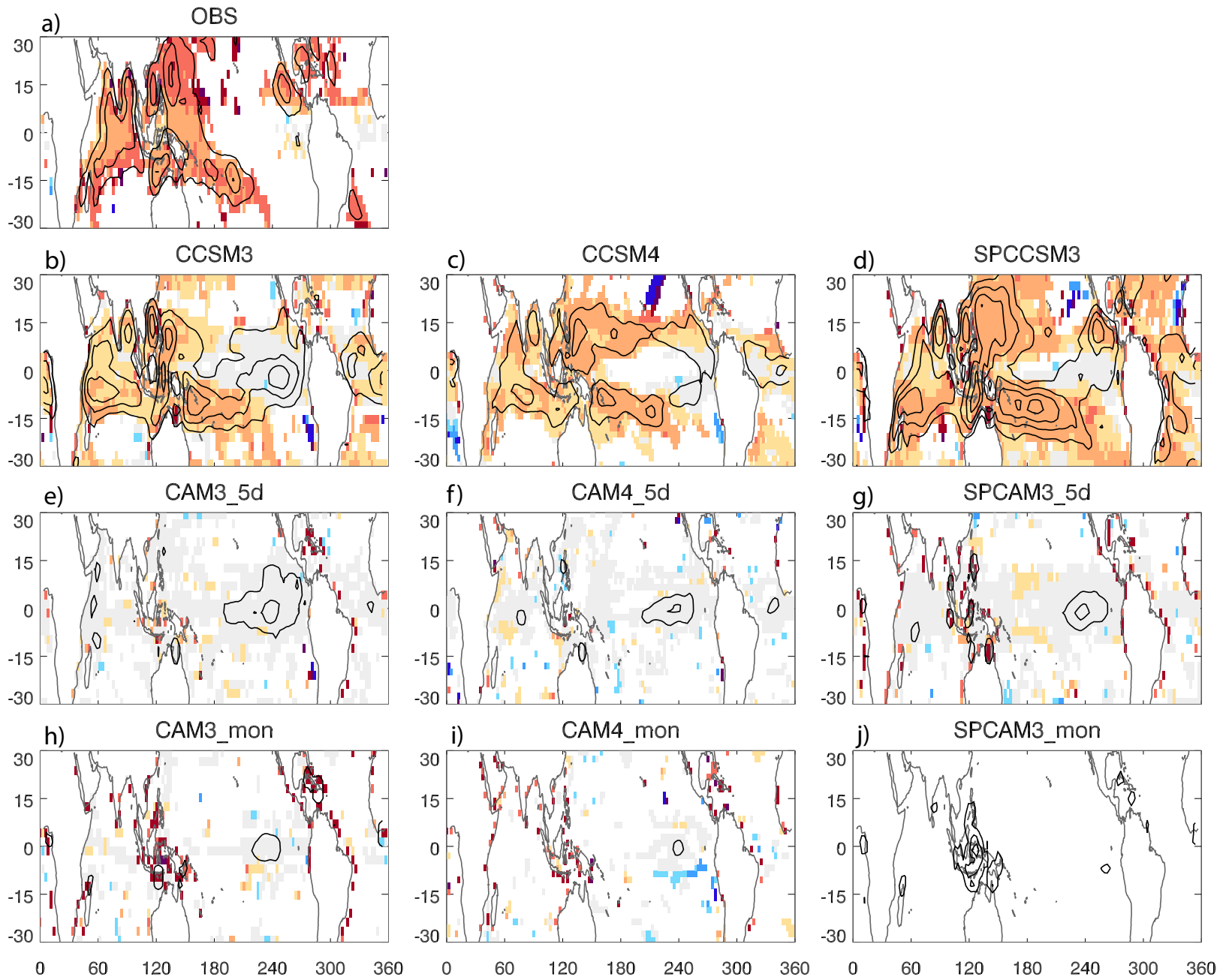1. **local** effects in Indian Ocean can "intensify" ISO convection, enhancing **remote** effects over MC.

2. air-sea coupling can lead to more robust convection, which drives stronger downstream advective moistening.

3. details of model physics may strongly influence points 2 and 3.

extra slides

P(SST≥29°C) ≥ 5%

ERAI DJF

a) normalized LHFLX'-Δq' regression [σ/σ]

b) normalized LHFLX'-|V|' regression [σ/σ]

c) (Δq)'-predicted LHFLX variance [(W/m²)²]

d) |V|'-predicted LHFLX variance [(W/m²)²]

e) variance ratio: (Δq)'-predicted / total

f) variance ratio: |V|'-predicted / total

g) ratio difference: |V|' - (Δq')

ERAI DJF

a) normalized SHFLX'-ΔT' regression [σ/σ]

b) normalized SHFLX'-|V|' regression [σ/σ]

c)(ΔT)'-predicted SHFLX variance [(W/m²)²]

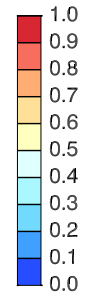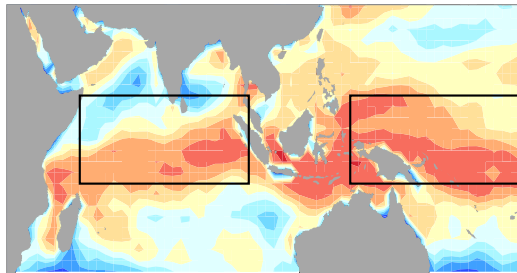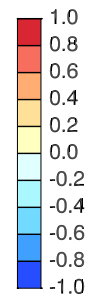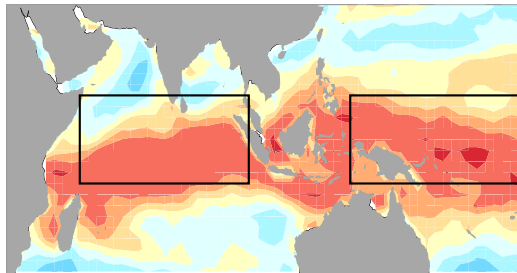d)|V|'-predicted SHFLX variance [(W/m²)²]

e) variance ratio: (ΔT)'-predicted / total

f) variance ratio: |V|'-predicted / total

g) ratio difference: |V|' - (ΔT')
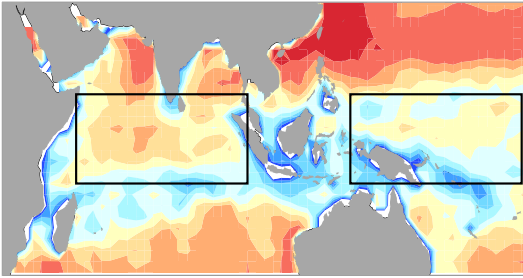
surface/lowest model level composites
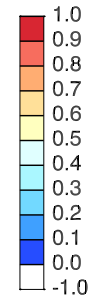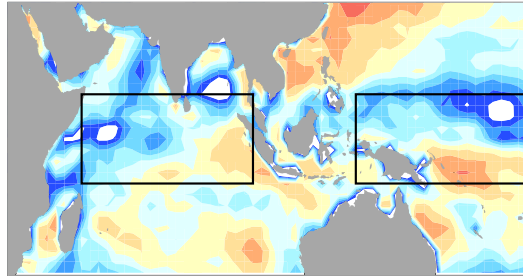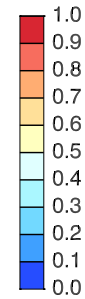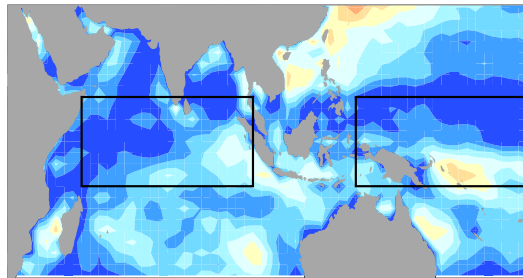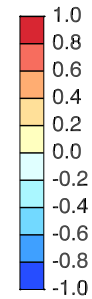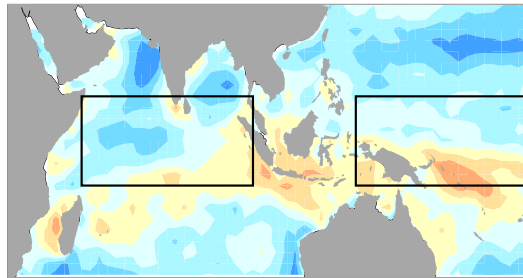all units [*x*] / [mm/day]

latent heat flux anomaly [(W/m**2)/(mm/day)]

-8 -7 -6 -5 -4 -3 -2 -0.5 0 0.5 2 3 4 5 6 7 8

ERAI

CCSM3   CCSM4   SPCCSM3

CAM3_5d   CAM4_5d   SPCAM3_5d

CAM3_mon   CAM4_mon   SPCAM3_mon

rainfall [0.5 (mm/day)/(mm/day)]   moisture convergence [1E-6 (g/kg/s)/(mm/day)]

AGCM

CGCM

a weak ISO

unorganized convection and circulation

b coupling

warm SST anomaly organizes convection, circulation. increases LHFLX and q-advection east of convection.

c better ISO

organized convection and circulation induces LHFLX and q-advection east of convection

d coupling

IO SST anomalies enhance convection & wind anomalies, increasing LHFLX and q-advection east of convection.

plan view of equatorial convection

regional and seasonal local SST sensitivity

DJF   JJA

a) IO DJF   b) WPac DJF   c) IO JJA   d) WPac JJA

lh

d) IO DJF   e) WPac DJF a)   f) IO JJA   g) WPac JJA

sh

regional and seasonal local SST sensitivity

- shflx is more sensitive to SST than lhflx.
- greater sensitivity to SST in IO than WPac.
- SST sensitivity is greatest in DJF.
- IMPORTANT: these results are for 10N-10S only!

ERAI

moistening rate

a) $\sigma(\text{rain}) = 3.19 \text{ mm/day}$

moistening by vertical advection

b)

moistening by horizontal advection

c)

residual

d)

Surface

e)   lh   sh   sst   ps

longitude

mean q[g/kg] (departure from zonal mean)

-1.0 -0.8 -0.6 -0.4 -0.2 0.2 0.4 0.6 0.8 1.0
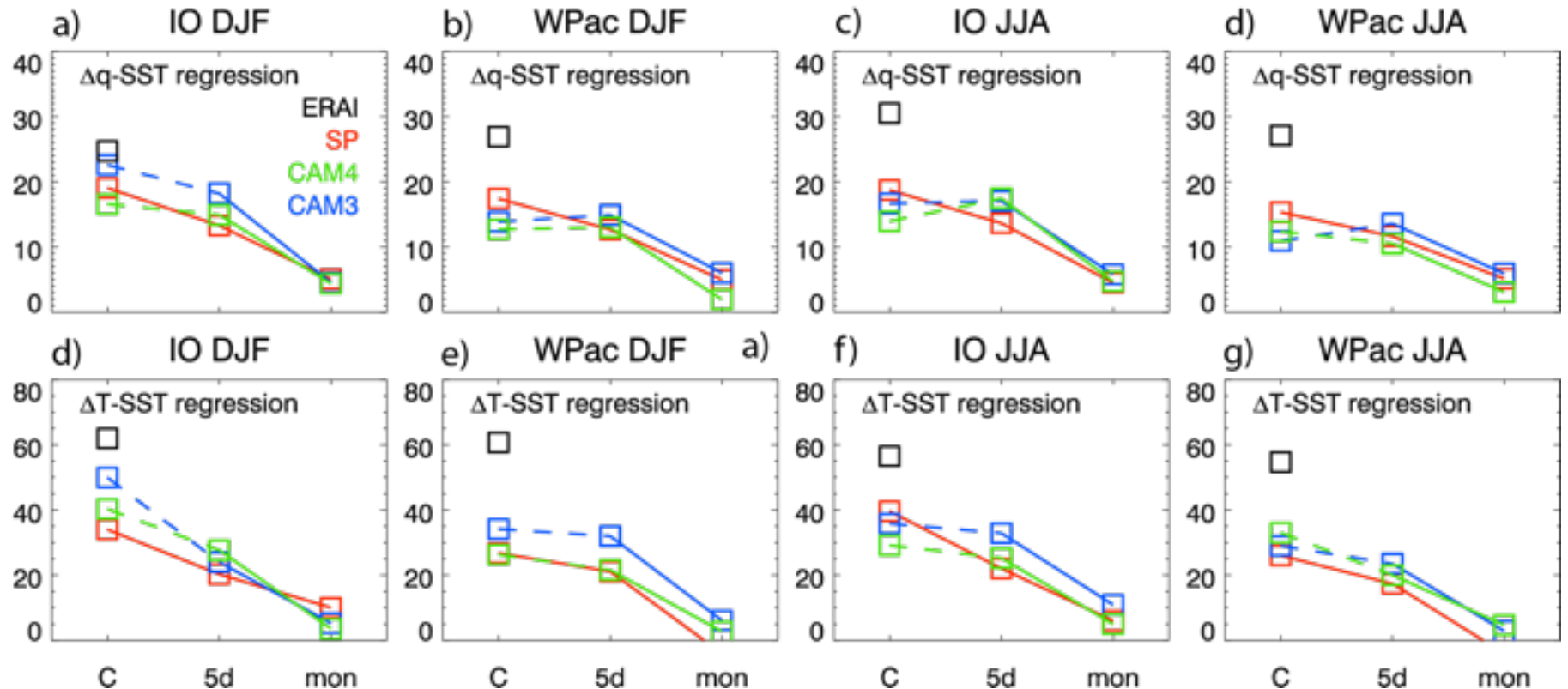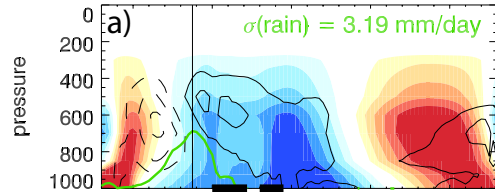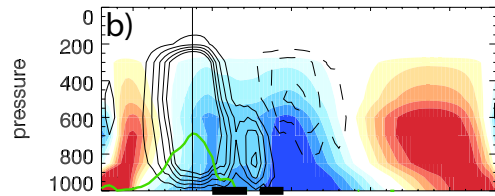
SPCCSM3   SPCAM3_5d   SPCAM3_mon

moistening rate

a) σ(rain) = 2.13 mm/day
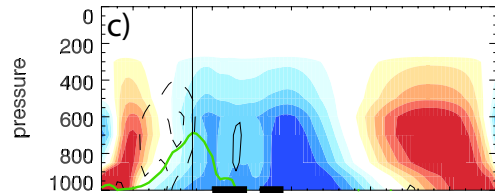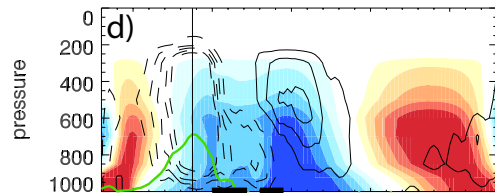f) σ(rain) = 2.04 mm/day
k) σ(rain) = 2.05 mm/day

moistening by vertical advection

b)   g)   l)

moistening by horizontal advection

c)   h)   m)

residual

d)   i)   n)

Surface

e)   j)   o)

lh   sh   sst   ps
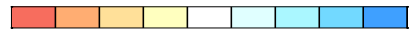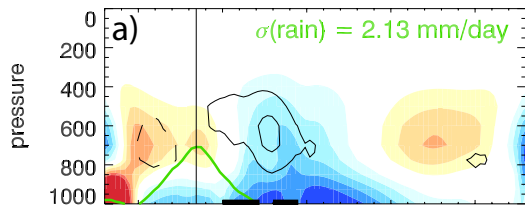
mean q[g/kg] (departure from zonal mean)

-1.0   -0.8   -0.6   -0.4   -0.2   0.2   0.4   0.6   0.8   1.0

CCSM4

moistening rate

a) σ(rain) = 1.43 mm/day

moistening by vertical advection

b)

moistening by horizontal advection

c)

residual

d)

Surface

e)   lh   sh   sst   ps

CAM4_5d

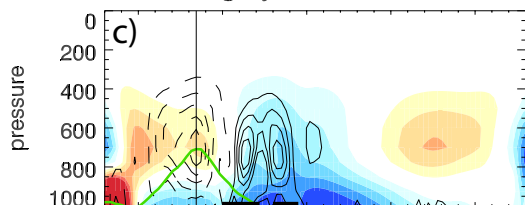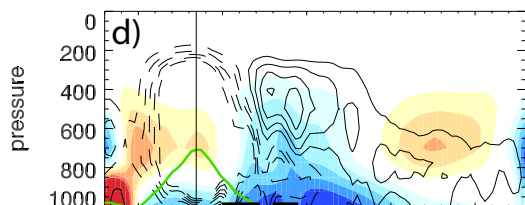moistening rate

f) σ(rain) = 1.54 mm/day

moistening by vertical advection

g)

moistening by horizontal advection

h)

residual

i)

Surface

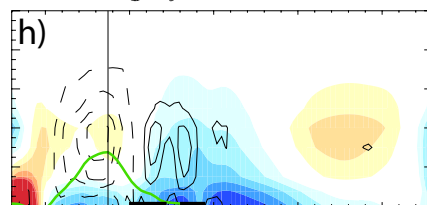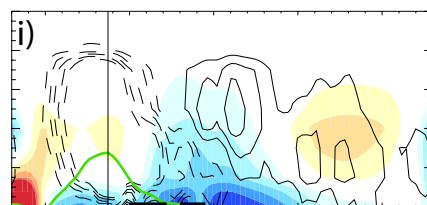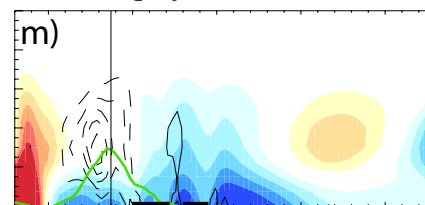j)   lh   sh   sst   ps

CAM4_mon

moistening rate
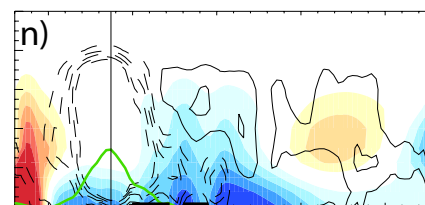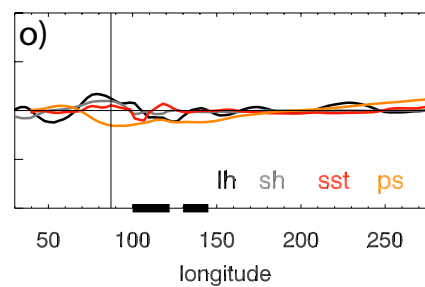
k) σ(rain) = 1.48 mm/day

moistening by vertical advection

l)

moistening by horizontal advection
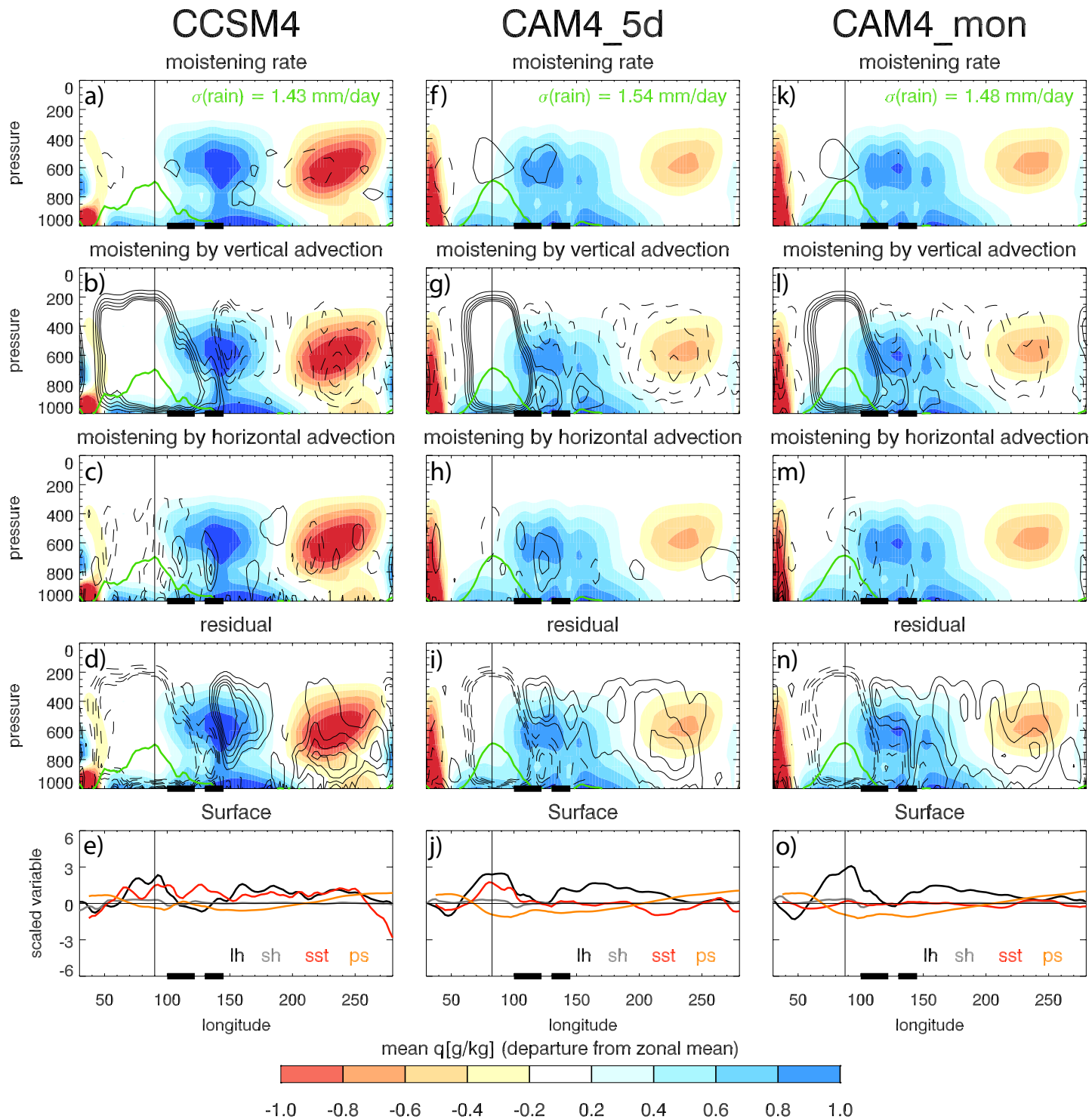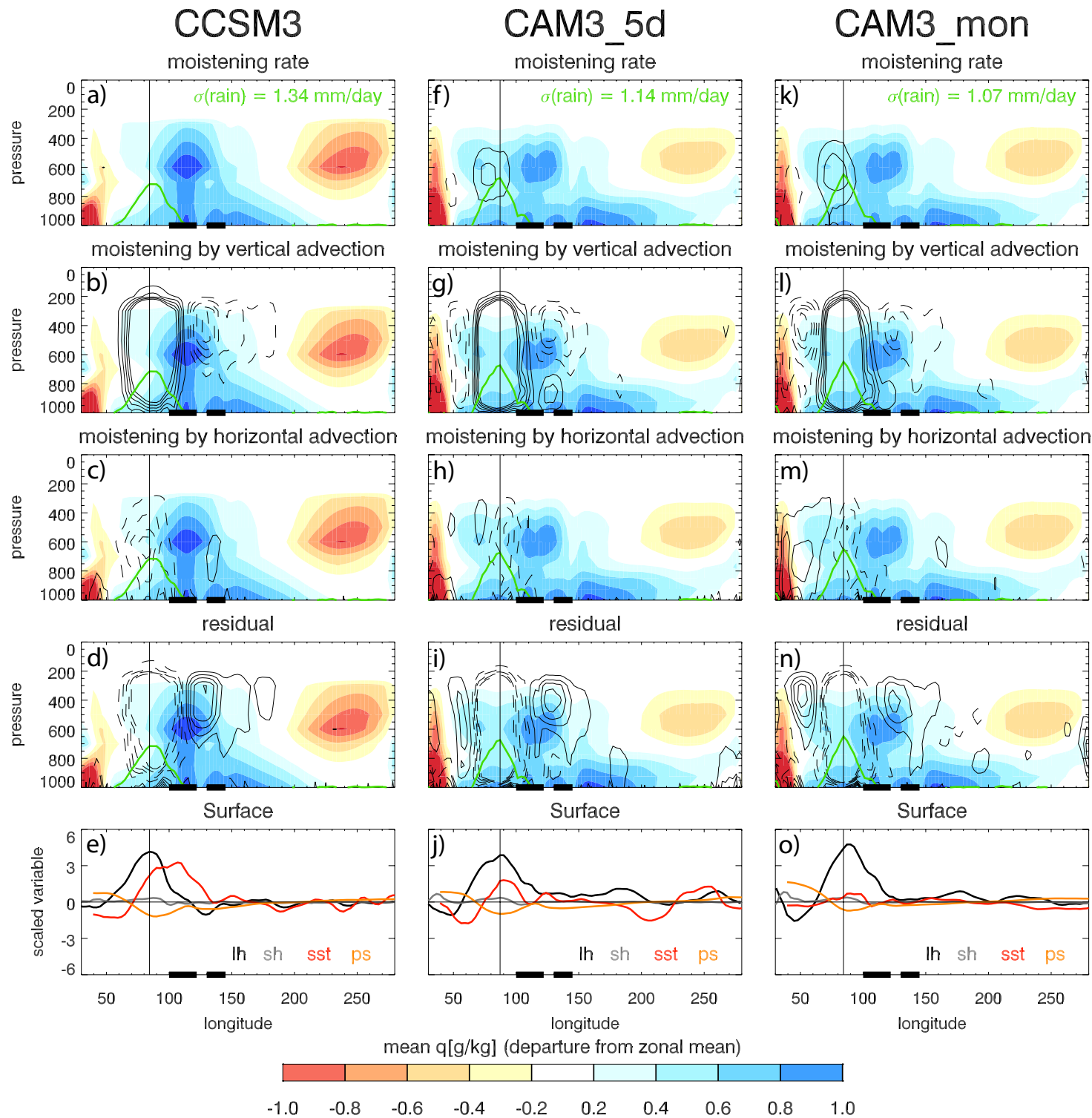
m)

residual

n)

Surface

o)   lh   sh   sst   ps
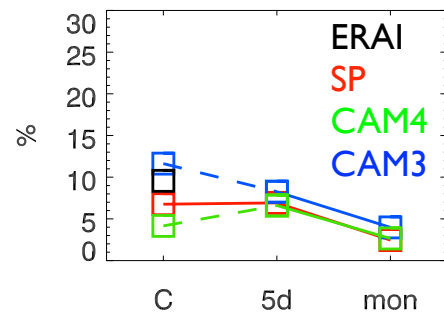
pressure

scaled variable

longitude

mean q[g/kg] (departure from zonal mean)

-1.0   -0.8   -0.6   -0.4   -0.2   0.2   0.4   0.6   0.8   1.0

mean q[g/kg] (departure from zonal mean)