

A satellite photograph of a tropical cyclone, showing a large, circular cloud system with a distinct eye in the center. The cyclone is set against the backdrop of the Earth's surface, which includes landmasses and oceans. The image is taken from a high altitude, providing a top-down view of the storm's structure.

Buoyancy in Tropical Cyclones

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source: NASA Space Photo Gallery

Importance of Eyewall convection

Pendergrass and Willoughby (2009):

- **Timing** and **Location** of repeated **convective bursts** seems critical in intensification efficiency
- Convective heating **inside the RMW** most efficient

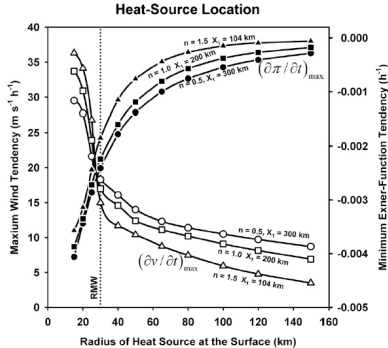
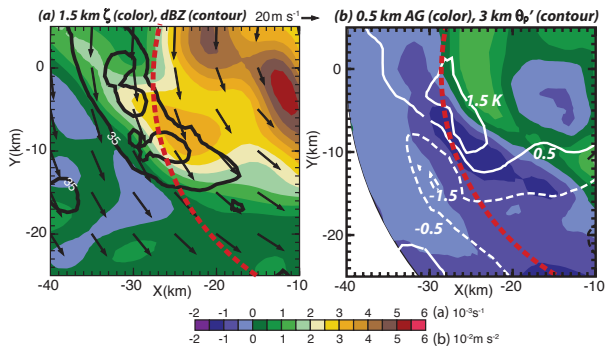


FIG. 12. Maximum rates to change of Exner function and swirling as a function of heat source radius for sharp ($n = 1.5$ and $X_1 = 104 \text{ km}$), medium ($n = 1.0$ and $X_1 = 200 \text{ km}$), and broad ($n = 0.5$ and $X_1 = 300 \text{ km}$) vortex wind profiles.

Possible Internal Forcing Mechanisms

- Boundary layer dynamics (Shannon McElhinney)
- Vorticity anomalies
- **Buoyancy**

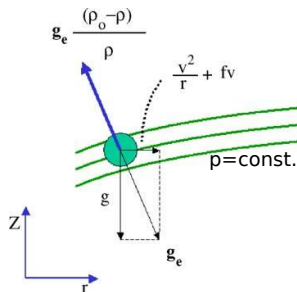


Buoyancy in a Tropical Cyclone

- Archimedes principle
 - Body lighter than environment
→ upward acceleration
 - $B = -g \frac{\rho'}{\rho_0}$
- More complicated in TC
 - Pressure not horizontally uniform
 - Choice of reference state critical for interpretation (see Zhang et al. (2000), Braun (2002) and Eastin et al. (2005))

$$\frac{Dw}{Dt} = -\frac{1}{\rho} \frac{\partial p}{\partial z} - g = -\frac{1}{\rho} \frac{\partial p'}{\partial z} + b$$

$$\text{with } p = p_0 + p'$$



Source: Smith et al. (2005)

Modelling studies

- **Zhang et al. (2000):**
 - MM5 simulation of Hurricane Andrew (1992) ($\Delta x = 6$ km)
 - Reference state: Running average over four neighboring points
- **Braun (2002):**
 - MM5 simulation of Hurricane Bob (1991) ($\Delta x = 1.3$ km)
 - Reference state: Wavenumber 0 and 1 of the Fourier decomposition (balanced vortex)
- **Comparison:**
 - Positive vertical accelerations (two large opposing terms)
 - Difference: **positively** buoyant (Braun), **negatively** buoyant (Zhang et al.)

Observation-derived Buoyancy

Ideally: High resolution thermodynamic and dynamic measurements with good spatial coverage

Flightlevel data (Eastin et al. 2005):

- Direct observations
- Limited spatial coverage
- Definition of reference state difficult
- Positive buoyancy

Radar data:

- Great spatial coverage
- No thermodynamic measurements
- Utilize radar data to estimate thermodynamic state of the atmosphere (following Gal-Chen (1978), Viltard and Roux (1998) and Liou (2001))

Thermodynamic Retrieval

- **Cylindrical** coordinate system
- Reference state is in **hydrostatic** and **gradient wind** balance

$$\left[\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial r} + v \frac{\partial u}{r \partial \lambda} + (w + q_r W) \frac{\partial u}{\partial z} - F_u - 2 \frac{\bar{v} v'}{r} - \frac{v'^2}{r} - f v' \right] = -c_p \bar{\theta}_\rho \frac{\partial \bar{\pi}}{\partial r} \left(\frac{\theta'_\rho}{\bar{\theta}_\rho} \right) - c_p \bar{\theta}_\rho \frac{\partial \pi'}{\partial r}$$

$$\left[\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial r} + v \frac{\partial v}{r \partial \lambda} + (w + q_r W) \frac{\partial v}{\partial z} + u \left(\frac{v}{r} + f \right) - F_v \right] = -c_p \bar{\theta}_\rho \frac{\partial \pi'}{r \partial \lambda}$$

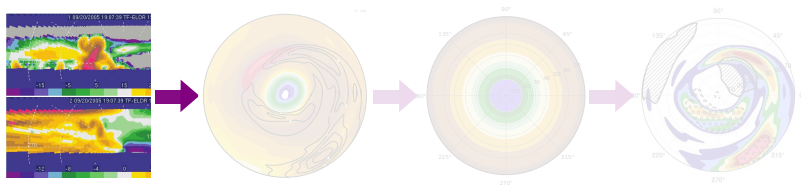
$$\left[\frac{\partial w}{\partial t} + u \frac{\partial w}{\partial r} + v \frac{\partial w}{r \partial \lambda} + (w + q_r W) \frac{\partial w}{\partial z} - F_w \right] = g \left(\frac{\theta'_\rho}{\bar{\theta}_\rho} \right) - c_p \bar{\theta}_\rho \frac{\partial \pi'}{\partial z}$$

$$\left[\frac{\partial \theta'}{\partial t} + u \frac{\partial \bar{\theta}}{\partial r} + v \frac{\partial \bar{\theta}}{r \partial \lambda} + w \frac{\partial \bar{\theta}}{\partial z} \right] = -u \frac{\partial \theta'}{\partial r} - v \frac{\partial \theta'}{r \partial \lambda} - w \frac{\partial \theta'}{\partial z} - F_\theta$$

- Using a variational approach to solve this set of equations

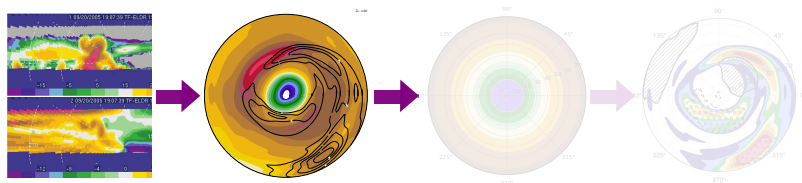
Retrieval Methodology

- 1) **Assimilate radar** and other observational data with variational approach
- 2) Calculate the **mean state** by integrating the thermal wind equation
- 3) Retrieve buoyancy and pressure perturbations with **thermodynamic retrieval**



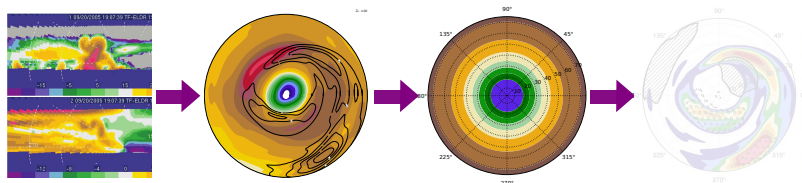
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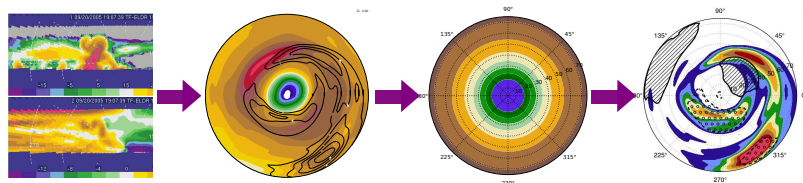
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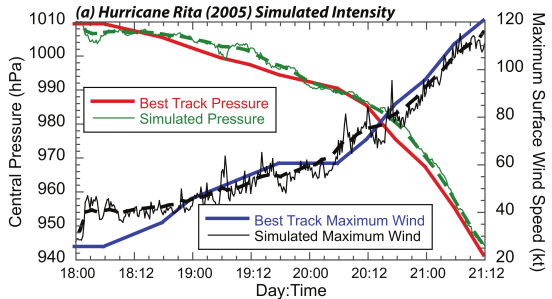
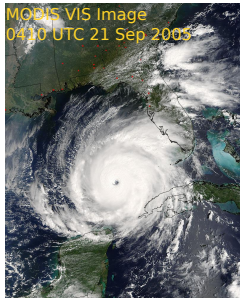
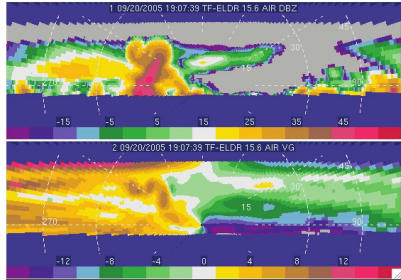
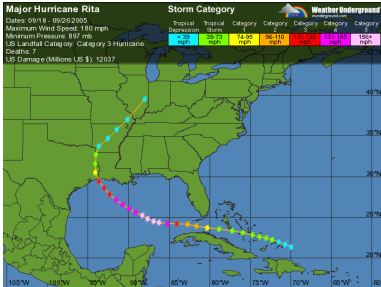


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RAINEX Hurricane Rita (2005)



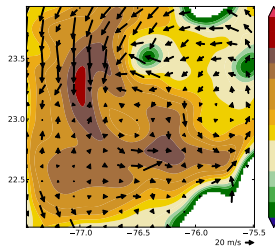
Step 1) Assimilation of Radar Data

SAMURAI (**S**pline **A**nalysis at **M**esoscale **U**tilizing **R**adar and **A**ircraft **I**nstrumentation)

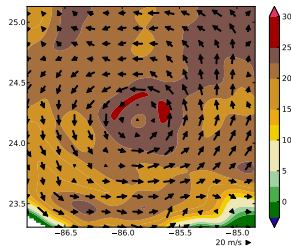
Bell et al. (2012)

- Maximum likelihood estimate by minimizing a cost function
- Galerkin approach (basis are cubic B-splines)
- Incorporating multiple data sources
- Analysis output on regular grid (Cartesian or cylindrical)

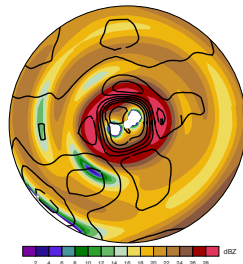
19 September 2005



21 September 2005



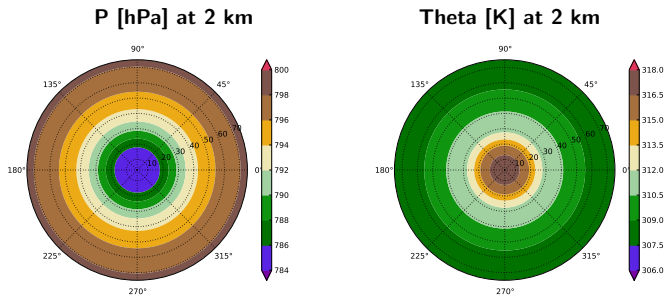
21 September 2005



Radar Reflectivity [dBZ] at 2 km

Step 2) Calculate Balanced Mean State

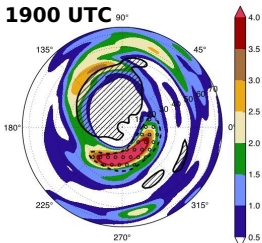
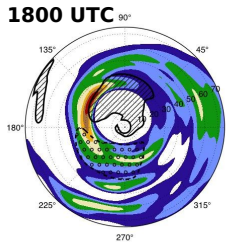
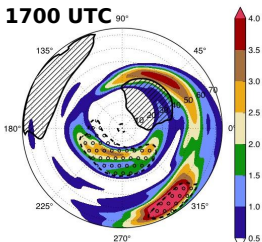
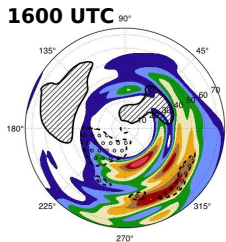
- Vortex in hydrostatic and gradient wind balance
- Start with sounding at the outer edge of the domain
- Integrate thermal wind balance inward along isobaric surfaces



- Retrieves pressure minimum and warm core
- Changes with radius and height

Step 3) Thermodynamic Retrieval (Synthetic Observations)

Precipitation mixing ratio [g/kg]

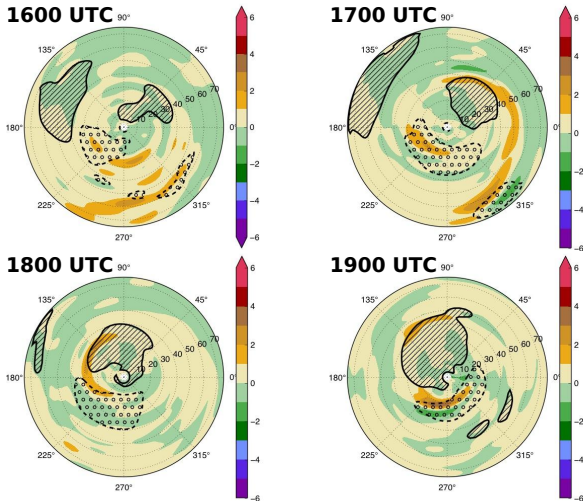


- Rainbands rotating cyclonically
- Buoyancy anomalies are associated with convection
- Mature convection shows waterloading

Buoyancy [m s^{-2}]: > 0.03 (hatched), < -0.03 (stippled)

Step 3) Thermodynamic Retrieval (Synthetic Observations)

Vertical motion [m/s]

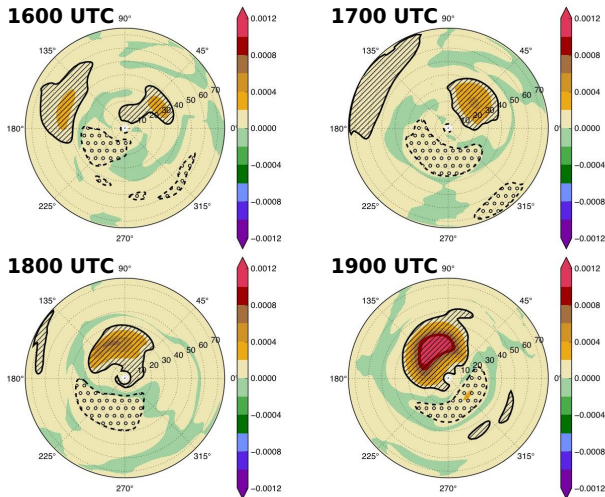


- Positive buoyancy anomalies are associated with upward motion
- Water loading is associated with downward motion

Buoyancy [m s^{-2}]: > 0.03 (hatched), < -0.03 (stippled)

Step 3) Thermodynamic Retrieval (Synthetic Observations)

Perturbation pressure gradient [m s^{-2}]



$$\frac{dw}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial z} - g$$

$$= -c_p \bar{\theta}_\rho \frac{\partial \pi'}{\partial z} + g \left(\frac{\theta'_\rho}{\bar{\theta}_\rho} \right)$$

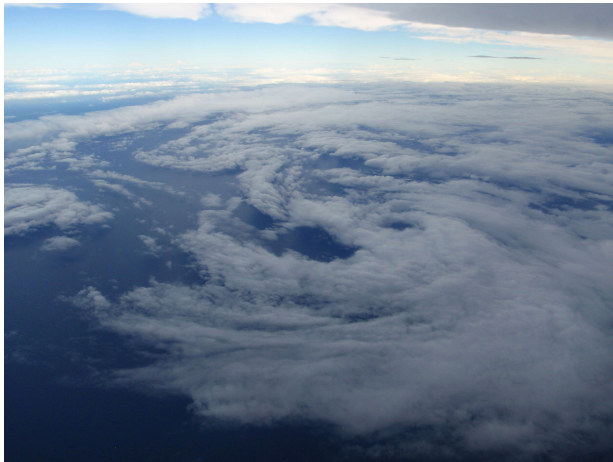
- Two large opposing terms

Buoyancy [m s^{-2}]: > 0.03 (hatched), < -0.03 (stippled)

Summary

- **Location** of bursts seems critical in **intensification efficiency**
- **Buoyancy** is one of the possible internal forcing mechanisms
- Thermodynamic retrieval of buoyancy from **radar data** using variational approach
- Preliminary results are promising
- Future work
 - Calculate buoyancy for different stages of the TC **lifecycle**
 - Compare results for **simulated** and **real** radar data
 - Compare rapid **intensifiers** with **steady state** storms
 - Investigate the interaction between **buoyancy**, **vorticity asymmetries** and **agradiant winds**

Thank you for your attention!
Questions?



Reference State

- **Zhang et al. (2000):**
 - MM5 simulation of Hurricane Andrew (1992)
 - Horizontal resolution $\Delta x = 6$ km
 - Running average over four neighboring points
- **Eastin et al. (2005):**
 - 1-Hz flight level data from 25 flights into 14 intense hurricanes
 - Running Bartlett filter with a 20-km window
- **Braun (2002):**
 - MM5 simulation of Hurricane Bob (1991)
 - Horizontal resolution $\Delta x = 1.3$ km
 - Wavenumber 0 and 1 of the Fourier decomposition (balanced vortex)